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TRANSMISSION LOSS IN RADIO PROPAGATION - II

BY KENNETH A. NORTON



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- (1) K. A. Norton, "Low and medium frequency radio propagation", Proc. of the International Congress on the Propagation of Radio Waves at Liege, Belgium, October, 1958, to be published by the Academic Press.
- (2) K. A. Norton, "System loss in radio wave propagation",
 J. Research, NBS, 63D, pp. 53-73, July-August, 1959.
- (3) K. A. Norton, "System loss in radio wave propagation", Letter to the Editor, Proc. I.R. E., to be published.

All of the material in reference (1) is included in this Technical Note. Some of the material in references (2) and (3) is new, particularly the definitions of the new terms "system loss" and "propagation loss". The transmission loss concept was adopted by the C.C.I.R. at its IXth Plenary Assembly in Los Angeles as is discussed more fully in reference (3) above.

TRANSMISSION LOSS IN RADIO PROPAGATION: II

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SUMMARY

In an earlier report with this title the concept of transmission loss was defined and its advantages explained. In this report a survey will be made of the transmission losses expected for a wide range of conditions, i.e., for distances from 10 to 10,000 statute miles; for radio frequencies from 10 kc to 100,000 Mc; for vertical or horizontal polarization; for ground waves, ionospheric waves, and tropospheric waves; over sea water or over land which may be either rough or smooth; and for various geographical and climatological regions.

Note: The attention of the reader is called to additional terms, discussed in appendix III, which must be added to the transmission losses shown in this report when the antennas are near the surface. These terms arise from changes in the antenna radiation resistances which occur when the antennas are near the surface, and represent important corrections to the transmission loss, particularly at the lower frequencies where the antennas, assumed to be 30 feet above the surface for many of the calculations, are only very small fractions of a wavelength above the surface.



by

Kenneth A. Norton

1. Transmission Loss in Radio Propagation

We will be concerned primarily with the transmission loss encountered in the propagation of radio energy between a transmitting and a receiving antenna. Simple methods will be given for determining the magnitude of this transmission loss and its variation in space and time (fading) for any frequency in the presently-used portion of the radio spectrum and for any kind of transmission path likely to be encountered in practice. In addition, methods will be given for estimating radio noise and interference levels. When combined, these two methods make possible the estimation of the transmitter power and antenna gain required for satisfactory communication, navigation, or other specific uses of the transmissions.

The transmission loss in a radio system involving propagation between antennas is simply the ratio of the radio frequency power, p_r , radiated from the transmitting antenna divided by the resulting radio frequency power, p_a , available from an equivalent loss-free receiving antenna; thus the system transmission loss = (p_r/p_a) . We see that the transmission loss of a system is a dimensionless number greater than unity, and that it will often be convenient to express this in decibels; the transmission loss, L, expressed in decibels, is thus always positive:

$$L = 10 \log_{10} (p_r/p_a) = P_r - P_a$$
 (1)

^{*} See references 1, 2 and 3.

[†] Throughout this report capital letters will be used to denote the ratios, expressed in decibels, of the corresponding quantities designated with lower-case type; e.g., P_r = 10 log₁₀P_r.

This particular choice of definition excludes from the transmission loss the transmitting and receiving antenna circuit losses and any loss which occurs in any transmission lines which may be used between the transmitter and the transmitting antenna or between the receiving antenna and the receiver. This exclusion has the advantage that it results in a measure of loss which is attributable solely to the transmission medium including the path antenna gain, Gp, which arises from the directivities of the transmitting and receiving antennas. In addition to the actual transmission loss, L, of the system, it is also convenient to define the basic transmission loss, Lb, to be the transmission loss expected if the actual antennas were replaced by isotropic antennas; this also serves to define the path antenna gain:

$$G_{p} \equiv L_{b} - L \tag{2}$$

Consider first an idealized isotropic transmitting antenna in free space radiating a power, p_r , expressed in watts. Such an antenna produces a field intensity of $p_r/4\pi d^2$ watts per square mile at a distance d expressed in miles provided $d > \lambda$. The absorbing area of a perfectly conducting, isotropic receiving antenna in free space is equal to $\lambda^2/4\pi$ where λ is the free-space wavelength expressed in miles; the resulting radio frequency power available from such a receiving antenna when placed at a distance $d > \lambda$ from the isotropic transmitting antenna is thus $p_a = p_r(\lambda/4\pi d)^2$. Thus we find

^{*} Antenna circuit loss includes the ground losses arising from the induction field of the antenna, but excludes losses occurring in the radiation field.

[†] In some of the past literature on radio wave propagation, the intensities of the expected fields have been given in terms of E, the field strength expressed in decibels above one microvolt per meter for one kilowatt effective power radiated from a half-wave dipole. It can be shown that L_b and E are simply and precisely related by $L_b = 139.367 + 20 \log_{10} f_{mc}$ - E.

that the basic transmission loss, $L_{\rm bf}$, for isotropic antennas in free space * is given by:

$$L_{bf} = 10 \log_{10}(p_r/p_a) = 10 \log_{10}(4\pi d/\lambda)^2 = 36.58 + 20 \log_{10}d + 20 \log_{10}f_{Mc}$$
(3)

In the above f denotes the radio frequency expressed in megacycles. Fig.1 shows this basic transmission loss for isotropic antennas in free space. For $d = 2\lambda$, $L_{bf} = 28$ db and thus (3) is only approximate when the indicated values of L_{bf} are less than, say, 30 db.

In fact, whenever the calculated transmission loss is less than, say 30 db, we must consider that the problem involves a transfer of an appreciable portion of the power between the transmitting and receiving antennas by other than radiation. For example, there is a direct coupling between the antennas via their induction and electrostatic fields, and this is a negligible factor in the calculation of the transmission loss only when L > 30 db. When high gain antennas are used, their separation must be much greater than 2λ in order to maintain the condition L > 30 db.

For an actual radio transmission system there will always be some path antenna gain so that the transmission loss $L = L_b - G_p$ will be less than the basic transmission loss. In some systems the free space gains G_t and G_r of both the transmitting and receiving antennas, respectively, will be fully realized so that $G_p = G_t + G_r$. For example, with half-wave dipoles having a common equatorial plane and separated by a distance $d >> \lambda$ in free space $G_t = G_r = 2.15$ db so that $G_p = 4.30$ db; the transmission loss for such a system is thus just 4.3 db less than that given by (3) and shown on Fig.1. Similarly, electrically short dipoles have gains

^{*} In some of the past literature on radio wave propagation, the intensities of the expected fields have been given in terms of $A \equiv L_b$ - L_{bf} , the attenuation relative to that expected for propagation in free space; in the case of surface wave propagation with vertical polarization, the attenuation has usually been expressed relative to an inverse distance field which is twice the free space field and thus $A' = L_b - L_{bf} + 6.021$ in this case.

BASIC TRANSMISSION LOSS IN FREE SPACE ISOTROPIC ANTENNAS AT BOTH TERMINALS $L_{bf} = 36.58 + 20 \log_{10} D + 20 \log_{10} f_{Mc}$

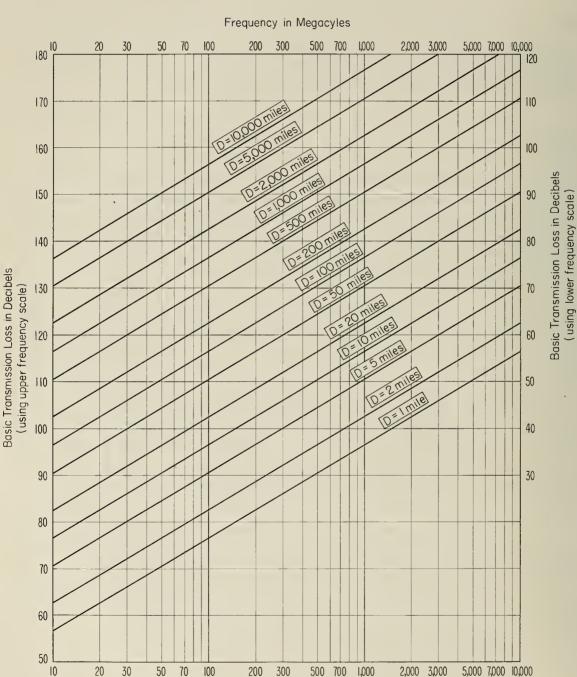


Figure I

Frequency in Kilocycles

 $G_t = G_r = 1.76$ db so that $G_p = 3.52$ db for propagation between appropriately oriented electrically short dipoles in free space.

The free space gain of a large receiving antenna with an effective absorbing area of a_e square meters will increase with increasing frequency at sufficiently high frequencies:

$$G_r = 10 \log_{10} a_e + 20 \log_{10} f_{Mc} - 38.54$$
 (4)

(For
$$f_{Mc} > 100/\sqrt{a_e}$$
)

For example, a large parabolic antenna will have an effective absorbing area a between 50 and 70 per cent of its actual area.

2. Transmission Loss in Free Space

Before considering the additional influences of the earth's surface and of its atmosphere on the propagation and transmission loss of the radio waves, it is instructive to consider first the characteristics of the transmission loss in free space for three kinds of systems which are typical of most of the applications encountered in practice.

Consider first a broadcast type of system in which essentially non-directional antennas are used at both terminals of the transmission path. For example, if half wave dipoles were used we have already seen that the system transmission loss, L, will be just 4.3 db less than that given by (3) and shown on Fig. 1. For such systems we see that the loss increases rapidly with increasing frequency because of the decreasing absorbing area of the receiving antenna. For this reason such systems should, in general, use the lowest available frequencies.

Consider next a type of broadcast service in which a directional array may be used at one end of the path: television is an example since the televiewers in remote areas consistently use high gain receiving antennas. If we assume that a half-wave dipole is used at the other terminal, we may combine (3) and (4) and obtain for the system transmission loss:

$$L_f = 72.97 + 20 \log_{10} d - 10 \log_{10} a_e$$
 (5)

Note that the free space transmission loss in this case is independent of frequency. In this case again, because of the additional loss arising from the effects of irregular terrain* which increase with increasing frequency, it is generally desirable to keep this kind of broadcasting service at the lowest available frequencies.

Finally consider a point-to-point type of service in which two identical high gain (and thus highly directional) antennas are used at each terminal of the transmission path. For such a system the free space transmission loss may be obtained from:

$$L_f = 113.67 + 20 \log_{10} d - 20 \log_{10} f_{Mc} - 20 \log_{10} a_e$$
 (6)

For services of this type it is clear that the highest frequencies free from the effects of atmospheric absorption are likely to be the most efficient. The above formula is applicable only to line-of-sight systems with first Fresnel-zone clearance over terrain which appears rough to the radio waves, and we will consider later within-line-ofsight smooth-terrain systems and beyond-the-horizon systems employing tropospheric scatter.

Rayleigh's criterion of the roughness may be used to determine whether a surface appears to the radio waves to be rough or smooth:

$$R = \frac{4\pi \sigma_h \sin \psi}{\lambda} \tag{7}$$

In the above equation σ_h denotes the standard deviation of the terrain heights relative to a smoothed mean height (see Fig. 2), $\psi = \psi_T = \psi_R$ denotes the grazing angle with the smoothed mean surface and λ is the wavelength expressed in the same units as σ_h . When R is less than 0.1, there will be a well defined specular reflection from the ground, but when R > 10, the reflected wave will be substantially weaker and will usually have a very small magnitude. †

The concept of first Fresnel-zone clearance provides a means of determining when the effects of the ground may be neglected so that the simple formula (6) may be used for determining the expected

^{*} See reference 2.

[†] See references 4, 5 and 6.

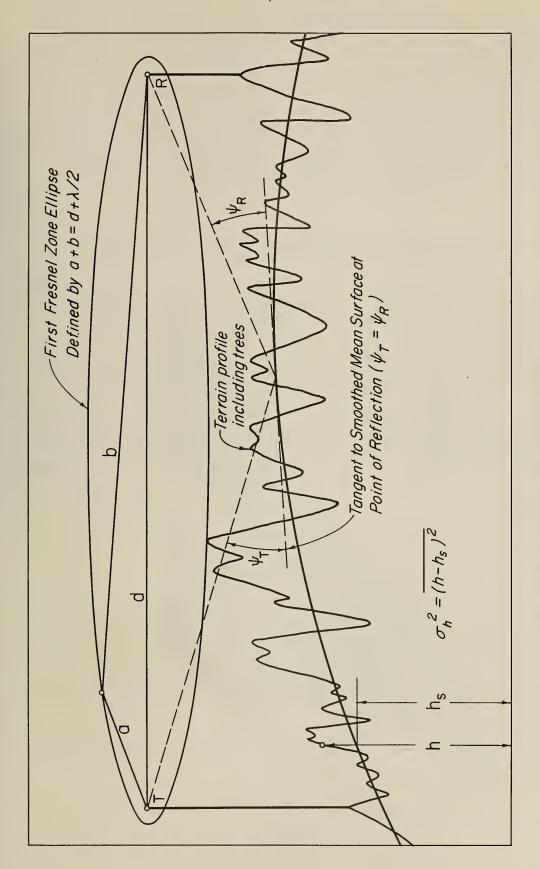


FIG. 2 PATH WITH FIRST FRESNEL ZONE CLEARANCE

transmission loss to a first approximation. Fig. 2 illustrates the first Fresnel zone concept. When the terrain along the path just touches the elliptical first Fresnel-zone defined by the locus of points such that $a+b=d+\lambda/2$, the path is considered to have first Fresnel-zone clearance for a system with wavelength λ . On Fig. 2, T and R represent the locations of the transmitting and receiving antennas. The presence of the ground will have only a small effect on the propagation, provided the antennas are sufficiently elevated so that none of the terrain lies within the first Fresnel zone and if, in addition, R > 10 so that the surface appears rough to the radio waves. *

3. Transmission Loss for Ground Wave Propagation

The ground wave is that component of the total received field which has not been reflected (or scattered) from either the ionosphere or the troposphere. It is convenient to divide the ground wave into two components: a space wave and a surface wave. **

The space wave is the sum of a direct wave and a ground-reflected wave. Figs. 3 and 4 give examples of space wave propagation. Near the radio horizon the ground-reflected wave is out of phase with the direct wave, and the received fields are quite weak; as the receiving antenna is raised, the relative phase increases until finally the direct and ground-reflected waves are in phase—at the lobe maxima shown on Figs. 3 and 4. At still higher heights the relative

^{*} See references 7 and 8.

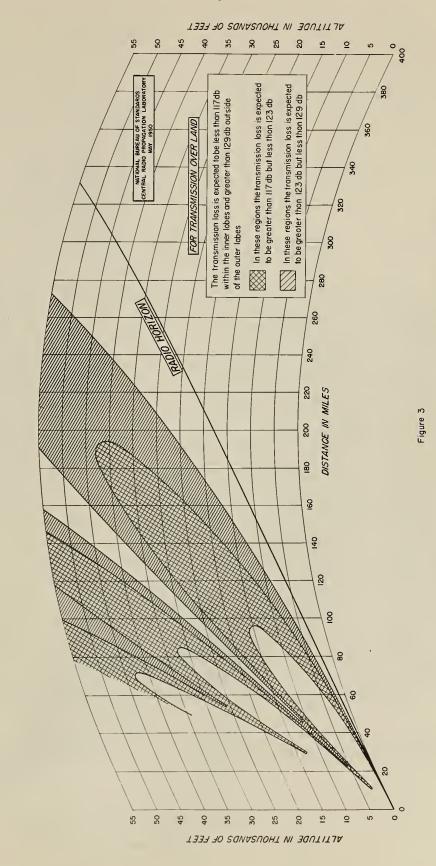
[†] See references 3, 9, 10, 11, 12, and 13.

^{**} The term Norton surface wave has been used in several recent papers in order to distinguish this component of the ground wave from the Zenneck surface wave with which it has sometimes been confused; the latter does not exist in practice as shown by Wise in reference 14. A recent discussion of surface waves by Wait in reference 15 further clarifies the physical nature of this and other surface wave components.

^{\$\}frac{1}{2}\$ See references 16 and 17 for a further discussion of air-to-ground propagation.

SPACE WAVE PROPAGATION BETWEEN VERTICAL HALF-WAVE DIPOLES OVER A SMOOTH SPHERICAL SURFACE

GROUND TERMINAL ANTENNA HEIGHT 35 FEET; FREQUENCY 328 MC



SPACE WAVE PROPAGATION BETWEEN VERTICAL HALF-WAVE DIPOLES OVER A SMOOTH SPHERICAL EARTH

GROUND TERMINAL ANTENNA HEIGHT 115 FEET; FREQUENCY 328 MC

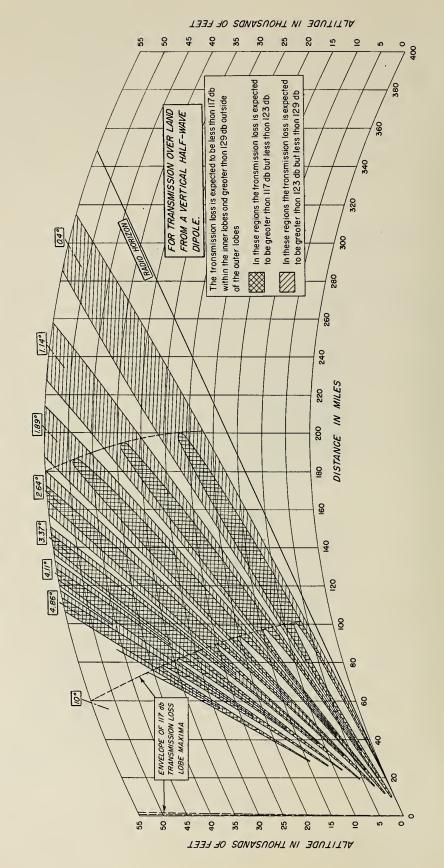


Figure 4

phase continues to increase, lobe minima and maxima occurring where the direct and ground reflected waves are out-of-phase and in-phase. Figs. 3 and 4 may be used to show approximately what happens at some other radio frequency, f_{Mc} , and ground antenna height, h, if we modify the transmission losses indicated on these figures by adding 20 log ($f_{Mc}/328$) and, at the same time, determine h for Fig. 3 by (h/35) = (328/ f_{Mc}) and for Fig. 4 by (h/115) = (328/ f_{Mc}).

Figs. 3 and 4 correspond to smooth earth conditions, i.e., for R < < 0.1. In this case the space wave field strength may be represented approximately by:

$$F = 2 F_0 \sin(2\pi h \sin \psi/\lambda)$$
 (8)

For example, the above equation represents very accurately the expected field strength for propagation from a horizontal dipole over a smooth, flat, perfectly conducting surface, where Fo is the field strength in free space, F is the expected field strength at a receiving point corresponding to a grazing angle ψ, and h is the height of the ground terminal antenna above the smooth surface. The maximum of the first lobe (F = 2F₀) occurs when $(2\pi h \sin \psi/\lambda) = \pi/2$; according to Rayleigh's criterion $(4\pi \sigma_h \sin \psi/\lambda)$ must be less than 0.1 for the surface to appear smooth to the radio waves. Combining these two results we find that σ_h must be less than $h/10\pi$, independent of the frequency, if we are to expect (8) to apply at the angle ψ corresponding to the maximum of the first lobe. At still higher angles the requirements for smoothness of the terrain are correspondingly more stringent. At lower grazing angles, however, the terrain may be correspondingly rougher; for example, at the angle below the first lobe maximum where $F = 0.2 F_0$ corresponding to a transmission loss 20 db greater than at the maximum of the lobe, σ_h must be less than 0.5h for the earth to be considered sufficiently smooth for (8) to apply and where F = 0.02 F_o corresponding to a transmission loss 40 db greater than at the lobe maxima, σ_h may be as large as 5h. Thus we see that the large reductions in the received field below the maximum of the first lobe as shown on Figs. 3 and 4 and indicated by (8) are expected to occur even over comparatively rough terrain.

For propagation conditions such that R is large, i.e., high frequencies, very rough terrain, or large grazing angles, the ground

reflected wave may be described statistically. It has been found that the Rayleigh distribution is appropriate for this purpose when R is very large, say R > 100, and that a combination of a constant specular component plus a random Rayleigh component is required for 0.01 < R < 100. Fig. 5 shows theoretical probability distributions for this case with the parameter K increasing from $(-\infty)$ for R<0.01 to values of K greater than 20 for R > 100. Here K is the level in decibels of the mean power in the random, Rayleigh distributed, component relative to that of the steady component. As an example of the use of probability distributions of this kind for describing space wave propagation conditions, suppose we have an air-to-air communication system operating at 328 Mc. As we fly over irregular terrain at a fixed high altitude away from another aircraft at the same altitude (See Fig. 6), the grazing angle ψ decreases from a comparatively large value to zero on the radio horizon, and this corresponds to a decrease of R from a very large value to zero on the radio horizon. Thus at short distances the ground-reflected wave will fluctuate in magnitude over a range indicated by the K = 20 curve on Fig. 5, while at larger ranges these fluctuations will occur over smaller and smaller ranges corresponding to the smaller values of K.

The above statistical description of space wave propagation over rough terrain is appropriate for propagation paths with Fresnel-zone clearance. For still smaller antenna heights involving propagation very near to or just below grazing incidence, the received space wave is log normally distributed. For example, a study ** of the fields received on over-land paths from television stations in the frequency range from 50 to 220 Mc and on receiving antennas with heights in the range from 12 to 30 feet indicates that the standard deviation of the received fields is of the order of 6 to 10 db about mean values of the order of magnitude expected for propagation over a smooth surface.

Finally, when the transmitting and receiving antennas are both actually on the surface, the received ground wave is a surface wave. Furthermore, when the transmitting and receiving antennas

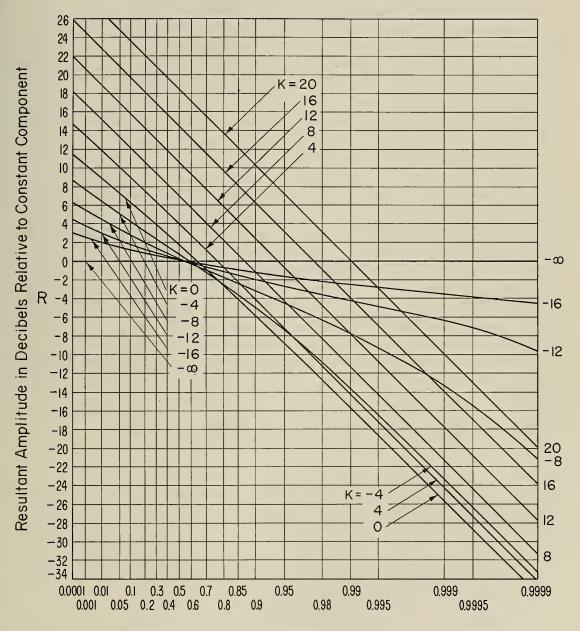
^{*} See references 4, 6 and 18.

[†] See references 2 and 19.

^{**} See reference 19.

DISTRIBUTION OF THE RESULTANT AMPLITUDE OF A CONSTANT VECTOR PLUS A RAYLEIGH DISTRIBUTED VECTOR

Power in Random Component is K Decibels Relative to Power in Constant Component



Probability that the Ordinate Value will be Exceeded

TRANSMISSION LOSS EXPECTED BETWEEN VERTICAL HALF WAVE DIPOLES IN AIR-TO-AIR PROPAGATION ON 328 MC BETWEEN TWO AIRCRAFT FLYING OVER IRREGULAR TERRAIN AT THE SAME ALTITUDE

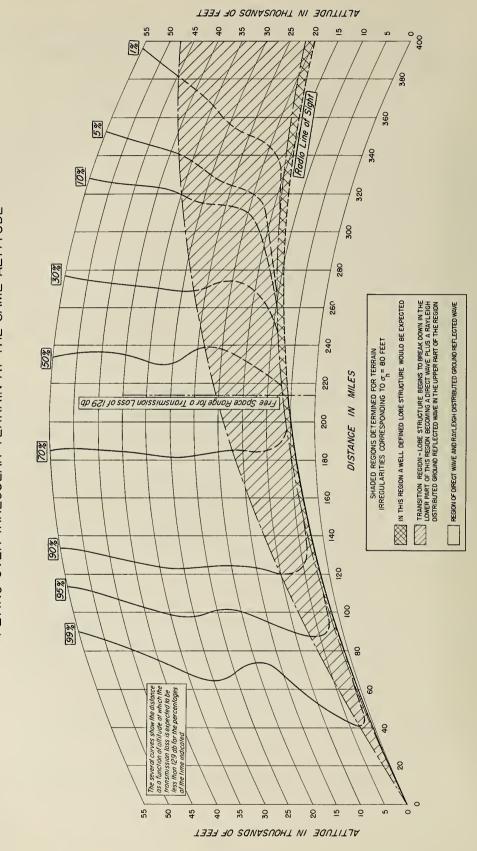


FIGURE 6

are both only a small fraction of a wavelength above the surface, the received ground wave is still primarily a surface wave together with a small space-wave component. The transmission loss in surface wave propagation* is very much influenced by the electrical constants of the ground, especially its conductivity. Although efforts have been made to correlate these ground constants with soil types so that predictions of the effective ground conductivity could be made, such studies have not been very successful so far. However, a publication of the National Bureau of Standards is available which gives the measured values of effective ground conductivity for various propagation paths in the United States. For propagation over average land one may use an effective ground conductivity of 5 milli-mhos per meter and an effective dielectric constant of 15 although individual over land paths may have substantially different ground constants. while over the sea the effective ground conductivity is of the order of 5 mhos per meter with an effective dielectric constant of 80.

Figs. 7, 8, 9, and 10 show the basic transmission loss expected for ground wave propagation over a smooth spherical earth with the transmitting and receiving antennas both at a height of 30 feet, for either vertical or horizontal polarization and with ground constants typical of over land and over sea water paths. At frequencies less than 10 Mc, the antenna heights are less than a wavelength and the ground waves shown for vertical polarization are primarily surface waves, whereas for frequencies greater than 100 Mc the ground waves with these antenna heights are primarily space waves with only a small surface wave component. Note that the proximity of the earth at low frequencies doubles the received fields for vertically polarized waves, but suppresses the propagation of horizontally polarized waves: i.e., horizontally polarized surface waves are highly attenuated. However, at the higher frequencies involving primarily space wave propagation, the expected transmission loss becomes independent of the polarization used.

On frequencies above 10,000 Mc the radio waves are appreciably absorbed by the oxygen and water vapor in the atmosphere. Fig. 11 shows the total gaseous atmospheric absorption near the surface at Washington, D.C. The absorption shown on Fig. 11 is the median value; for small percentages of the time the absorption will be considerably greater as a result of absorption

^{*} See references 9, 10, 11, and 12.

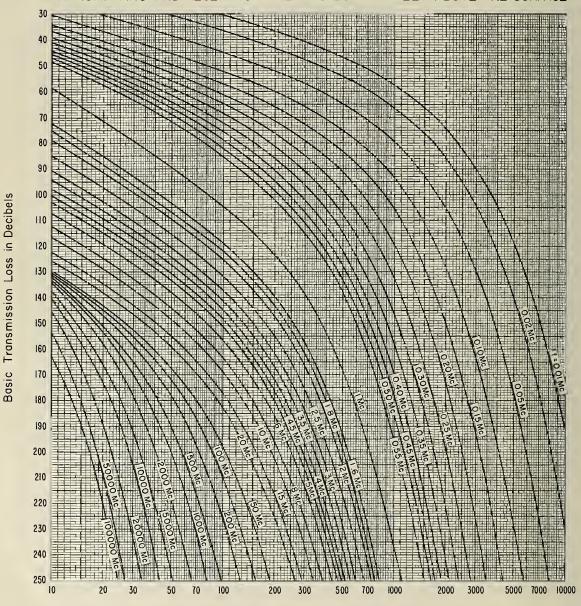
[†] Sce reference 20.

BASIC TRANSMISSION LOSS EXPECTED FOR GROUND WAVES PROPAGATED OVER A SMOOTH SPHERICAL EARTH

OVER LAND: $\sigma = 0.005$ MHOS/METER, $\epsilon = 15$

POLARIZATION: VERTICAL

TRANSMITTING AND RECEIVING ANTENNAS BOTH 30 FEET ABOVE THE SURFACE



Distance in Statute Miles

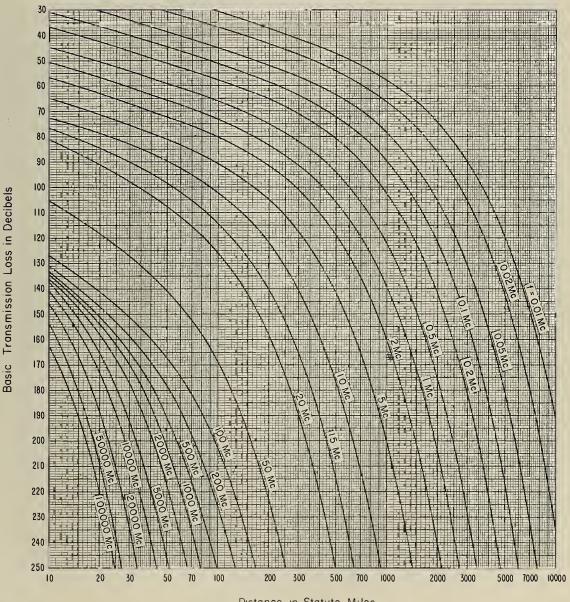
Figure 7

BASIC TRANSMISSION LOSS EXPECTED FOR GROUND WAVES PROPAGATED OVER A SMOOTH SPHERICAL EARTH

OVER SEAWATER: σ = 5 MHOS/METER, ϵ = 80

POLARIZATION: VERTICAL

TRANSMITTING AND RECEIVING ANTENNAS BOTH 30 FEET ABOVE THE SURFACE



Distance in Statute Miles

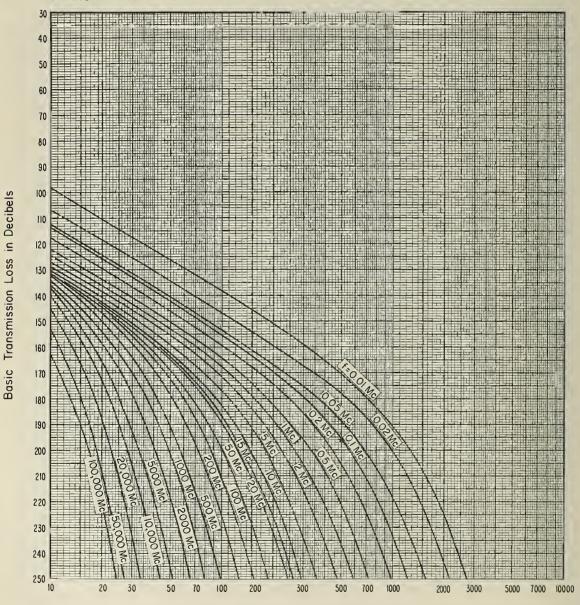
Figure 8

BASIC TRANSMISSION LOSS EXPECTED FOR GROUND WAVES PROPAGATED OVER A SMOOTH SPHERICAL EARTH

OVER LAND: σ =0.005 MHOS/METER, ϵ =15

POLARIZATION: HORIZONTAL

TRANSMITTING AND RECEIVING ANTENNAS BOTH 30 FEET ABOVE THE SURFACE



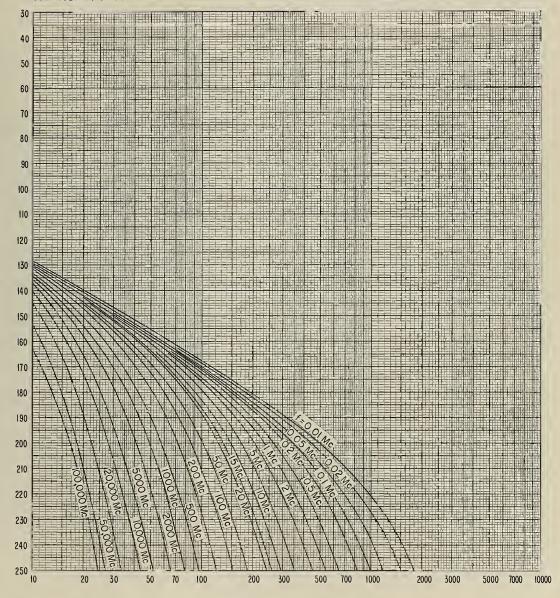
Distance in Statute Miles

Figure 9

BASIC TRANSMISSION LOSS EXPECTED FOR GROUND WAVES PROPAGATED OVER A SMOOTH SPHERICAL EARTH

OVER SEAWATER: σ = 5 MHOS/METER, ϵ = 80 POLARIZATION: HORIZONTAL

TRANSMITTING AND RECEIVING ANTENNAS BOTH 30 FEET ABOVE THE SURFACE

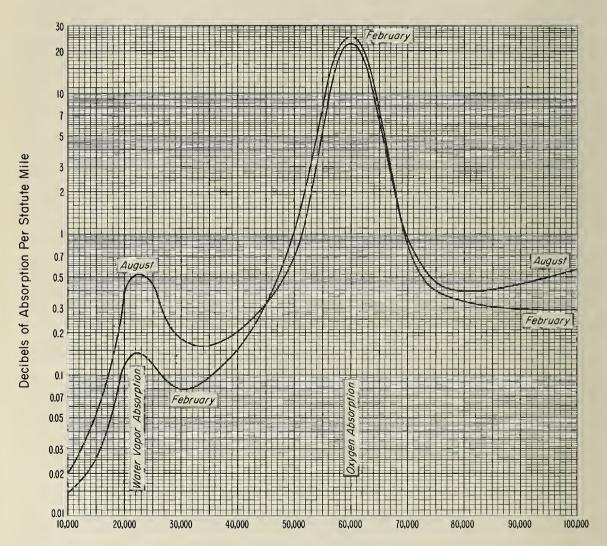


Basic Transmission Loss in Decibels

Distance in Statute Miles

Figure 10

AVERAGE GASEOUS ATMOSPHERIC ABSORPTION NEAR THE GROUND AT WASHINGTON, D.C. Meteorological Data From the Ratner Report July II, 1957



Frequency in Megacycles

Figure II

by rain. The transmission losses shown on Figs. 7, 8, 9, and 10 for frequencies above 10,000 Mc were estimated by using the August absorption shown on Fig. 11.

The heights of the antennas are very important in ground wave propagation at the higher frequencies. This is illustrated on Fig. 12 which shows for a frequency of 50 Mc the great reduction in transmission loss expected when the antenna height at one terminal is increased from zero up to 10,000 feet while the other antenna height is increased from zero up to 30 feet.

4. Transmission Loss for Ionospheric Propagation

Radio waves with frequencies less than the maximum usable frequency for a given transmission path are reflected by the ionized regions of the upper atmosphere with sufficient intensity so that they often provide a mode of transmission with less loss than that involved in ground wave propagation. The maximum frequency usable on a given ionospheric transmission path depends upon the length of the path, its geographical location, the time of day, the season of the year, and the phase of the sunspot cycle. Predictions of these maximum usable frequencies are published regularly three months in advance by the Central Radio Propagation Laboratory. † Fig. 13 is an example of these predictions showing how the maximum usable frequencies vary with local time and with geographical location for a path 4,000 kilometers long. This particular chart is for February, 1957, a period near sunspot maximum as may be seen on Fig. 14 which shows the smoothed Zurich sunspot numbers from 1750 to 1957. The sunspot numbers shown on Fig. 14 are averaged over a period of 13 months, but the author has shown in unpublished work that the sunspot numbers obtained by averaging over a period of three months are just as well correlated with ionospheric propagation conditions and thus provide a more useful index for prediction purposes. Fig. 15 shows a typical correlation between the observed maximum usable frequencies and these three-months-smoothed sunspot numbers.

^{*} See references 21, 22 and 23.

See references 24 and 25.

BASIC TRANSMISSION LOSS EXPECTED IN PROPAGATION OVER A SMOOTH SPHERICAL EARTH AT 50 MEGACYCLES

HORIZONTAL POLARIZATION; σ =0.005 MHOS/METER; ϵ = 15 FOR THE TERMINAL ANTENNA HEIGHTS INDICATED

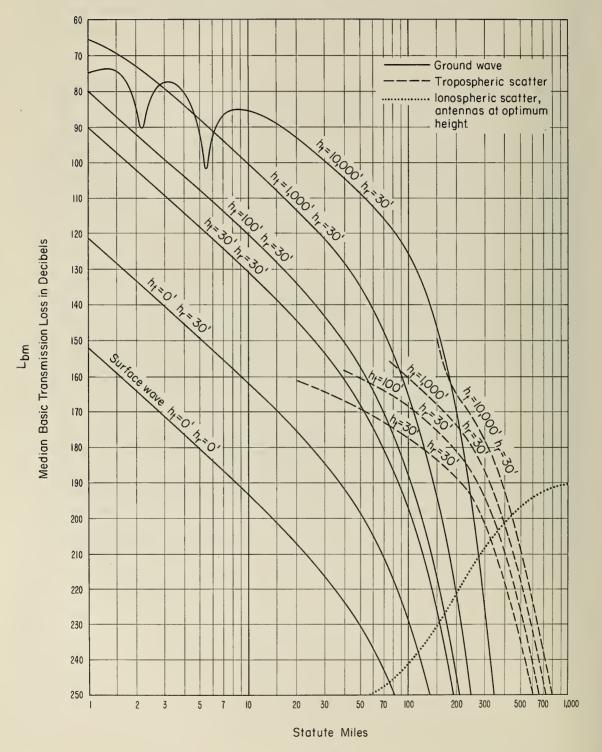
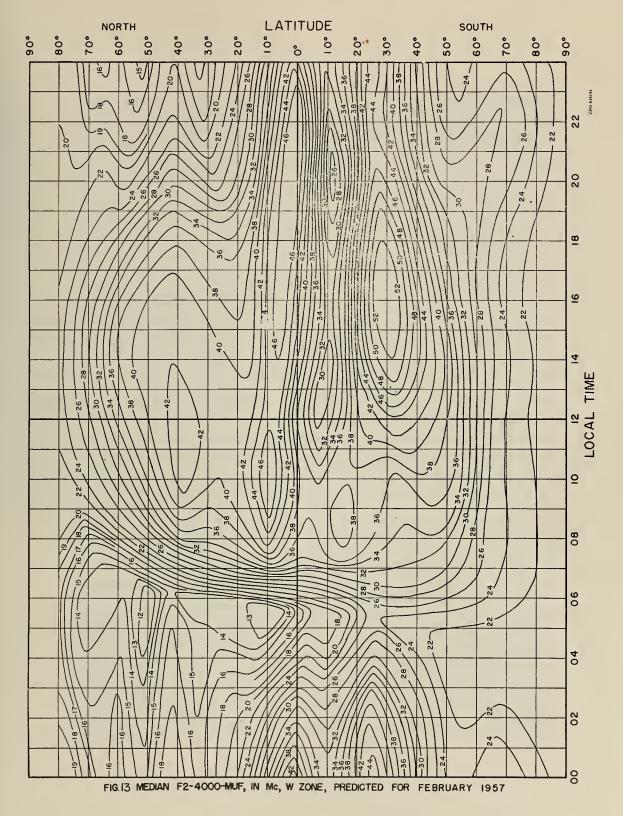


Figure 12



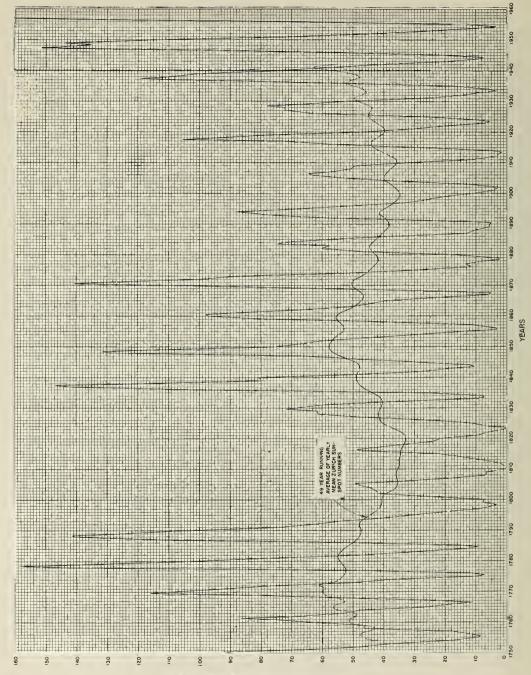
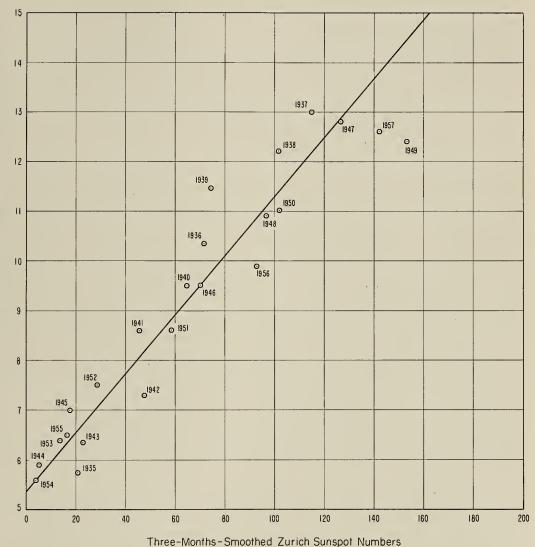


Figure 14

ZURICH AMONETISM AND ATMOSPHERIC ELECTRICITY" "LERRESTRIAL MY BRUNNER IN THE SEPTEMBER 1939 ISSUE OF "TERRESTRIAL MAGNETISM AND ATMOSPHERIC ELECTRICITY"

CORRELATION OF THE MONTHLY MEDIAN WASHINGTON 2.P.M. MAXIMUM USABLE FREQUENCIES FOR FEBRUARY WITH THE THREE-MONTHS-SMOOTHED ZURICH SUNSPOT NUMBERS



Monthly Median Maximum Usable Frequency at Vertical Incidence Expressed in Megacycles

•

Figure 15

4.1 Very Low Frequency Ionospheric Propagation

At the very low frequencies below 30 kc, the ionosphere reflects the waves at relatively low heights, about 70 km in the daytime and 90 km at night. At these low heights the ionization gradients are sufficiently large so that the ionosphere behaves as a sharp boundary, and it is convenient to use wave guide theory for determining the phase and amplitude of the received waves; good discussions of this theory are presented in the June, 1957 issue of the Proceedings of the Institute of Radio Engineers. Figures 16, 17, 18, and 19 give examples for this frequency range of the transmission loss expected in propagation between vertical electric dipoles over land and over sea and for day and night conditions. The values shown on these four figures were computed by the methods described by Wait; 26/ 27/ the minima and maxima shown are caused by interference between the ground wave and ionospheric wave modes at distances less than about 1,000 miles, and are caused by interference between the several ionospheric modes at the larger distances. At these long wavelengths the fading of the received waves is caused by a gradual shift from midday to midnight conditions and consequently has a very long period; thus at certain distances the received field may remain weak throughout the day or the night. The comparison between the calculated and observed locations and magnitudes of such anomalies provides a useful means for determining the effective constants of the ionosphere. The dimensionless constant L/H provides a measure of the effective conductivity of the ionospheric boundary, and the values of this constant assumed in these examples were determined by a comparison with observations of transmission loss. It is expected that L/H will also vary somewhat with the geomagnetic latitude of the receiving point, but such variations are not expected to have a large influence on the transmission loss.

It should be noted on Figs. 16 - 19 that values are shown for the transmission loss expected at distances beyond the antipode of the transmitter (about 12,500 miles), and at these larger distances a stronger signal would be expected from the shorter great circle path corresponding to transmission in the opposite direction; when short pulses are transmitted, these signals traveling in opposite directions will interfere with each other, and the results given on these figures should be useful in determining the magnitude of this multipath problem.

TRANSMISSION LOSS EXPECTED BETWEEN SHORT VERTICAL ELECTRIC DIPOLE ANTENNAS

Day Over Land $\sigma = 0.00477$ Mhos/Meter; lonospheric Constant L/H=0.1; h=70 km

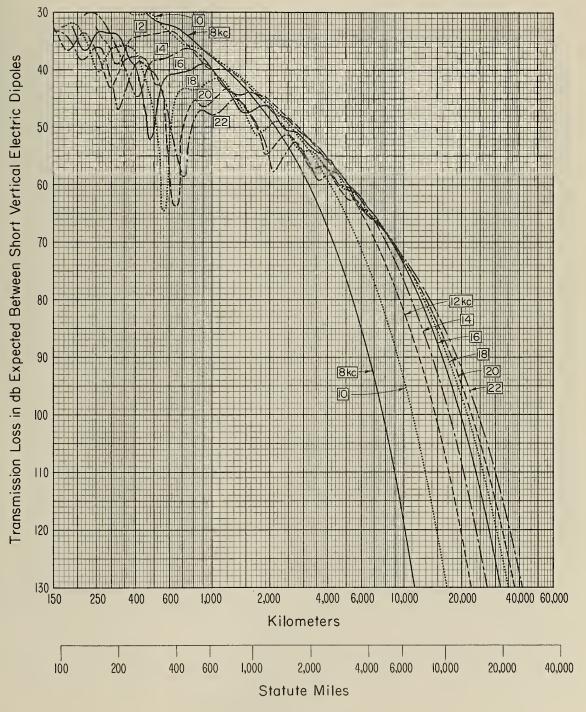


Figure 16

TRANSMISSION LOSS EXPECTED BETWEEN SHORT VERTICAL ELECTRIC DIPOLE ANTENNAS

Day Over Sea $\sigma = \infty$ lonospheric Constant L/H = 0.1; h = 70 km

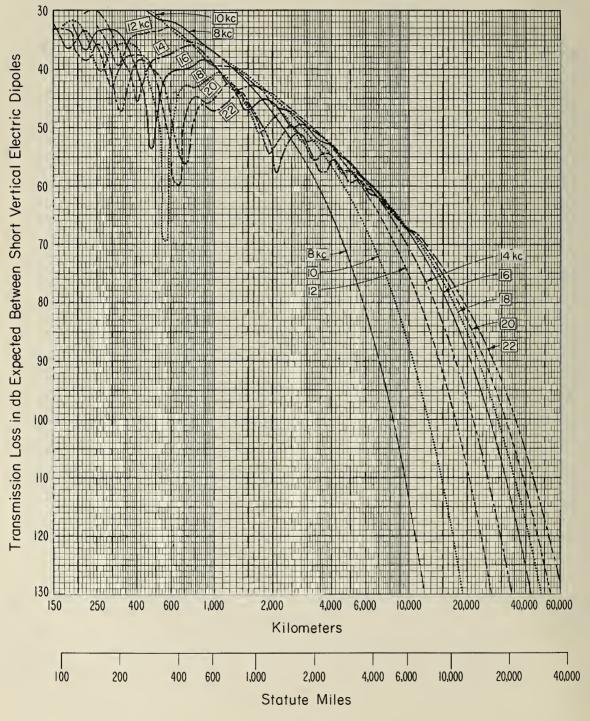


Figure 17

TRANSMISSION LOSS EXPECTED BETWEEN SHORT VERTICAL ELECTRIC DIPOLE ANTENNAS

Night Over Land $\sigma = 0.00371$ Mhos/Meter lonospheric Constant L/H=0.05; h=90 km

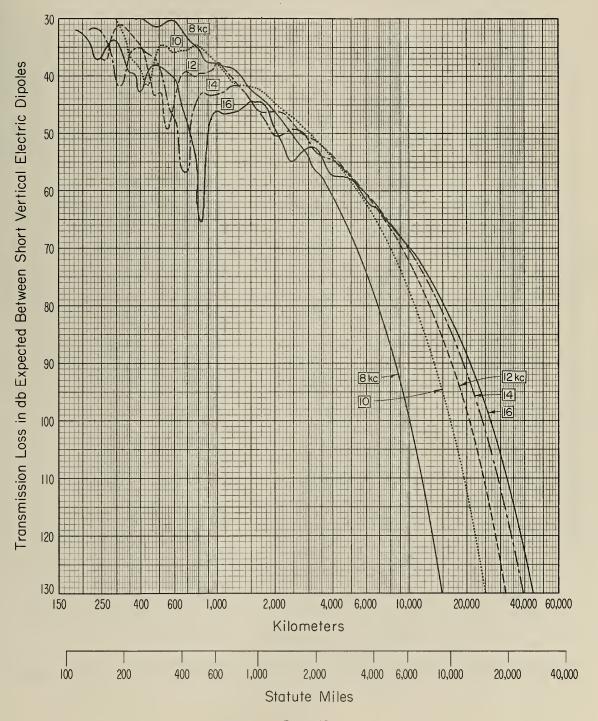


Figure 18

TRANSMISSION LOSS EXPECTED BETWEEN SHORT VERTICAL ELECTRIC DIPOLE ANTENNAS

Night Over Sea $\sigma = \infty$ Ionospheric Constant L/H = 0.05; h = 90 km

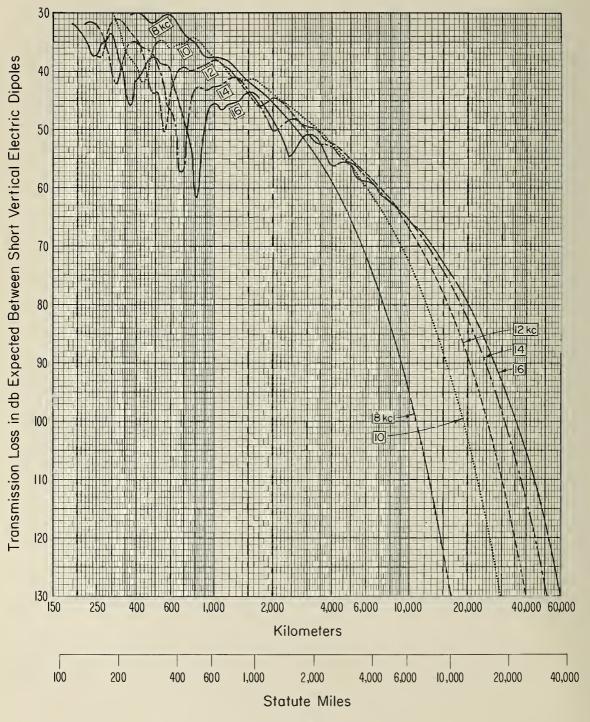


Figure 19

We will see in a later section how convergence at the curved surface of the ionosphere is expected to affect the transmission loss; in calculating the losses shown on Figs. 16-19, allowance was made only for the convergence expected in the vertical plane and on the assumption that the ionosphere is smooth. Allowance for convergence in the horizontal plane leads to large additional reductions in the transmission loss expected near the antipode-12,441 miles-and near 24,881 miles and 37,322 miles. For a smooth, spherical ionosphere, we should subtract the horizontal plane convergence C_h = 10 log₁₀[d(miles)/3960 sin{d(miles)/3960}] from the transmission losses shown; at distances, ρ , from the antipodal points less than 100 miles but greater than 0.1 λ , we may write C_h = 40.948 -10 log₁₀ ρ (miles); the value right at the antipodal points is given by C_h = 34.210 + 10 log₁₀ f_{kc} , but this latter includes both long and short great circle path energy.

The available experimental data indicate that the parameters of the ionosphere chosen for these calculations lead to about the right conclusions over the range of frequencies from 10 to 20 kc, but at 8 kc the calculated losses are somewhat greater than those observed.

Finally, it should be noted that mixtures of day and night and land and sea conditions are to be expected over these long paths, and suitable methods of calculation have yet to be developed for such mixed paths. It will likely be possible, however, to develop empirical methods for combining the results given here to obtain good estimates of the transmission loss expected on such mixed paths in much the same manner as has been used for estimating the transmission loss expected in ground wave propagation over mixed paths. 29/30/

4.2 Low and Medium Frequency Ionospheric Propagation

As we increase the radio frequency well above 30 kc, the ionosphere behaves much less as a sharp boundary and instead gradually refracts the waves back to the receiving point only after they have penetrated many kilometers into it, this penetration being greater the higher the radio frequency. The available evidence appears to indicate that the D and E regions of the ionosphere, which extend from 70 to 110 km, are turbulent, consisting of "blobs" of ionization which drift with the mean wind with velocities often in excess of 100 miles per hour. An interesting discussion of these irregular ionospheric motions is given in a recent article by Gautier. The radio waves will travel along many different paths through this turbulent ionized medium, the received field being the resultant vector sum of the waves received after propagation along these different paths. At sufficiently high frequencies, the relative phases of these waves will be random, and the resultant received field will have a Rayleigh distributed amplitude as shown on Fig. 5; on this figure K represents the ratio in decibels between the field intensity of the random ionospheric waves and a steady ground wave or, in the case of a single ionospheric mode, K represents the ratio in decibels between the field intensity of the random ionospheric waves and the steady, specularly-reflected component. Thus it becomes convenient, particularly for frequencies above 30 kc, to determine the transmission

loss separately for the ionospheric and ground wave modes of propagation. Figs. 20 and 21 show the transmission loss expected at 100 kc between short vertical electric dipole antennas for the ground wave and several ionospheric wave modes of propagation over land and over the sea, and for day and night conditions. The method of calculation used in determining the results shown on Figs. 20 and 21 involves a combination of ray and wave theory. Fig. 22 illustrates the geometry of our model and some of the assumptions made in the calculations. The waves are refracted in the troposphere down towards the earth and, as a consequence, the distance, d₁, traveled for a given ray angle of elevation, ψ , before the waves arrive back at the earth is substantially larger than if there were no atmosphere.

We have idealized our problem by assuming for all points along the path that the ionosphere has the same height, h, and the same reflection coefficient, while the ground is assumed to have the same electrical constants even for propagation all the way to the antipode at a distance of about 12,500 miles. The actual ionosphere and ground reflection conditions over particular propagation paths are obviously much different from these idealized paths, but our present model seems better for expository purposes. The principle of stationary phase (essentially the same as Fermat's principle) leads to the conclusion that the received waves may be considered to travel along several discrete ray paths between the transmitter and the receiver. All of these paths are great circle paths, the shortest corresponding to the ground wave mode of propagation. The other paths involve m reflections at the ionosphere, and the waves propagated along these other paths arrive at the receiving point at successively later times. By transmitting short pulses, it is possible to observe these several modes independently at a distant receiving point, and in this way their physical reality has been verified. The term "mode of propagation" here, and in the remainder of the ionospheric propagation discussions, refers to the waves propagated along one of these ray paths, and has a distinctly different meaning from the usage in the previous section where the modes of propagation were the wave guide modes which are simply the successive terms in a mathematical expression for the field. The use of short pulses to make possible the separate reception of each of these modes is a very useful device for radio navigation and, in this connection, the estimation of the time of arrival of the successive modes becomes of great practical importance. These time delays have been studied both theoretically 32/33/34/ and experimentally 35/36/37/38/ and the reader is referred to the references for information of this kind; here we will be primarily interested only in their transmission losses.

MEDIAN TRANSMISSION LOSS OVER LAND AT 100 kc

 σ = 0.005 Mhos/meter; ϵ = 15 Day h = 70 km; Night h = 90 km

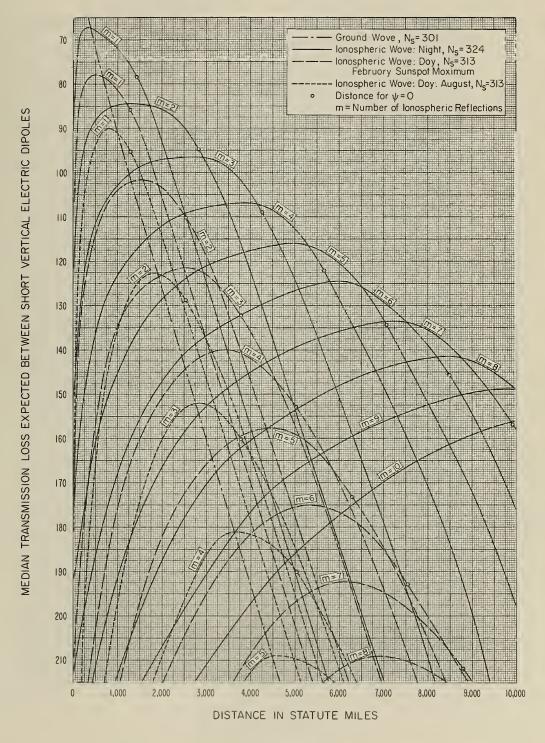


Figure 20

MEDIAN TRANSMISSION LOSS OVER SEA AT 100 kc

 σ = 5 Mhos/meter; ϵ = 80 Day'h = 70 km; Night h = 90 km

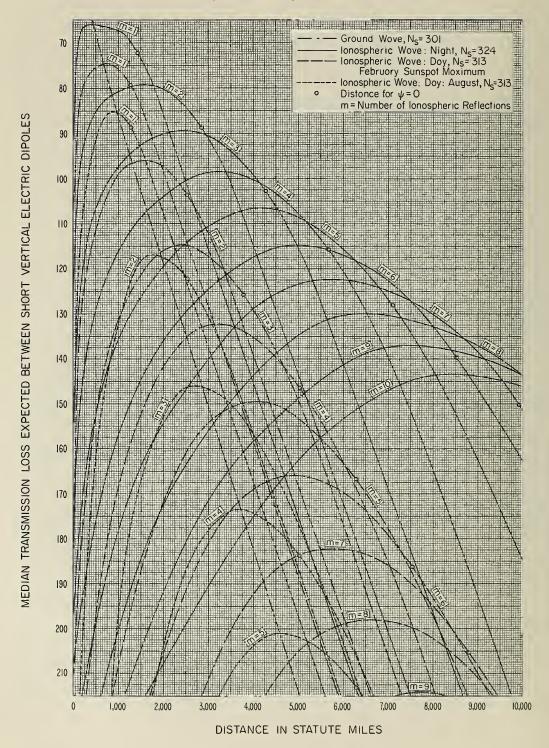
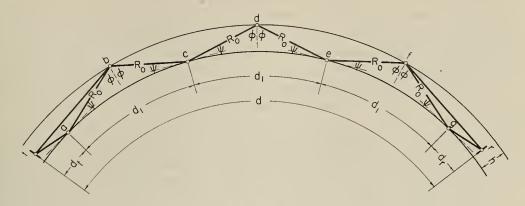
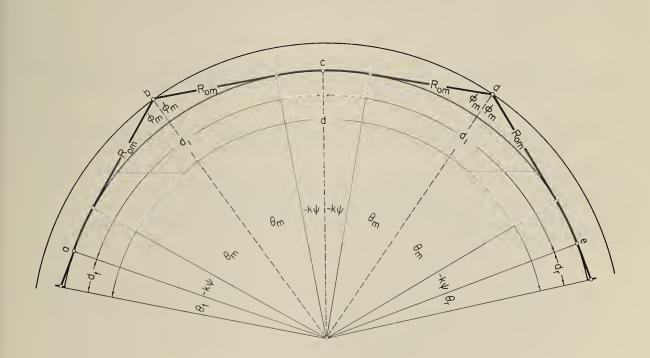


Figure 21



Case (a) ψ Positive and m = 3



Case (b) ψ Negative and m=2; k=I $-\frac{d\tau}{d\psi}$

Figure 22

Consider the phase of a radio wave for a given ray path, $\Omega(\psi) = 2\pi R/\lambda$, where λ is the wavelength and R is the total length of the ray path between the transmitter and the receiver. Now consider the variation of this phase for all possible adjacent paths between the transmitter and the receiver as we vary the angle of elevation, ψ , and azimuth, χ ; it should be clear (a) that $\Omega(\psi)$ will be a minimum with respect to variations in x for the set of rays lying in the great circle plane, i.e., for x = 0 and (b), of this set there will be m + 1 points of stationary phase, $\Omega'(\psi) = 0$, for the ray representing the ground wave and for the m rays reflected at the ionosphere in such a way that the angle of incidence, ϕ , at the ionosphere (for example, at b, d, and f on Fig. 22) is equal to the angle of reflection, and also (for m > 1) that the angle of incidence, $(90^{\circ} - \psi)$, at the ground (for example, at a, c, e, and g on Fig. 22) is equal to the angle of reflection at the ground. When the angle of elevation, ψ , is positive, the waves may be considered to travel both along the direct ray path, tb, from the transmitter to the ionosphere and along the ground-reflected ray path, tab; the reflection points b and f at the ionosphere and c and e at the ground will be very slightly different for the direct and ground-reflected ray paths, but this small difference is ignored in our calculations. that the angle of elevation, ψ , can be negative as is illustrated on Fig. 22, case (b).

The following formula may be used to calculate the median transmission loss of an ionospheric mode of propagation involving m reflections at the ionosphere and a ray path of length, R:

$$L_{m} = L_{bf}(R) + A_{t}(\psi) + A_{r}(\psi) + (m - 1) A_{g}(\psi) - C_{m}(R, 0.5) + P + mA(\phi, 0.5)$$
(9)

Each of the terms in the above is expressed in decibels; $L_{bf}(R)$ denotes the basic free space transmission loss (Set d = R in (3) in Section 1) for the ray distance, R. We see by Fig. 22 that we may calculate R as follows:

$$R \cong d_t + d_r + 2m R_0 \qquad (\psi > 0)$$
 (10)

$$R \cong d_t + d_r + 2m (R_{om} - ka\psi) \qquad (\psi \leq 0)$$
 (11)

$$d_{t} = \sqrt{(k a \tan \psi)^{2} + 2 k a h_{t} - k a \tan \psi}$$
 (12)

The distance, d_r , is determined by a formula similar to (12) with h_t replaced by h_r ; in this equation ka is the effective earth's radius, and k has been chosen equal to 4/3 in our ionospheric examples. Methods for estimating k as a function of time and geographical location are given in a later section. It is convenient to choose several values of ψ at conveniently spaced intervals and then to calculate all of the remaining factors at these particular values of ψ .

The space wave radiation factors, $A_t(\psi)$, and $A_r(\psi)$ include the gains of the transmitting and receiving antennas, respectively, relative to that of an isotropic antenna in free space, and allow for the radiation patterns of the antennas and the loss arising from the proximity of the antennas to the curved earth. The magnitude of $A_t(\psi)$ can be determined from $A_t(\psi) = L_1(\psi) - L_{bf}(R_0 + d_t)$, where $L_1(\psi)$ denotes the transmission loss expected for the ground wave mode propagated between the actual transmitting antenna and an isotropic receiving antenna placed at the first point of reflection in the ionosphere, while $L_{bf}(R_0 + d_t)$ is the corresponding basic free space transmission loss at this distance. Figs. 23 and 24 give typical values of $A_t(\psi)$ expected for short vertical electric dipoles 30 feet above the ground; in this case we may express $A_t(\psi)$ as follows:

$$A_t(\psi) = 20 \log_{10} |F| - 1.761 - 20 \log_{10} \cos \psi - 20 \log_{10} f(q)$$
 (13)

In the above |F| is a "cut-back" factor. When ψ is large and positive, |F| is just $|1+R|(\psi)|$ where $R|_V$ is the complex Fresnel reflection coefficient for plane vertically polarized waves incident on the ground at the grazing angle ψ ; when ψ is small or negative, the curvature of the earth becomes important and the values of |F| have then been determined by formulas recently developed by Wait. $\frac{39}{40}$ The term 1.761 is just the gain of the short dipole; the term 20 $\log_{10} \cos \psi$ allows for the cosine pattern of the dipole; and finally f(q) is the height gain factor given by equation (19) in reference (11) which allows for the effect of the height, $h|_V$ of the antenna above the surface. The "cut-back" factor |F| was calculated for a spherical surface of radius, $a|_V$ with k=4/3; this provides approximately for the effect of air refraction.

The factor $(m-1)A_g(\psi)$ allows for loss on reflection at the ground, for example at c and e on Fig. 22 in Case (a) and at c in Case (b). The amount of this loss will depend on the polarization of

EFFECTIVE SPACE WAVE RADIATION AT A LARGE DISTANCE FROM A SHORT VERTICAL DIPOLE 30 FEET ABOVE A SPHERICAL EARTH

Earth's radius 4/3 actual value to allow for the bending near the surface in a standard atmosphere; ground constants σ =0.005 mhos/meter, ϵ =15

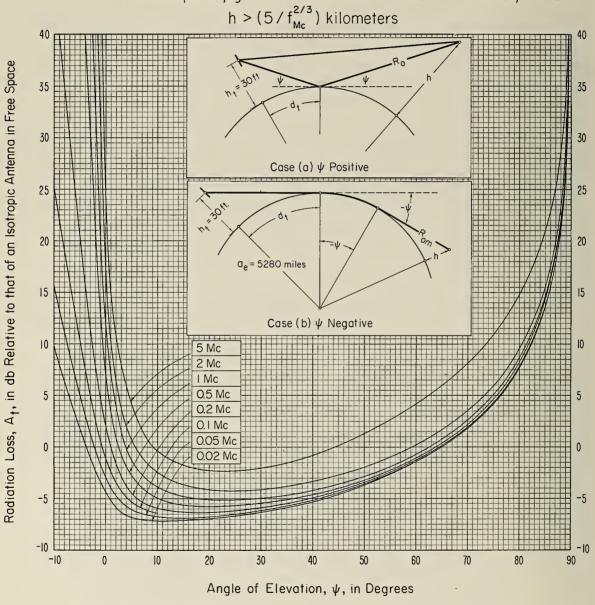


Figure 23

EFFECTIVE SPACE WAVE RADIATION AT A LARGE DISTANCE FROM A SHORT VERTICAL DIPOLE 30 FEET ABOVE A SPHERICAL EARTH

Earth's radius 4/3 actual value to allow for the bending near the surface in a standard atmosphere; ground constants σ = 5 mhos/meter, ϵ = 80

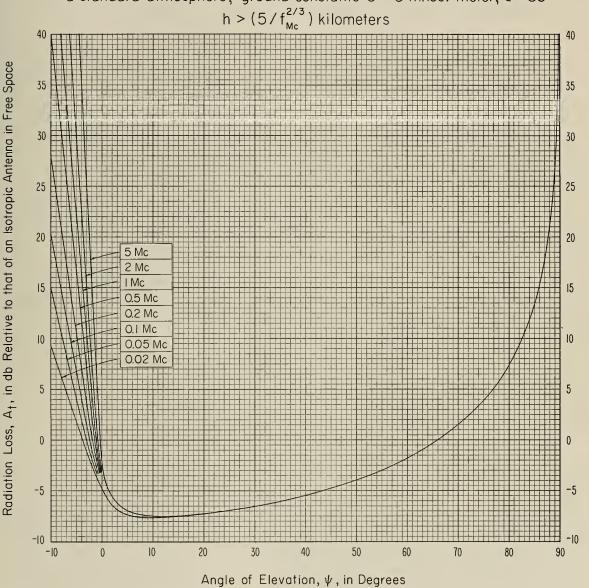


Figure 24

the downcoming waves. Since the waves reflected from the ionosphere will have both vertically and horizontally polarized components, even when the incident waves are linearly polarized, it becomes necessary to know the relative amounts of energy associated with each polarization in the downcoming waves. This problem is not easy to solve precisely, and we have obtained an estimate for A (ψ) by assuming, quite arbitrarily, that the energy in the downcoming waves is equally divided between the two polarizations. Thus, for angles $\psi > 2^{\circ}$, we have:

$$A_g(\psi) \cong -10 \log_{10} [(|R_v^2| + |R_h^2|)/2] \qquad \psi > 2^{\circ}$$
 (14)

where R_v and R_h are the complex Fresnel reflection coefficients for vertical and horizontal polarization, respectively. Since the value determined by (14) represents only a few decibels, the use of the above approximate expression will not lead to serious errors.

It has been shown by Rice $\frac{41}{}$ and by Fock $\frac{42}{}$ that $A_g(0) = 6.021\,db$ when ψ equals zero (in the limit as R_{om}/λ is very large, i.e., for $h > (5/f_{Mc}^{2/3})$ kilometers) regardless of the polarization or ground constants; their results can also be used to compute $A_g(\psi)$ for other values of ψ , but we have instead assumed that $A_g(\psi)$ can be calculated for $\psi \leq 0$ by the following approximate formula:

$$A_g(\psi) \cong 6.021 - 2 A_t(0) + 2 A_t(\psi) \qquad (\psi \leq 0)$$
 (15)

For values of ψ between 0 and 2°, it is easy to sketch in a smooth curve between the results given by (14) and (15).

We turn next to a consideration of the convergence factor $C_m(R, p)$ which provides a measure of the focusing of the energy on reflection at the curved surface of the ionosphere exceeded with probability p. Fig. 25 is a geometrical construction which demonstrates the nature of this focusing of rays in the vertical plane for ψ near zero. At the antipode of the transmitter, half way around the earth, the rays are also focused in the horizontal plane. A detailed discussion of this phenomenon is given in Appendix I. For rays leaving the earth's surface at grazing incidence ($\psi \leq 0$) and at the antipode ($m \theta = 90^{\circ}$), it is necessary to use a wave treatment of the problem, the amount of the focusing then being a function of the frequency. At points substantially removed from these caustics, geometrical optics leads to the following

GEOMETRICAL CONSTRUCTION DEMONSTRATING THE CONVERGENCE OF RAYS NEAR GRAZING INCIDENCE

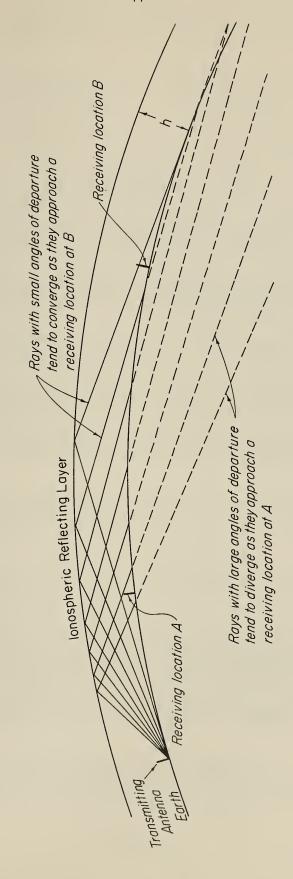


Figure 25

formula for C_m(R) which provides for a smooth ionosphere a measure in decibels of the expected increase in the received field due to this focusing.

$$C_{m}(R) = 10 \log_{10} c_{m} = 10 \log_{10} \left\{ \frac{R_{o}}{-a \sin \psi (d\theta/d\psi)} \right\}_{v} \left\{ \frac{2m R_{o} \cos \psi}{a \sin 2m \theta} \right\}_{h}$$
(16)

$$\frac{d\theta}{d\psi} = -\frac{\sin(\theta - \tau)}{\cos\psi\cos\phi} + \frac{d\tau}{d\psi} \tag{17}$$

In the above τ denotes the total bending of a radio wave in passing through the troposphere, and methods of calculating τ are given in a later section. The two factors in (16) correspond to the focusing in the vertical and horizontal planes, respectively. Equation (16) may be used except near the caustics ($\psi \leq 0$) and (m $\theta = 90^{\circ}$). When $\psi \leq 0$ and m $\theta \neq 90^{\circ}$, we may use:

$$C_{\rm m}(R) = 20 \log_{10} [R_{\rm om} - 92.153 \, \psi \, ({\rm degrees})] + (10/3) \log_{10} f_{\rm kc}$$

- $10 \log_{10} [\sin 2 m \, (\theta_{\rm m} - 4/3 \, \psi)] + (40/3) \log_{10} m - 60.694$ (18)

When $\psi > 0$ and we are at a distance in wavelengths (ρ/λ) from the antipode, we may use:

$$C_{m}(\rho) = 20 \log_{10}(R_{a}/\pi a) + 10 \log_{10}\left\{\frac{\cos^{2}\psi}{-\sin\psi(d\theta/d\psi)}\right\} - 10 \log_{10}m$$

$$+ 10 \log_{10}f_{kc} + 10 \log_{10}[J_{o}(2\pi\cos\psi\rho/\lambda)]^{2} + 36.172 \qquad (\psi > 0)$$
(19)

The above may be used for $m \ge 9$ at night since ψ is greater than zero for these modes with h = 90 km. For $m \le 8$, $\psi < 0$ near the antipode at night, and we may then use the following formula:

$$C_{\rm m}(\rho) = 20 \log_{10} (R_{\rm a}/\pi a) - (20/3) \log_{10} m + (40/3) \log_{10} f_{\rm kc}$$

$$+ 10 \log_{10} [J_{\rm o} (2\pi \rho/\lambda)]^2 + 44.422 \qquad (\psi \le 0)$$
(20)

In (19) and (20), R_a denotes the ray distance to the antipode. If a horizontal magnetic dipole is used for reception of the field radiated from the vertical electric dipole, then the J_o in (19) and (20) is to be replaced by J_1 ; here J_o and J_1 are Bessel functions. Very near $\psi = 0$, the values of $C_m(R)$ determined by (16) will exceed those given by (18), particularly at the lower frequencies, and in this region a smooth curve may be drawn between the values for $\psi > 0$ and the values given by (18); similarly, at the antipode the values of $C_m(\rho)$ given by (19) will exceed those given by (20), and a smooth curve may be drawn between the values for $\psi > 0$ and those given by (20). Fig. 26 gives examples of $C_m(R)$ calculated in this way for m = 1, 2 and 9 for daytime propagation (h = 70 km), for a smooth ionosphere, and typical refraction conditions.

It is clear from Fig. 25 that this focusing will be fully realized in practice only to the extent that the ionosphere presents a smooth surface to the radio waves. A discussion is presented in Appendix I which indicates how allowance may be made for ionospheric roughness. It is shown that an individual ionospheric mode of propagation consists of a steady specularly-reflected component plus a random Rayleigh distributed component. If we let k^2 denote the ratio of the power in the random component relative to that in the specularly-reflected component, then we may estimate the convergence $C_m(R, p)$ exceeded 100 p% of the time in terms of the values of $k^2(1-p)$ exceeded 100(1-p)% of the time:

$$C_{m}(R, p) = 10 \log_{10} \left[\frac{c_{m} + k^{2}(1-p)}{1+k^{2}(1-p)} \right]$$
 (21)

Here c (see 16) denotes the ratio of the received power with and without focusing at a smooth ionosphere. As the probability varies from 0 to 1, $k^2(1-p)$ will vary from zero for a smooth ionosphere to ∞ for a perfectly rough ionosphere, and $C_m(R, p)$ will vary from the values given by (16), (18), (19), and (20) for a smooth ionosphere to zero for a perfectly rough ionosphere. The transmission losses given in this report correspond to median values, i.e., p = 0.5. The random variable k^2 depends upon the radio frequency, angle of incidence, ϕ , and time of day.

As an illustration of the effects of focusing near the antipode and of the influence of ionospheric roughness, Fig. 27 shows the

CONVERGENCE FACTOR IN IONOSPHERIC PROPAGATION $h = 70 \text{ kilometers}; N_s = 313 \text{ and } h_s = 0 \text{ and } k^2(l) = 0$

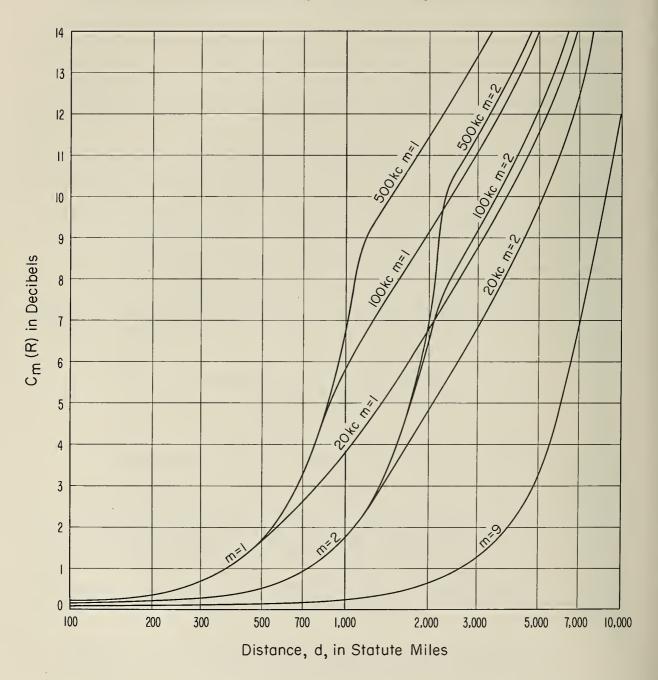


Figure 26

TRANSMISSION LOSS EXPECTED NEAR THE ANTIPODE AT NIGHT FOR 100 kc

Over Land = σ = 0.005 Mhos/meter, ϵ = 15

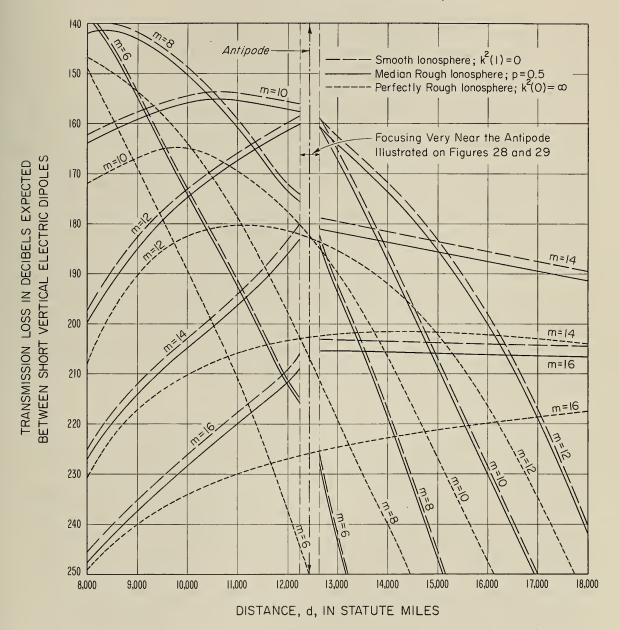


Figure 27

transmission loss expected between vertical electric dipoles in the range from 8,000 to 18,000 miles for 100 kc at night over land. Three values are shown for each of the modes m = 6, 8, 10, 12, 14, and 16 corresponding (a) to a smooth ionosphere (p = 0, $k^2(1) = 0$), (b) to an ionosphere of median roughness (p = 0.5), and (c) to a perfectly rough ionosphere (p = 1, $k^2(0) = \infty$).

In the immediate vicinity of the antipode, i.e., within a few wavelengths, the focusing for a smooth concentric ionosphere is very large. Thus Fig. 28 shows the value of Cm(Ra, p) expected right at the antipode at night for a smooth, concentric ionosphere and for a rough, concentric ionosphere; the values expected for h = 70 km during the daytime would be only slightly different; Fig. 29 shows the rapid decrease of the focusing as we leave the antipode and, at distances greater than 100\lambda, the envelope will be just 6 db above the values of transmission loss shown on Fig. 27, i.e., the values shown on Fig. 27 correspond to the single wave expected from a directive transmitting antenna with an infinite front-to-back ratio. We see on Fig. 29 that the field expected from the non-directive dipole oscillates with increasing distance from the antipode; this oscillation is caused by the interference between the waves arriving at the receiving point along the short and long great circle paths. Thus there will be concentric rings around the antipode at which the expected field will be equal to zero. The radii of these concentric rings are the same for a concentric ionosphere, regardless of the number, m, of ionospheric reflections, and are determined by the zeros of the Bessel functions; for the electric field, the first two such rings have radii equal to 0.38λ and 0.88λ.

The actual ionosphere will never be concentric with the surface of the earth. In practice, as the sun rises and sets, or as the geomagnetic latitude of the reflection point is varied, the surface of the ionosphere will undoubtedly change in such a way that its radius of curvature and slope relative to a tangent plane on the earth will vary over appreciable ranges, and this will cause $C_m(R)$ to vary up and down relative to the values expected on the basis of the above analysis. However, except near the antipode, it seems plausible to assume that the median values of $C_m(R)$ may not be much influenced by such changes. The magnitude of the antipodal anomaly will be substantially reduced by these macroscopic perturbations of the spherical concentric shell model assumed for our calculations. Also, for the actual non-concentric ionosphere, the geographical location of the antipode may be expected to vary with time over a fairly large

EXPECTED FOCUSING AT THE ANTIPODE FOR A CONCENTRIC IONOSPHERE

 $h=90KM; N_s=324$

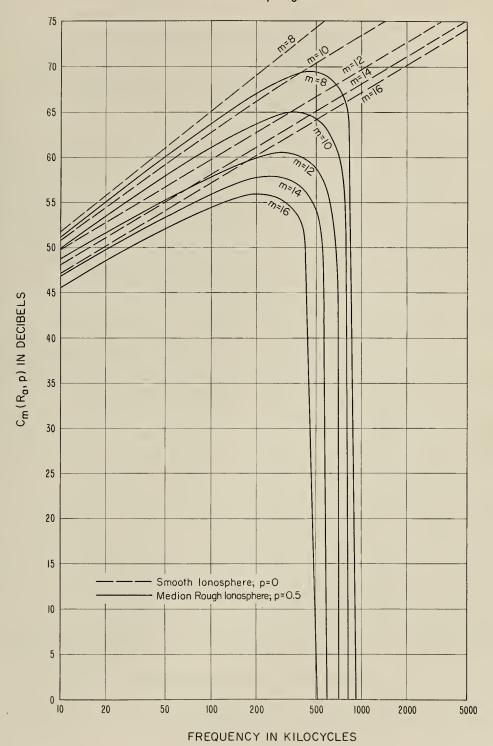


Figure 28

THE BEHAVIOR NEAR THE ANTIPODE OF THE CONVERGENCE FACTORS FOR THE RADIATION FROM A VERTICAL ELECTRIC DIPOLE FOR A CONCENTRIC IONOSPHERIC SHELL AROUND A SPHERICAL EARTH

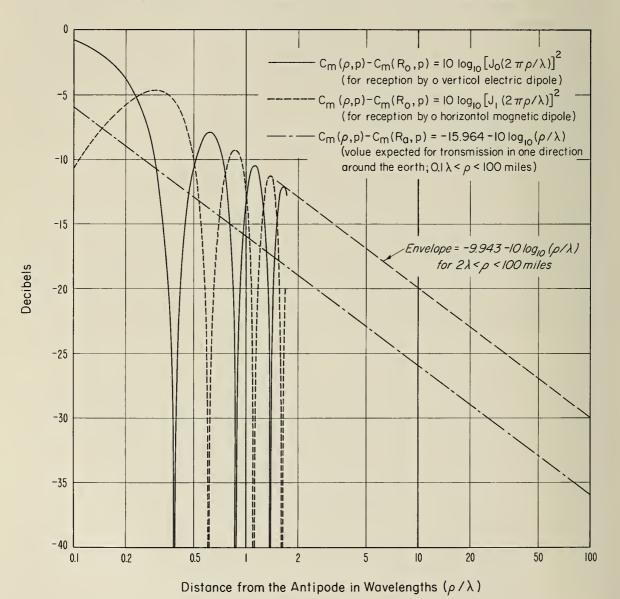


Figure 29

area, and this should result in a net increase in the fading range in this region. Furthermore, the antipodal locations may be expected at any given time to be different for the different modes, and thus the zeros predicted by (19) and (20) and shown on Fig. 29 are not likely to be observable unless some means is used to exploit the different times of arrival of the individual modes.

Consider next the loss, P, arising from the polarization characteristics of the downcoming ionospheric waves. An incident linearly-polarized wave will be reflected as two component waves. an ordinary and an extraordinary wave, each of which will be elliptically polarized. These two component waves, which will have roughly the same amplitudes except on frequencies near the gyrofrequency (near 1.5 Mc in the United States) will mutually interfere, and this causes the rapidly varying polarization characteristics of the observed downcoming waves. The polarization loss, P, arises from the fact that typical receiving antennas will respond to only one polarization. The amount of this loss will depend principally upon the transmission frequency, the penetration frequency for the layer involved, and the intensity and direction of the earth's magnetic field relative to the path; it can be calculated $\frac{13}{34}$ / 43/ with some accuracy when these parameters are known. For more than one reflection at the ionosphere, the polarization loss becomes a very complex function of the reflection coefficients for the parallel and perpendicular components of the incident fields and is difficult to separate from the absorption loss, $A(\phi, 1 - p)$. All of the low frequency examples of ionospheric wave propagation in this report have been obtained using the empirical estimates described below for P + A(ϕ_0 1 - p) which thus includes the polarization loss P; consequently we have calculated the total reflection loss as $m\{P + A(\phi, 1 - p)\}$ - (m - 1)P for m > 1. The calculations at f = 100 kc have been made in this report by setting P = 3.01 db, but this estimate is now believed to be substantially too large, except near vertical incidence. The calculations for all of the other frequencies from 20 kc to 1,000 kc were calculated with P = 0 for all values of m, although this assumption probably leads to somewhat more transmission loss than would be expected for m > 1 since the appropriate value of P probably lies between 0 and 3 db, approaching the latter value near vertical incidence; however, we are usually more interested in the values near oblique incidence for our applications, and this latter assumption should yield more nearly correct results for the solution of these problems.

Finally we will consider the loss, $A(\phi, 1 - p)$, on reflection at the ionosphere, exceeded for 100(1 - p)% of the time; this depends

MEDIAN LOSS FOR ONE REFLECTION AT THE IONOSPHERE

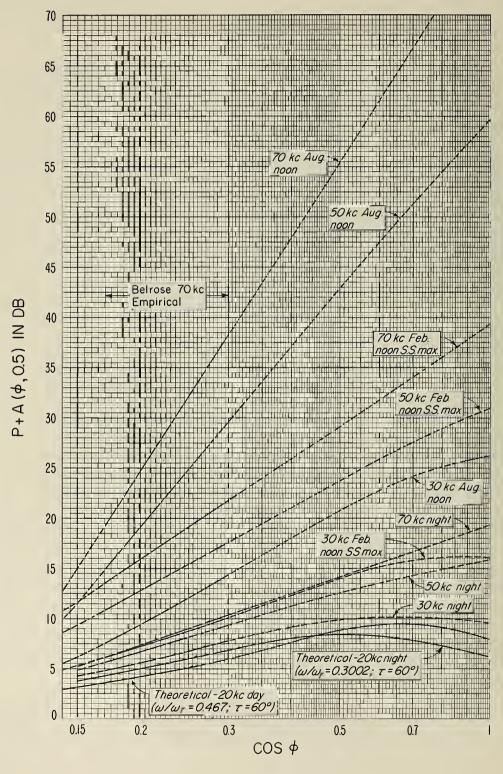


Figure 30

on the angle of incidence, ϕ , at the ionosphere, the radio frequency, time of day, season of the year, phase of the sunspot cycle and the geomagnetic latitude of the reflection point. For the examples developed in this report, we have used some empirical evaluations of P + A(ϕ , 1 - p) made by Belrose $\frac{45}{}$ on transmission paths between stations in England, the Scandinavian countries, and Germany. Using the results in his doctoral thesis, we may express the median values of P + A(ϕ , 0.5) as follows:

$$P + A(\phi, 0.5) = 17.2 \log_{10} (f_{kc} \cos \phi) - 12.4$$
 (70 < $f_{kc} < f_D$) (22) (NIGHT)

P + A(
$$\phi$$
, 0.5) = 30.8 log₁₀ (f_{kc} cos ϕ) - 22.6 (70 < f_{kc} < f_D) (23) (FEB., NOON, SUNSPOT MINIMUM)

P + A(
$$\phi$$
, 0.5) = 33.6 log₁₀ (f_{kc} cos ϕ) - 22.6 (70 < f_{kc} < f_D) (24) (FEB., NOON, SUNSPOT MAXIMUM)

P + A(
$$\phi$$
, 0.5) = 77.3 log₁₀ (f_{kc} cos ϕ) - 64.0 (70 < f_{kc} < f_D) (25) (AUG., NOON)

The above formulas were determined empirically from data extending only over the range of frequencies from 70 - 250 kc and the range of distances from 350 to 900 miles; a recent analysis of lower frequency data by Watt, Maxwell and Whelan $\frac{46}{}$ indicates that the absorption is greater at frequencies less than 70 kc than would be predicted by the above formulas. Consequently, as shown on Fig. 30, we have used the theoretical results of Wait and Murphy $\frac{34}{}$ at 20 kc and then interpolated linearly on a logarithmic frequency scale to obtain values for intermediate frequencies; the same ionospheric parameters L/H = 0.1 by day (i.e., $\omega/\omega = 0.467$) and L/H = 0.05 at night (i.e., $\omega/\omega = 0.3002$) were used in these calculations at f = 20 kc as for those leading to Figs. 16 to 19, but the index τ for the earth's magnetic field was set equal to 60° in the present calculations, whereas τ was set equal to zero in the calculations leading to Figs. 16 to 19.

Figs. 31 to 36 give the median transmission loss expected in accordance with the above methods of calculation at 20, 50, and 200 kc in over-land and over-sea propagation and for day and night conditions.

MEDIAN TRANSMISSION LOSS OVER LAND AT 50 kc σ = 0.005 Mhos/meter; ϵ =15; h_t = h_r =30 feet Day h = 70 km; Night h = 90 km

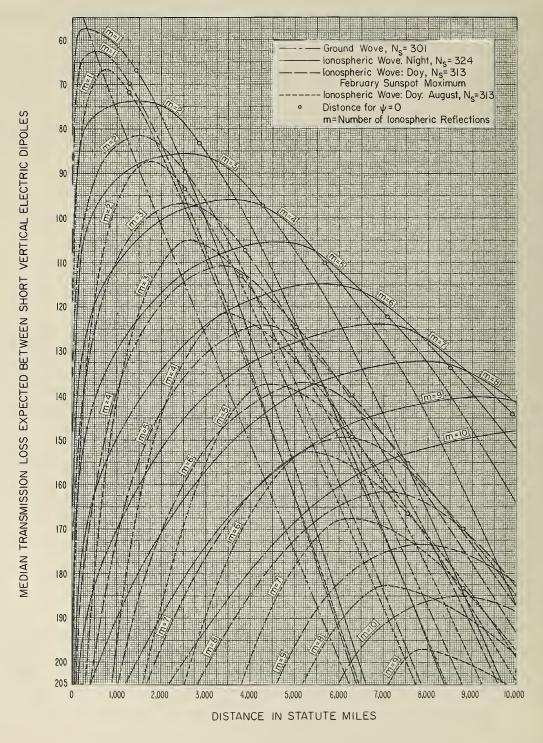


Figure 33

MEDIAN TRANSMISSION LOSS OVER SEA AT 50 kc

 σ = 5 Mhos/meter; ϵ = 80; h_t = h_r = 30 feet Day h = 70 km; Night h = 90 km

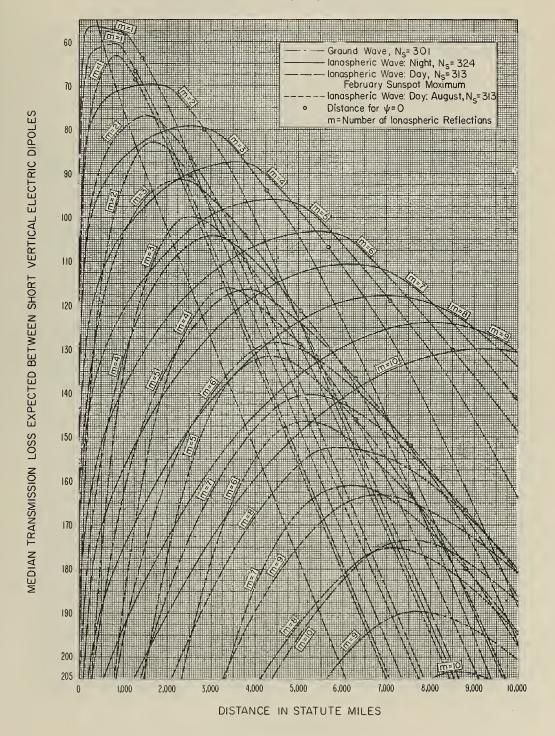
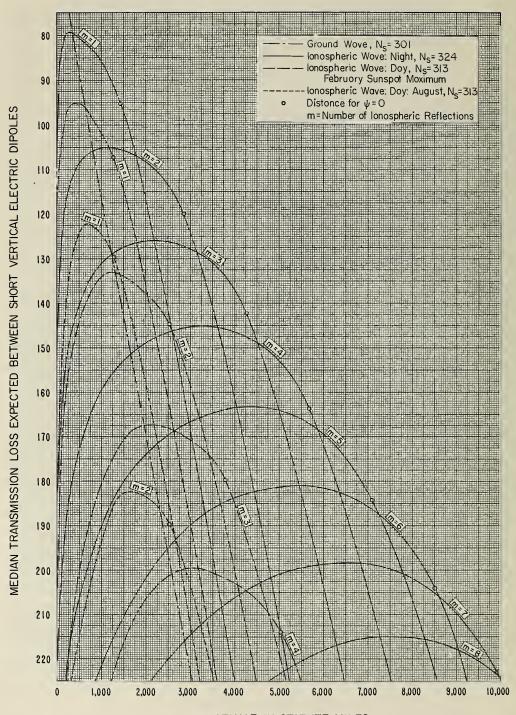


Figure 34

MEDIAN TRANSMISSION LOSS OVER LAND AT 200 kc σ = 0.005 Mhos/meter, ϵ =15; h_t=h_r = 30 feet Day h = 70 km; Night h =90 km



DISTANCE IN STATUTE MILES

Figure 35

MEDIAN TRANSMISSION LOSS OVER SEA AT 200 kc σ = 5 Mhos/meter; ϵ = 80; h_1 = h_r = 30 feet Day h = 70 km; Night h = 90 km

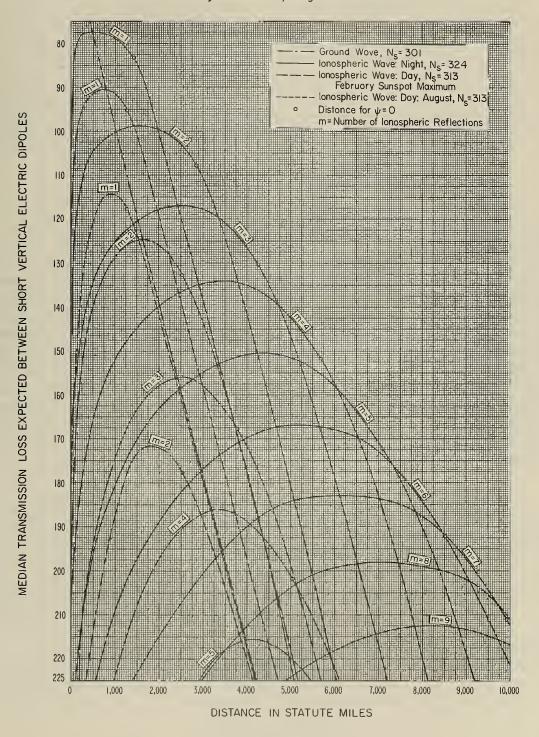


Figure 36

Figs. 31 and 32 show the large decrease in the transmission loss near the antipode, and illustrate the fact that each mode of propagation has two important branches at points somewhat removed from the antipode, corresponding to propagation via the short and long great circle paths, respectively; actually there are still other branches for each mode corresponding to propagation more than once around the earth, but these are not shown. The decrease in transmission loss shown at the antipode is the value for an idealized concentric ionosphere and will, in practice, undoubtedly be somewhat smaller. Transmission loss curves for the separate modes are not given at 20 kc for sea water since they differ so little from those for overland. It should be noted that the curves on Figs. 16 and 17 for 20 kc will be more useful for most applications than those on Fig. 31 since they combine the separate modes with proper relative phases.

Throughout this section the formulas and graphs refer to the median values of transmission loss for individual modes of propaga-This form of presentation was used since it is more useful in applications such as the design of navigation systems or of systems to avoid multipath distortion. To determine the expected median transmission loss for a continuous wave transmission, it is necessary to convert the transmission losses for the individual modes to power ratios, and then add these power ratios; at 500 kc, and possibly even as low as 50 kc, it is reasonable to assume that the several ionospheric modes will have random relative phases so that the median power of the resultant will be equal to the sum of the median powers of the individual modes. For example, if there were two modes with equal median transmission losses, the median transmission loss for the sum of these two modes would be 3 db less, and for three equal modes the sum would have 4,77 db less transmission loss than each individual mode.

The author has studied the behavior of $P + A(\phi, 0.5)$ at night in the United States for frequencies in the standard broadcast band from 500 to 1,500 kc over a very wide range of distances, and has found the following semi-empirical formula:

P + A(
$$\phi$$
, 0.5) = $\frac{26 \cos \phi}{(f_{\text{mc}} \cos \phi)^{0.4}}$ (26)

(NIGHT)

Note that (22) indicates an increasing loss with increasing frequency, presumably because of a deeper penetration of the D layer as the frequency is increased, while (26) indicates that the loss decreases with increasing frequency. At the higher frequencies where (26) was established, the waves penetrated the D layer and, as shown by Martyn $\frac{47}{}$ this behavior of P + (ϕ , 0.5) with frequency and angle of incidence is to be expected. By virtue of the method used for its determination, (26) includes the polarization loss P; since the extraordinary waves are much weaker than the ordinary waves in this frequency range, there will be additional polarization loss at each reflection from the ionosphere. With the above discussion in mind, it seems appropriate to assume that the penetration frequency of the D layer at night is effectively defined by the following relation:

$$P + A(\phi, 0.5)_{(26)} = P + A(\phi, 0.5)_{(22)}$$
 (27)

(At the D layer penetration frequency, f_{D} , at night)

As determined in this way, the D layer penetration frequency at night varies from about 500 kc at vertical incidence to about 250 kc with $\cos \phi = 0.164$, the minimum value expected for a 90 km layer height; the anomalous behavior of this penetration frequency suggests that neither of our empirical absorption formulas are very dependable in this intermediate range of frequencies. Since nothing better is readily available, it was decided to calculate the transmission loss at night at frequencies greater than the above-defined D layer penetration frequency by using m 26 cos $\phi/(f_{MC}\cos\phi)^{0.4}$ as the total loss on reflection; the reflection height was assumed to be 110 km at night, but the value of $\cos \phi$ to be used in the absorption equation was determined on the assumption that the absorption takes place at a height of 100 km.

Figs. 37 to 40 give the median transmission loss expected between short vertical electric dipoles at 500 kc and at 1,000 kc in over-land and over-sea propagation and for day and night conditions. The absorption at night was determined by (26) as described above, but, in the daytime, (24) and (25) were used since radio waves in this frequency range are then presumably reflected and absorbed by the D layer at an assumed height of 70 km.

At still higher frequencies during the daytime, the radio waves will penetrate the D layer and be reflected by the E layer at a height of about 110 km. The ionospheric absorption is so great during the daytime in the range of frequencies from, say 500 kc to 4 Mc, and

MEDIAN TRANSMISSION LOSS OVER LAND AT 500 kc σ = 0.005 Mhos/meter, ϵ = 15; h_t = h_r = 30 feet Day h = 70 km; Night h = 110 km

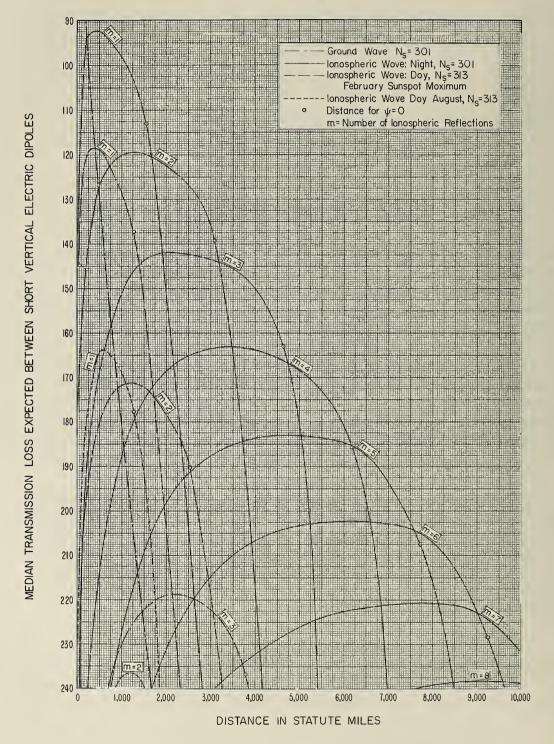


Figure 37

MEDIAN TRANSMISSION LOSS OVER SEA AT 500 kc σ = 5 Mhos/meter , ϵ = 80; h = h = 30 feet Day h = 70 km , Night h = 110 km

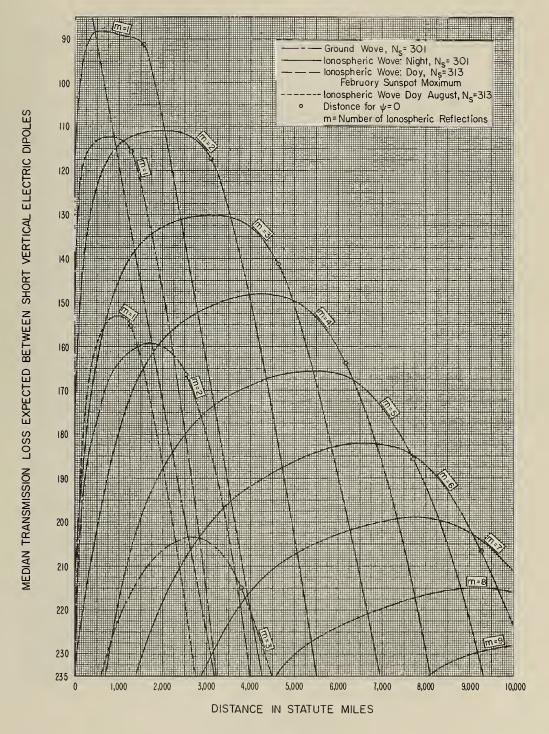


Figure 38

MEDIAN TRANSMISSION LOSS OVER LAND AT 1000 kc σ = 0.005 Mhos/meter, ϵ = 15; h_t = h_r = 30 feet Day h = 70 km; Night h = 110 km

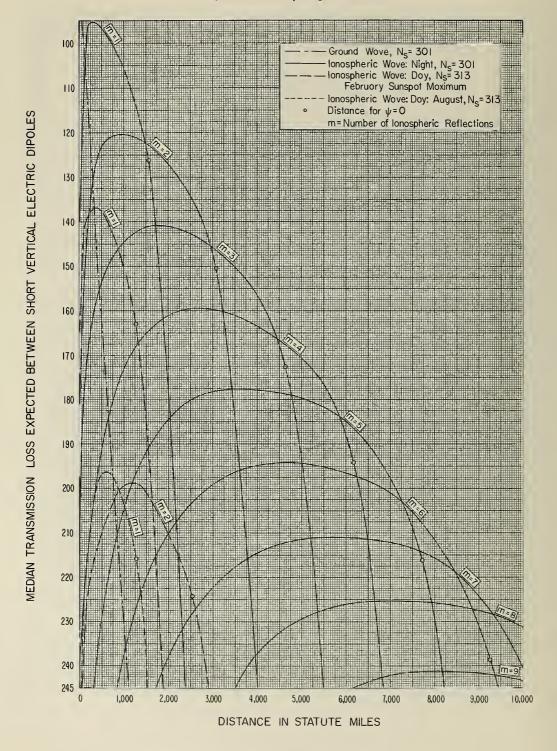


Figure 39

MEDIAN TRANSMISSION LOSS OVER SEA AT 1000kc σ = 5 Mhos/meter; ϵ = 80, h_1 = h_r = 30 feet Day h = 70 km; Night h =110 km

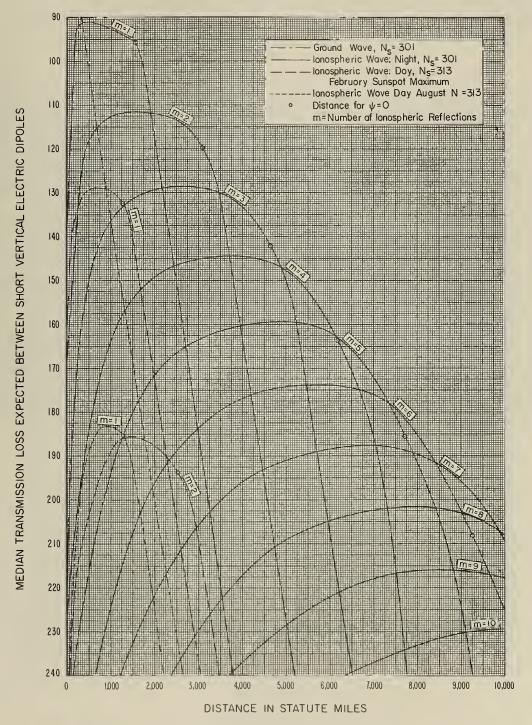


Figure 40

the reflection phenomena so complex, with reflections taking place sometimes at the D layer, sometimes at the E layer and sometimes at the F layer, that useful, simple absorption formulas are not available. For this reason the calculations of transmission loss given in this report for this range of frequencies have been made by extrapolating (23), (24) and (25) to higher frequencies, and by extrapolating to lower frequencies the absorption formulas applicable to the high frequency band as discussed in the next section.

Figs. 41 to 45 give transmission losses for propagation over land, based on the above-described methods of computation and on the methods described in following sections, and are designed to show more clearly the effect of radio frequency for day and night, for two seasons and for minimum and maximum sunspot conditions. Only one set of curves are presented for propagation at night since the seasonal and sunspot cycle effects on the transmission loss are comparatively small at night. The curves on these figures give the transmission loss expected between short electric dipole antennas, oriented vertically at frequencies less than 5 Mc and horizontally for frequencies greater than 5 Mc, for the ground wave and for the particular sky wave mode with a minimum transmission loss at the distances 200, 500, 1,000, 2,000, 5,000 and 10,000 miles. At each distance we have shown only the value expected for the single sky wave mode with the minimum transmission loss; with continuous wave transmission, the losses would be several db less than these values, particularly at the larger distances where several sky wave modes with comparable intensities are expected. Note that ground proximity losses L and L, as discussed in Appendix III, have been omitted in calculating the values shown on Figs. 41 to 45.

MEDIAN TRANSMISSION LOSS FOR MIDNIGHT AT THE RECEIVING ANTENNA February; Sunspot Maximium; σ =0.005 Mhos/meter, ϵ =15; h_{t} = h_{r} =30 feet West to East Transmission Path with Washington, D.C. at the Midpoint

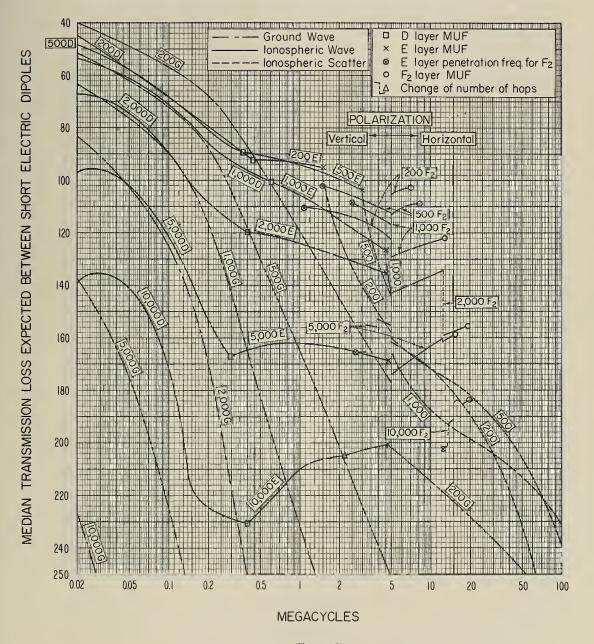


Figure 41

MEDIAN TRANSMISSION LOSS FOR NOON AT THE RECEIVING ANTENNA February; Sunspot Minimum; σ =0.005 Mhos/meter; ϵ =15; h_t=h_r=30 feet West to East Transmission Path with Washington, D.C. at the Midpoint

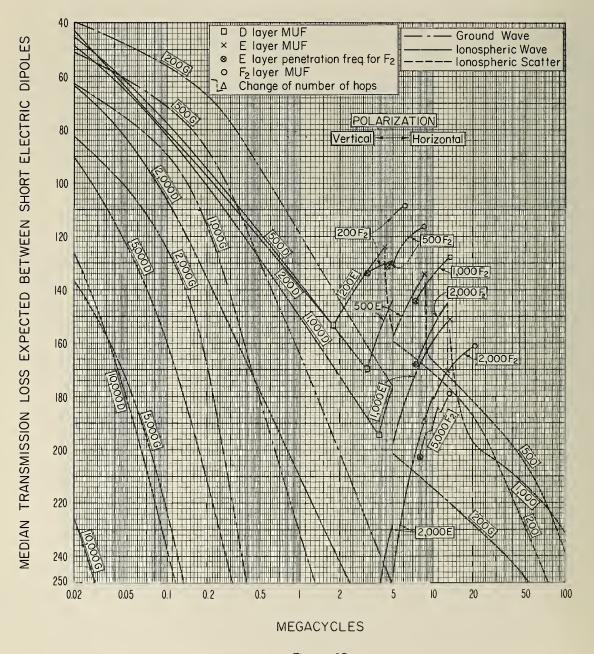


Figure 42

MEDIAN TRANSMISSION LOSS FOR NOON AT THE RECEIVING ANTENNA February; Sunspot Maximium; σ = 0.005 Mhos/meter, ϵ = 15; h_{t} = h_{r} = 30 feet West to East Transmission Path with Washington, D.C. at the Midpoint

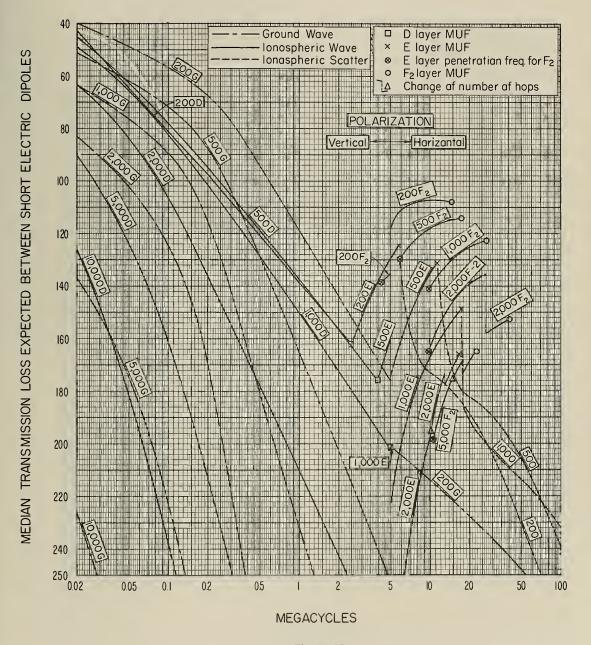


Figure 43

MEDIAN TRANSMISSION LOSS FOR NOON AT THE RECEIVING ANTENNA August; Sunspot Minimum; σ =0.005 Mhos/meter; ϵ =15; h_t=h_r=30 feet West to East Transmission Path with Washington, DC at the Midpoint

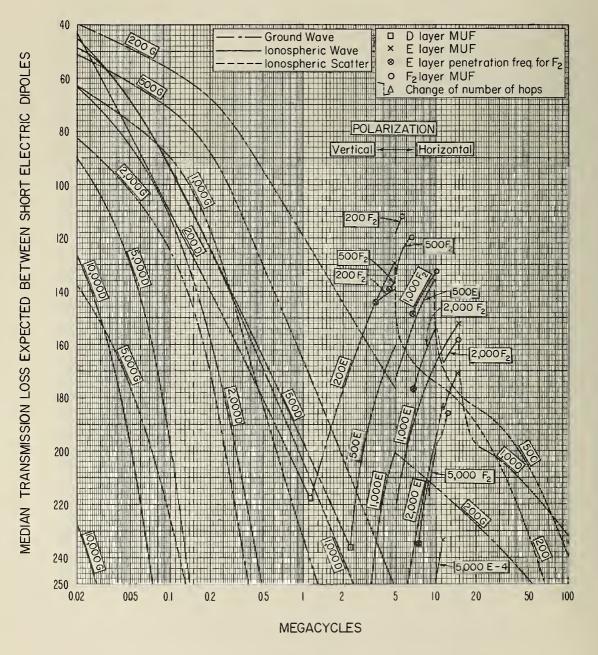


Figure 44

MEDIAN TRANSMISSION LOSS FOR NOON AT THE RECEIVING ANTENNA August; Sunspot Maximum; σ =0.005 Mhos/meter; ϵ =15; h_t = h_r =30 feet West to East Transmission Path with Washington, D.C. at the Midpoint

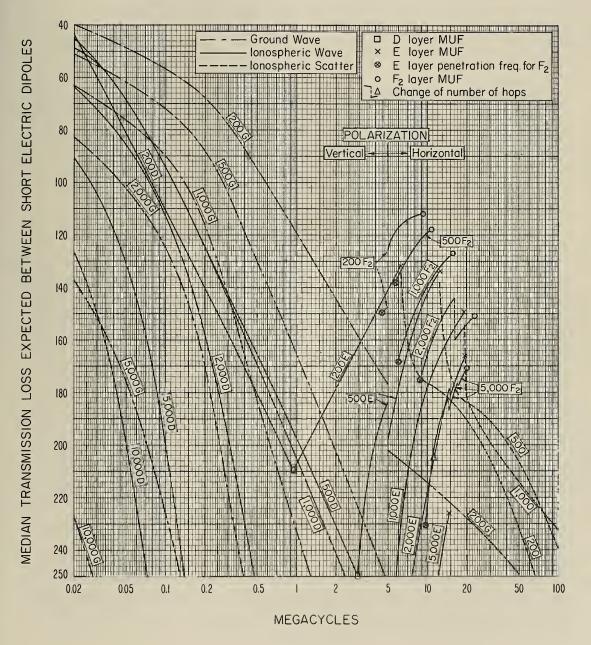


Figure 45

4.3 High Frequency Ionospheric Propagation

We see on Figs. 41 to 45 that the variation of transmission loss with frequency changes as the various layers of the ionosphere are penetrated. The method of determining the D layer penetration frequency was described in the preceding section, and the penetration frequencies for the higher layers were determined by the methods described in references 24 and 25; for this purpose the propagation path was assumed to be from West to East with its midpoint at Washington, D.C., and with either noon or midnight at the eastern end of the path.

During the daytime the polarization and absorption losses in the high frequency band have been calculated by the methods described in a Signal Corps report. 48/ Thus the constant attenuation of 8.9 db found in their analysis has been somewhat arbitrarily attributed to a polarization loss P, and the following semi-empirical formula used for calculating the daytime absorption:

$$A(\phi, 0.5) = \frac{615.5 \left\{\cos (0.881 \chi)\right\}^{1.3} (1 + 0.0037 s)}{\left(f_{Mc} + f_{H}\right)^{1.98} \cos \phi}$$
(DAY) (28)

In this formula x denotes the zenith angle of the sun at the reflection point, and s denotes the smoothed Zurich sunspot number; for sunspot minimum, s was set equal to 10, and for sunspot maximum, s was set equal to 150. The gyrofrequency, $f_{LI} = 1.5 \text{ Mc}$, on the average in the United States. The angle of incidence, ϕ , to be used in (28) refers to the value this angle will have at the absorption level, and this angle will be systematically larger, for a given angle of elevation ψ , than the angles of incidence at the higher layers where the reflections take The Signal Corps analysis was based on the assumption that the absorption takes place at a height of 100 km and this same height was used in our calculations. In the Signal Corps report, convenient graphical methods are given for determining many of the factors involved in calculating (28). Prof. A. Kazantsev 49/50/ has proposed a method for calculating $A(\phi, 0.5)$ which, in essence, involves the replacement of the numerator of (28) by a constant times the square of the penetration frequency of the E layer at vertical incidence; this method appears to have considerable merit, but a critical determination of its accuracy compared to that of the Signal Corps method has not yet been published.

Thus, during the daytime we have used (9) for calculating the median transmission loss in the high frequency band with P = 8.9 db and $A(\phi, 0.5)$ determined by (28). Note that the focusing is probably negligible in this frequency range since the ionosphere will very likely appear rough to these radio waves most of the time; in the absence of quantitative information on ionospheric roughness at these higher frequencies, we have arbitrarily set $C_{m}(R, 0.5) = 0$ at all distances for $f \geq 2$ Mc both day and night.

At night the absorption in the high frequency band is quite small; it has been estimated in this report by means of (26). We have already noted, however, that (26) includes a polarization loss, P, and the Signal Corps report indicates, in effect, that the absorption plus the polarization loss at night is equal to 8.9 db; thus it appears to be appropriate to use (9) for calculating the transmission loss with $P + A(\phi, 0.5)$ calculated by (26) for the lower frequencies where $m\{P + A(\phi, 0.5)\} > 8.9$ db and to set $m\{P + A(\phi, 0.5)\} = 8.9$ db for all higher frequencies. This is the method used for the examples presented in this report.

4.4 Ionospheric Scatter Propagation

At frequencies above the penetration frequency of the E layer, the radio waves are scattered forward with sufficient intensity to be usable for communications over distances of the order of 600 to 1,400 miles. $\frac{51}{52}$ $\frac{53}{53}$ The transmission losses shown on Figs. 41 to 45 for this mode of propagation are based on the measurements reported by Bailey, Bateman and Kirby $\frac{51}{5}$ for the Fargo, North Dakota to Churchill, Manitoba path as extrapolated to other distances and frequencies by means of a theory developed by Wheelon. $\frac{54}{55}$ $\frac{56}{56}$ Thus Wheelon attributes the scattering to turbulence in the D and E regions of the ionosphere and, on the assumptions (1) that the spectrum of this turbulence may be determined by the mixing in gradient hypothesis and (2) that the viscosity cut-off has a characteristic scale 1.5 meters, is able to develop a formula for the transmission loss expected with this mode of propagation.

Wheelon's analysis leads directly to the spectrum of the turbulence, but the turbulence may also be characterized in the range of wave numbers smaller than the viscosity cut-off by the correlation function $(r/l_0)K_1(r/l_0)$ which describes the degree of correlation in the fluctuations in electron density at points a distance r apart; K_1

denotes the modified Bessel function of the second kind and ℓ_0 is a characteristic scale of the turbulence, set equal to 100 meters in our subsequent analysis. It is interesting to note that this same correlation function is applicable for describing tropospheric turbulence as well. $\frac{57}{58}$

The following formula gives the median transmission loss expected for the ionospheric scatter mode of propagation, i.e., on frequencies above the effective maximum usable frequency, f MUF, of the scattering region:

$$L_{ms} = L_{bf}(R) + A_{t}(\psi) + A_{r}(\psi) - S(0.5) - 10 \log_{10} \sec \phi$$

$$+ B(k^{2}, \ell_{o}, \ell_{s}) + P + A(\phi, 0.5) \qquad (f_{Mc} \ge f_{MUF}) \qquad (29)$$

In the above, $L_{\rm bf}(R)$ denotes the free space transmission loss for waves traveling a distance corresponding to an average scatter path of length R; this path length has been determined in this report on the assumption that the mean layer height h = 87 km both day and night. For horizontally polarized waves the following formula may be used to estimate the median value for the sum of the space wave radiation factors:

$$A_{t}(\psi) + A_{r}(\psi) = -20 \log_{10} [2 \sin (2\pi h_{t} \sin \psi/\lambda)]$$

$$-20 \log_{10} [2 \sin (2\pi h_{r} \sin \psi/\lambda)] - G_{p}(0.5) \quad \text{(Horizontal polarization)}$$
(30)

In the above, h and h denote the heights of the transmitting and receiving antennas above the local terrain, and G (0.5) denotes the median path antenna gain. For the high gain antennas normally used for communication by scatter, there will usually be a substantial "loss in gain" relative to the value G would be expected to have for communication between similar antennas in free space. For example, on the Fargo to Churchill path, G as determined for successive half-hour periods of time, was found to be a random approximately normally distributed variable with a median value G (0.5) = 25.7 db and a standard deviation of 5.85 db; the sum of the free space gains in this case was about 40 db. It has been found that some of this "loss in gain" can be recovered by directing the antenna towards the better scattering regions. The values of transmission loss shown on Figs. 41 to 45 are

for propagation between short horizontal electric dipoles, and in this case $G_p = 3.52$ db. The antenna heights, h_t and h_r , were taken to be 30 feet except at the higher frequencies where somewhat lower values of h_t and h_r were chosen so that $4h_t$, $\sin \psi/\lambda = 1$; this choice of h_t and h_r effectively minimized the loss and in this case h_t $h_$

The factor S involves the intensity and scales of the turbulence and, together with G, exhibits most of the variability of the transmission loss. The median value S(0.5) undoubtedly varies somewhat diurnally, seasonally, and with the sunspot cycle but, since such changes are not large with a probable extreme range of the monthly medians at a given time of day of less than 20 db, we have calculated all of the examples in this report by setting S(0.5) = -8.4 db, the value obtained for the Fargo-Churchill path. An analysis is presented in Appendix II which shows that this value of S(0.5) is not inconsistent with what is presently known about ionospheric turbulence.

The factor $10 \log_{10}$ sec ϕ provides a measure of the size of the effective scattering volume for transmission paths of various lengths.

The transmission loss factor $B(k^2, \ell_0, \ell_s)$ may be expressed: $B(k^2, \ell_0, \ell_s) = 25 \log_{10} \left[1 + k^2 \ell_0^2\right] + 20 \log_{10} \left[1 + (k^2 \ell_s^2)^{2/3}\right] + (40/3 \log_{10} \left[1 + (k^2 \ell_s^2)^2\right]$ (31)

Although the characteristic scale lengths, ℓ_0 and ℓ_s , are likely to be somewhat variable diurnally and seasonally, we have, for the purpose of the calculations in this report, taken them to be equal to the constant values $\ell_0 = 100$ meters and $\ell_s = 1.5$ meters; k^2 is defined as follows:

$$k^{2} = \left[\frac{4\pi}{\lambda} \cos \phi\right]^{2} \left(1 - \frac{f_{MUF}^{2}}{f^{2}}\right)$$
 (32)

When f is equal to the maximum usable frequency, f_{MUF} , $k^2 = 0$, the wavelength in the medium increases without limit, and the scattering is no longer directed forward, but occurs uniformly in all directions. In this limiting case, $B(k^2, l_0, l_s) = 0$ and (29) indicates a scatter loss exceeding that predicted by (9) for normal E layer propagation at the MUF by only (8.4 - 10 \log_{10} sec ϕ) decibels. For the calculations in this report, we have somewhat arbitrarily used the median E layer MUF as a measure of the MUF of the scattering region.

For ionospheric scatter, we have taken P = 3 db for both day and night propagation conditions. During the day, the absorption term $A(\phi, 0.5)$ was computed by (28), but at night $P + A(\phi, 0.5)$ was determined by (26) up to frequencies for which the resulting value is greater than 3 db, and at higher frequencies $P + A(\phi, 0.5)$ is set equal to 3 db.

There is no present evidence for the existence of F layer ionospheric scatter or for multi-hop E layer scatter and, for this reason, the transmission loss curves for d = 2,000, 5,000, and 10,000 miles stop abruptly at the MUF; the transmission loss is expected to increase very rapidly indeed at frequencies just above the F layer MUF.

5. The Bending of Radio Waves by the Troposphere

Since the density as well as the absolute humidity of the air decrease with the height, h, above sea level, the refractive index, n, also usually decreases with h, and this causes radio waves leaving an antenna at a given angle, ψ , to bend down towards the earth, the amount of this bending being larger, the smaller the value of ψ . This is illustrated on Fig. 46 which shows the total bending, τ , of a radio wave traveling entirely through the troposphere and subsequently being reflected at an ionospheric layer.

Since the refractive index, n, departs from unity by only a few parts in 10⁻⁴, it is convenient to describe n in terms of the refractivity, N, which is defined:

$$N = (n - 1) \times 10^{6}$$
 (33)

If the value of N were known as a function of time at every point in the atmosphere between two radio antennas, it should be possible, in principle, to predict the instantaneous behavior of the transmission

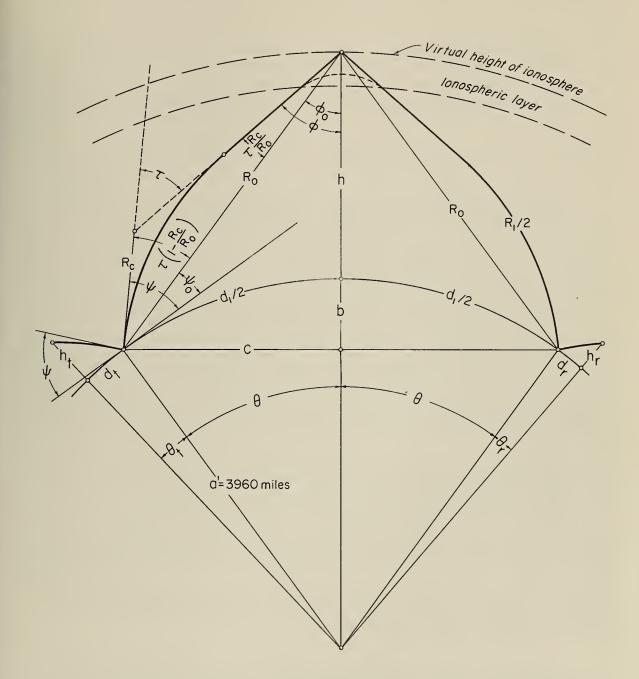


Figure 46

loss in propagation between these antennas. Actually, of course, this is not feasible because of the complexity of the solution of such an electromagnetic problem. Furthermore, even if the engineer could be provided with this instantaneous information, he would normally be forced to describe it in some statistical terms before he could use it effectively in the design or use of radio systems. Consequently, we are led to the description of N in statistical terms, with the hope that these statistical characteristics of N may be used for the prediction of the more important statistical parameters describing the transmission loss. At a given instant N will vary considerably with height above the surface, and, to a lesser extent, with distance along the path. If, however, we average the values of N over a period of an hour, then N will, on typical propagation paths, be more nearly constant along the path at a given height, but will normally decrease monotonically with increasing height above the surface. For the solution of most radio prediction problems, it is permissible and, indeed, desirable to average N over still longer periods of time in order to obtain mean conditions useful in radio systems design. Thus, for predicting the diurnal, seasonal, or geographical variations of the median transmission loss, we may further average the values of \overline{N} as determined from day to day over a period of many years for a particular hour of the day, month of the year and geographical location. The resulting values \overline{N} will vary still less along the path and, on most paths, \overline{N} , for a given time of day and season of the year, may be taken to be a function only of the height, h - h, above the earth's surface where h represents the height, expressed in kilometers, above sea level, while hs represents the height of the surface above sea level. There will be some paths for which N will also vary appreciably in the horizontal plane; examples are paths with one terminal over land and the other over the sea, and these will undoubtedly require special treatment. Since the average values N tend for most paths to be very nearly horizontally homogeneous, it should only be necessary to know the vertical profile of $\overline{\overline{N}}$ at one point along the path for a successful prediction of the median transmission loss at a particular time of day and season of the year. For very long paths on which N does vary appreciably along the path, we may base our radio predictions on the two N profiles at the intersections of the two radio horizons with the great circle path.

From the above discussion it appears to be desirable to study these N profiles, and it will be convenient in the following analysis to omit the superscripts and simply let N(h) denote these long-term average values. The most generally reliable single parameter for the description of the profile as it affects radio propagation is the difference, ΔN ,

in the values, N_1 , at a height of one kilometer above the surface and N_s the value at the surface:

$$\Delta N \equiv N(h_s + 1) - N(h_s) \equiv N_1 - N_s \tag{34}$$

Note that $\triangle N$ is a negative quantity. Most of the diurnal, seasonal, and geographical variations in propagation between antennas at heights of less than one kilometer above the surface may be predicted on the assumption that N(h) decreases linearly with height above the surface up to a height of one kilometer:

For radio propagation predictions at the higher frequencies above, say, 50 Mc, the above assumption of linearity for the initial decrease of N(h) with height is not adequate for some times of the day or for some geographical locations; in some of these special cases there may be ducting with a resulting substantial increase $\frac{59}{}$ in the transmission loss for paths just short of the radio horizon, and a very large decrease $\frac{60}{}$ in the transmission loss on paths just beyond the radio horizon. Since these appreciably non-linear profiles occur in only a very small percentage of all cases, $\frac{61}{}$ we will not consider them further in this survey report. Furthermore, although the principles of duct propagation are well understood, $\frac{62}{}$ $\frac{63}{}$ $\frac{64}{}$ there are nevertheless no very satisfactory formulas for predicting the transmission loss for the large variety of non-linear profiles typically encountered in practice.

The assumption of a linear profile makes possible the introduction of a great simplification in radio propagation predictions. Thus it has been shown $\frac{65}{}$ that the behavior of radio waves in an atmosphere with a linear gradient is the same as that expected with no atmosphere for an earth with effective radius a \equiv ka' where a' denotes the actual earth's radius, expressed in kilometers, and a is defined by:

$$\frac{1}{a} = \frac{1}{ka'} = \frac{1}{a'} + \frac{\Delta N}{(1 + N_g \cdot 10^{-6})}$$
 (36)

Bean and Meaney $\frac{66}{}$ demonstrate that there is a high correlation between the monthly median transmission loss and the monthly median values of ΔN , and give maps of the monthly median values of ΔN for the United States for several months of the year. The values of ΔN

must be determined from radio-sonde observations and are, as a consequence, not as readily available at all hours of the day nor for as many geographical locations as the surface values, N_s . Fortunately, ΔN may be predicted $\frac{67}{}$ with quite good accuracy from N_s and, in the absence of observations of ΔN , the following empirical formula may be used to determine the predicted value $\Delta N'$:

$$\Delta N' = -7.32 \exp\{0.005577 N_s\}$$
 (37)

If now we combine (35) and (37), we have the following expression for the initial behavior of N(h) in what will be referred to as the CRPL Standard Radio Refractivity Atmosphere as recently proposed and studied by Bean and Thayer: $\frac{67}{}$

$$N(h) = N_s - (h - h_s) 7.32 \exp \{0.005577 N_s\}$$
 $[h_s \le h \le h_s + 1] (38)$

Note that the only parameters in (38) are the surface refractivity, N_s , and the height, h_s , of the surface above sea level.

Note that (37) may also be used to predict N_s in terms of known values of ΔN :

$$N'_{s} = 412.87 \log_{10} (-\Delta N) - 356.93$$
 (39)

When values of ΔN and of N_s are both available, and when the actual value of N_s differs from the value predicted by (39), it is better to use N_s^t for predictions when $\psi < 3^\circ$ rather than the actual value of N_s ; in other words, ΔN is slightly better than N_s as a predictor of propagation conditions for small values of Ψ . On the other hand, it is at present easier and usually also more accurate to predict Ns for some particular time of day, season of the year and geographical location, and then use (37) for determining $\Delta N'$, than it is to use the available maps $\frac{66}{}$ directly for the prediction of ΔN . It should also be noted that, even when ΔN is available, N_s is a better predictor $\frac{67}{}$ of the bending at high elevation angles: $\psi > 3^{\circ}$. It is expected that the recent study program 68/ proposed by the International Radio Consultative Committee (CCIR) will tend to expedite the gathering of the data on AN required for the development of suitable prediction methods; however, it is unfortunate that emphasis was given in that proposal to the gathering of data at only two hours of the day, 0200 and 1400 U.T., since it is precisely the large diurnal variation of AN, occurring at many locations, which is most difficult at present to predict with adequate accuracy.

The effect on N_s of the height of the surface above sea level may be determined from the relation:

$$N_{s} = N_{o} \exp(-c_{s}h_{s}) \tag{40}$$

where $c_s = 0.1057/kilometer = 0.1701/statute mile = 0.03222/thousand feet. Bean and Horn <math>\frac{69}{}$ give maps which may be used to estimate the value of N_0 averaged throughout the day for the months of February and August at any geographical location in the world, together with a map of the annual range of N_s . A comprehensive climatological study of N_s for the United States is in preparation at C. R. P. L.; this study will make available charts useful for the prediction of N_0 (and thus of N_s by means of (40) above) at 0200, 0800, 1400 and 2000 for February, May, August and November and, in addition, gives the detailed statistical characteristics of N_s at several representative weather stations in the United States.

Summarizing the above, we see that the average value of N up to a height of one kilometer above the surface may be predicted preferably, when ψ is small, in terms of a measured mean gradient, ΔN , or, alternatively and with only slightly less accuracy, in terms of the mean value of the surface refractivity, N_g . Above one kilometer, N decreases exponentially with height, and Bean and Thayer $\frac{67}{}$ give the following formulas for N(h) in this range:

$$N(h) = N_1 \exp[-c_i(h - h_s - 1)]$$
 $(h_s + 1 \le h \le 9 \text{ km})$ (41)

$$c_i = \frac{1}{8 - h_g} \log_e (N_1/105)$$
 (42)

$$N(h) = 105 \exp[-0.1424 (h - 9)]$$
 $h \ge 9 \text{ km}$ (43)

Since the constant c_i depends only on N_s and h_s (see Table 5.1 below), it appears that the mean atmosphere may be described in most cases in terms of these two parameters or, alternatively, in terms of h_s and N_s' when ΔN is known. Note that h_s influences the description of the atmosphere [see (42)] only in the range from one kilometer above the surface to 9 km above sea level, and then only slightly for the range of values of h_s normally encountered in practice; consequently, we may, for practical purposes, consider that the atmosphere is well defined by the single parameter N_s . The success of this model in

Table 5.1

Constants for the CRPL Reference Atmospheres

| N · s | h s feet | a' MILES | - ΔN [†] | k | a MILES | c i per kilometer |
|----------|----------------|-------------|-------------------|---------|------------|-------------------------|
| 0 | 0 | 3960 | 0 | 1 | 3960.00 | 0 |
| 200 | 10,000 | 3961.8939 | 22.3318 | 1.16599 | 4619.53 | 0.106211 |
| 250 | 5,000 | 3960.9470 | 29.5124 | 1.23165 | 4878.50 | 0.114559 |
| 301 | 1,000 | 3960. 1894 | 39.2320 | 1.33327 | 5280.00 | 0.118710 |
| 313 | 700 | 3960.1324 | 41.9388 | 1.36479 | 5404.57 | 0.121796 |
| 350 | 0 | 3960 | 51.5530 | 1.48905 | 5896.66 | 0.130579 |
| 400 | 0 | 3960 | 68.1295 | 1.76684 | 6996.67 | 0.143848 |
| 450 | 0 | 3960 | 90.0406 | 2.34506 | 9286.44 | 0.154004 |

predicting the bending of radio waves has been examined by Bean and Thayer 67/ and leaves little to be desired except in the small percentage of cases involving non-linear profiles.

For some mathematical analyses of radio propagation, the above-described CRPL Model Radio Refractivity Atmospheres* have the undesirable characteristic of having discontinuities in the gradient at one kilometer above the surface and at 9 km above sea level. The following exponential model is free of this defect and, although it does not fit the meteorological data above one kilometer as well as the CRPL Model Radio Refractivity Atmospheres, it does nevertheless provide a representation useful for many applications:

$$N(h) = N_s \exp[-c_e(h - h_s)]$$
 (44)

^{*} The CRPL Model Radio Refractivity Atmospheres have arbitrary values of N_s , h_s , and ΔN ; in the CRPL Standard Radio Refractivity Atmospheres N_s and ΔN are related by (37) and (39), but h_s is arbitrary; and in the CRPL Reference Radio Refractivity Atmospheres N_s , ΔN_s , and h_s have the values given in Table 5.1.

The constant c_e is defined in terms of ΔN and N_s :

$$\exp(-c_e) = 1 + \frac{\Delta N}{N_g} \tag{45}$$

Table 5.2 gives the values of c_e for several values of ΔN for the particular case in which N_s' is determined by (39); the value $N_s = 313$ represents the average of the observed values of N_s in the United States.

| N(h) = 1 | Ng | e xp[- | ce(h | - h |)] |
|----------|----|---------------|------|-----|----|
|----------|----|---------------|------|-----|----|

| | | c |
|---------|-----------------------------|---------------|
| ΔΝ | N [†] _S | per kilometer |
| 0 | 0 | 0 |
| 22.3318 | 200 | 0.118400 |
| 29.5124 | 250.0 | 0.125625 |
| 30 | 252.9 | 0.126255 |
| 39.2320 | 301.0 | 0.139632 |
| 41.9388 | 313.0 | 0.143859 |
| 50 | 344.5 | 0.156805 |
| 51.5530 | 350.0 | 0.159336 |
| 60 | 377.2 | 0.173233 |
| 68.1295 | 400.0 | 0.186720 |
| 70 | 404.9 | 0.189829 |
| 90.0406 | 450.0 | 0.223256 |

COMPARISON OF RAYS IN THE CRPL REFERENCE REFRACTIVITY ATMOSPHERES-1958 AND THE 4/3 EARTH ATMOSPHERE

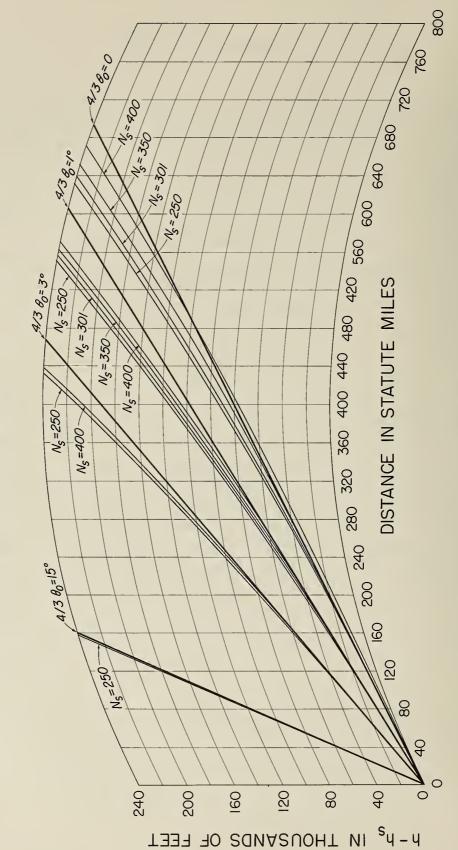


Figure 47

Fig. 47 compares the paths followed by radio rays leaving the earth at several selected elevation angles, ψ , for several of the CRPL reference atmospheres with the paths expected in a four-thirds earth atmosphere. Thus, the graph paper used for tracing these rays was so designed that they are straight lines for the linear gradient atmosphere corresponding to an effective earth's radius of 5,280 statute miles, i.e., (4/3), 3,960 miles. Note the very large departures at large heights of the rays in all of these representative atmospheres from the rays in the usually assumed four-thirds' earth atmosphere. Note that there are large departures from the four-thirds' earth atmosphere at large heights for N_s = 301, even though the bending in this atmosphere is correct for heights h - hg less than one kilometer. An appropriate allowance for this difference in bending is made in the tropospheric curves presented in subsequent sections of this report, but no such allowance was made in preparing Figs. 3, 4 and 6. A good estimate of the correction which should be made to the altitudes of the contours on Figs. 3 and 4 is simply the difference in heights, at the appropriate range, between the 4/3 earth rays and the N_s atmosphere rays. For a detailed discussion of such corrections with appropriate graphs, see a recent report by Rice, Longley and Norton. 10

The Bean and Thayer report gives the elevation angle error $\epsilon \equiv \psi - \psi_0$ as a function of electrical path length, R_e , for rays in the reference atmospheres as well as $\Delta R_e \equiv R_e - R_o$.

$$(R_1/2)$$
 $R_e = \int_0^1 n dR = ct_e$ (46)

where c is the velocity of light in a vacuum and te is the time of transit along the ray path.

5.1 The Influence of Tropospheric Bending on Ionospheric Propagation

The bending of the radio waves by the troposphere has the effect of extending the range of ionospheric propagation and, although the effect is small, it is not negligible near oblique incidence. Table 5.3 shows the influence of N_g on R_1 , d_1 , and $\cos \phi$ for ionosphere heights h=70, 90, 110, 225, 350, and 475 km. The calculations were made by tracing rays in the CRPL reference atmospheres with $N_g=0$, 301 and 400. Since the variations with N_g are comparatively small, it would

Table 5.3: h = 70 km

| | | ф воз | | • 144319 | • 144320 | • 144323 | • 144333 | .144350 | • 144374 | • 144404 | • 144441 | • 144486 | .144536 | • 144594 | • 144658 | • 144807 | • 145081 | •145670 | • 146424 | .147340 | • 149646 | • 152558 | • 157985 | • 164564 | • 174892 | • 206627 | 4 | •326100 | •411481 | • 577069 | • 788447 |
|---------|---------|-------------------|-------|-----------|-----------|------------|----------|-----------|-----------|-----------|----------|-----------|-----------|----------|-----------|----------|----------|---------|----------|----------|----------|------------|-----------|-----------|-----------|----------|-----------|-----------|-----------|-----------|-----------|
| M - 400 | l co | d- | miles | 130881 54 | 130186 54 | 129500 54 | 8156 | 26848 5 | 25575 | 124336 54 | 23130 | 21954 | 120809 54 | 19690 | 66 | 7 | 32 | | ١.+ | | 33056 | 868558 53 | 83948 | 711348 53 | 629690 53 | 5 | 82478 | 264968 53 | 198145 53 | 124489 53 | 680906 52 |
| | | R ₁ /2 | miles | 5 | N | 651804 53 | 45083 5 | 3 5 | | 625984 53 | 9952 5 | 4075 5 | 5 | 2754 5 | 7294 5 | 7 5 | 71708 5 | 3 | 5 | 5 5 | 7084.1 5 | 4 5 | 396365 53 | N | 19468 5 | 5 | 97128 5 | 7 | 108695 53 | 762154 52 | 553508 52 |
| | | φ soo | | • 144985 | • 144986 | . 144989 | • 144999 | • 145016 | 3 | • 145070 | • 145107 | .145151 | .145201 | • 145259 | • 145323 | • 145471 | • 145743 | 7 | 14 | • 147992 | • 150287 | • 153187 | • 158591 | • 165145 | • 175437 | . 207082 | 442 | • 326370 | •411679 | .577183 | • 788494 |
| N - 301 | 1 100 | d ₁ | miles | 124865 54 | 124339 54 | 123816 54 | 783 | 21766 | 120763 54 | 2 | | | | | 115047 54 | | 3 | | ı | | | 859148 53 | | 706578 53 | 626450 53 | 9160 | 1746 | 64710 | 198029 53 | 124453 53 | 680800 52 |
| | | R ₁ /2 | miles | 628625 53 | 5 3665 | 623384 53. | | 613131 53 | 08118 | | 98313 5 | 593518 53 | 88791 5 | 84132 5 | 79538 5 | 7 5 | 57475 5 | 36771 5 | 17255 5 | 98799 5 | +702 5 | 4,33914 53 | 93083 5 | N | 17856 5 | 44755 5 | 196770 53 | 40003 5 | 108642 53 | 62006 5 | 553476 52 |
| | | φ soo | | • 147007 | •147007 | .147010 | .147020 | .147036 | .147060 | .147090 | •147126 | • 147169 | •147219 | •147276 | • 147339 | •147485 | • 147753 | .148331 | .149071 | .149971 | •152235 | .155096 | .160433 | • 166911 | •177094 | 847 | .245413 | .327194 | .412287 | .577531 | • 788639 |
| 2 | 100 | ф | miles | 6852 | 9 29 59 | 6063 5 | 15279 5 | 500 5 | 13 | 959 5 | 961 | 439 5 | 587 5 | 94.0 5 | 198 ء | 731 5 | 5 70 5 | 073 5 | 362 5 | 5 66 5 | 508 5 | 837342 53 | 346 5 | 550 5 | 320 5 | 288 5 | 728 5 | 992 5 | 703 | 4352 5 | 80510 5 |
| | | R ₁ /2 | miles | 88558 | 86581 | 84611 | 80691 | 76797 | 572930 53 | 06069 | 65277 | 61490 | 57729 | 53996 | 50288 | 42953 | 32148 | 14666 | 97832 | 81640 | 51128 | 23022 | 85085 | 51773 | 13660 | 42840 | 95780 | 39660 | 8492 | 61588 | 3385 |
| | | -) | m.r. | • | • | | | • | 0 • 5 | • | • | • | | • | 0 | 2. | 5. | 0 | 5 | 0 | 0 | • | 5 | 0 | 0 | 50 | 20000 | 00 | 0 | | 0.006 |
| | | | | | | | | | | | | | | | | | 8 | 4 | | | | | | | | | | | | | |

588558 53 is to be read 0. $588558 \times 10^3 = 588.558$ miles

Table 5.3: h = 90 km

| | ф воз | • 163948 | • 163948 | • 163951 | • 163960 | • 163974 | • 163995 | • 164022 | • 164054 | • 164093 | • 164137 | • 164188 | • 164244 | • 164374 | • 164614 | .165130 | • 165792 | • 166596 | • 163627 | • 171200 | •175024 | • 181916 | • 191253 | .220478 | •255514 | • 334460 | • 417671 | • 580629 | |
|---------------|----------------------------|------------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|------------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|---|
| N = 400 | d | 146615 54 | | 145234 54 | | 142579 54 | 141304 54 | 140062 54 | | 137672 54 | | 135398 54 | | 132174 54 | | 124433 54 | + | | 108532 54 | 101818 54 | | 850888 53 | 761498 53 | 592876 53 | 77668 5 | 335116 53 | 252026 53 | 159068 53 | |
| | $R_1/2$ miles | 739332 53 | 5 | 732426 53 | 5 | 9151 | 712775 53 | 706565 53 | 700515 53 | 694617 53 | 88862 | 683243 53 | 677752 53 | | | | 606622 53 | | | 515380 53 | | 431879 53 | 48 5 | 303653 53 | 5 26 | 177752 53 | 138671 53 | 976845 52 | |
| | ф воз | .164530 | .164530 | • 164533 | . 164542 | • 164556 | • 164577 | • 164604 | • 164636 | • 164675 | • 164719 | • 164709 | • 164825 | • 164955 | • 165194 | . 165708 | • 166367 | • 167169 | • 169192 | • 171757 | • 176564 | .182438 | • 191747 | . 220901 | • 255873 | • 334721 | • 417856 | . 580 741 | |
| $N_{s} = 301$ | d miles | 140534 54 | 140007 54 | | 138451 54 | 137432 54 | 136427 54 | 135436 54 | 134459 54 | 133496 54 | 132546 54 | | 130684 54 | 128869 54 | 126231 54 | 122033 54 | | 114280 54 | 5 | | 921768 53 | 845662 53 | | 591206 53 | 76786 | 34196 | | 159022 53 | |
| | R ₁ /2 miles | 708924 53 | 706293 53 | 1 5 | 698511 53 | 693415 53 | 688390 53 | 683437:53 | | 9 | 68985 | | 659673 53 | 650600 53 | | | 596546 53 | | | 510401 53, | 467239 53 | | 85544 | 302828 53 | 46466 | 177599 53 | 5 | 976655 52 | • |
| | φ soo | ,166304 | .166304 | .166306 | •166315 | • 166330 | .166350 | .166377 | • 166409 | • 166447 | •166491 | .166540 | • 166596 | • 166724 | .166960 | .167469 | •168120 | .163913 | •170915 | .173452 | •178211 | .184029 | .193256 | .222196 | .256972 | .335520 | | .581086 | |
| 0 = 8 s | $d_{\frac{1}{n}}$ | 132327 54 | 319 | 1537 | 52 | 129972 54 | 9616 | 128425 54 | 659 | | | 89 | | 160 | 4260 | 117423 54 | 3986 | 10663 | 352 | | 4212 | 832262 53 | 48280 | 586604 53 | 312 | 33886 | 51460 | 3890 | |
| | R ₁ /2 miles | 667884 53* | 5 9069 | 63935 5 | 60010 5 | S | 2231 5 | 648377 53 | 9 9 9 | 38 5 | 36954 5 | 633194 53 | 594 | 622052 53 | 1121 5 | 93366 5 | 76188 5 | 9579 5 | 8039 5 | 98669 5 | 458477 53 | 22594 5 | 2 0 2 2 | 00551 5 | 245253 53 | 177164 53 | 38410 | 976108 52 | |
| | ψ m.r | 0.0 | 0.5 | 1.0 | | 3.0 | 0 • 7 | 5.0 | 0 • 9 | 7.0 | > ® | 0.6 | 0 | 12.0 | 10 | 0 | 25.0 | \circ | 0 | 20.0 | 5 | 80.0 | 100.0 | 150.0 | 200•0 | . 300.0 | 0.004 | 0.009 | |
| | | | | | | 100 | | | | | | | | | 8 | 35 | | | | | | | | 1 | | | | | |

* 66788453 is to be read 0. $667884 \times 10^3 = 667.884$ miles

| = 400 | d ₁ cos φ | 1111cb | ± 1 | † · | 54 | 54 | 156538 54 • 181335 | 54 | 154017 54 •181378 | 152805 54 • 181407 | 151622, 54 • 181442 | | 149340 54 • 181527 | 148238 54 • 181578 | 146102 54 • 181695 | | 138302 54 • 182375 | 133891 54 • 182970 | 129752 54 • 183695 | 122132 54 • 185527 | 115226 54 • 187855 | 105947 54 • 192234 | 53 | 882178 53 • 206186 | 697892 53 •233391 | 568162 53 • 266544 | 403140 53 • 342542 | 304796 53 • 423716 | 193226 53 • 584135 | |
|---------|----------------------|------------|--------|-----------|----------|-----------|--------------------|-----------|-------------------|--------------------|---------------------|------------|--------------------|--------------------|--------------------|-----------|--------------------|--------------------|--------------------|--------------------|--------------------|--------------------|-----------|--------------------|-------------------|--------------------|--------------------|--------------------|--------------------|--|
| Z | R ₁ /2 | | 0 0 | 7 0 | 55 | 53 | | 53 | 778522 53 154 | 772459 53 15 | 766546 53 15. | | 755135 53 149 | 53 | 738940 53 140 | 53 | 699940 53 138 | 677885 53 13. | 657194 53 129 | 619108 53 12 | 584600 53 11 | 538266, 53 10 | 53 | 449883 53 88 | 53 | 294506 53 56 | 5 53 | 168211 53 30 | 119023 53 193 | |
| | oso o | | 101033 | • | • | • | 4 • 181857 | • | + • 181900 | + • 181929 | + • 181964 | + • 182004 | • | + 182099 | + •182216 | | + • 182894 | + • 183487 | + •184210 | + • 186037 | + • 188358 | + • 192725 | • | 3 • 206641 | 3 .233788 | 3 • 266886 | 3 .342795 | 3 • 423906 | 3 • 584246 | |
| N = 301 | dı | 20 077701 | • | | - 1 | | 151345 54 | 150339 54 | 149346 54 | 148367 54 | 147401 54 | 146447 54 | 145506 54 | 144576 54 | 142751 54 | | 135857 54 | 131831 54 | 127991 54 | 120802 54 | 114188 54 | | 971920 53 | 878252 53 | 696020 53 | 567150 5 | 402760 53 | 304620 5 | 193169 53 | |
| | R ₁ /2 | Salling Co | - 1 | | | 770263 53 | 765161 53 | 760128 53 | 755164 53 | 750267 53 | 745436 53 | 740667 53 | 735960 53 | 731313 53 | | 708903 53 | 687717 53 | 66 1861 53 | 648391 53 | ł | 579417 53 | 534541 53 | 494590 53 | 447930 53 | 357490 53 | 294010 53 | 214284 53 | 168130 53 | 119000 53 | |
| | os o | 1000 | 104010 | 1037.37 | • 182424 | .183442 | .183455 | • 183473 | .183497 | •183526 | •183560 | • 183600 | • 183644 | •183694 | .183810 | .184023 | • 184482 | 0/0581. | .185786 | • 187596 | • 189897 | 194527 | ,199546 | • 208034 | ,235005 | • 56 / 934 | .343570 | 685575 | • 584586 | |
| 0 = Z | 4. | 12.010.67 | | 74 710971 | 5 T 4 | 529 | 143747 54 | 142970 54 | | | 140664 54 | | 148 | 138396 54 | 9 | 701 | 131111 54 | 679 | 124242 54 | 117781 54 | 111717.54 | 331 | 485 | 867754 53 | 810 | 564280 53 | 401672 53 | 304110 53 | 193007 53 | |
| | R ₁ /2 | 720070 02* | - 1 | 72,000 52 | | 731071 53 | 727164 53 | ω | | ı | 711748 53. | L | 704167 53 | | | 681933 53 | 663.984 53 | l | 629649 53 | _ | _ | ١. | _ | 2 | | 292603 53 | | 167896 53 | 118933 53 | |
| | 3 | iller. | | | D • T | 2.0 | 3.0 | 0 • 7 | 5.0 | 0.9 | 0.7 | 8.0 | 0.6 | 10.0 | 12.0 | | 50.0 | 25.0 | 30.0 | 0.04 | 20.0 | 65.0 | 0.08 | 10000 | 150,0 | .0002 | 3000 | 0.004 | 0.009 | |

* 738948 53 is to be read 0.738948 $\times 10^3 = 738.948$ miles

Table 5.3: h = 225 km

| | | 0 = N | | | $N_s = 301$ | | | $N_{\rm g} = 400$ | |
|--------|----------------------------|-------------------------|---------|--------------------|--|----------|--------------------|--------------------|---|
| ÷ + | R _l /2 miles | d _l miles | d soo | $ m R_{l}/2$ miles | $\frac{\mathrm{d}_{\mathrm{l}}}{\mathrm{miles}}$ | ф воэ | $ m R_{l}/2$ miles | $d_{ m l}$ $miles$ | ф воз |
| | */1 01:/0: | 007400 | 00000 | | | 70230 | ш | 1 | ,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,, |
| 0 0 | -1 | 074/07 | 026020 | | 2157.0 54 | 75075 | | | 0 2 2 / 4 / 4 |
| 0 0 | | | フログイ | | 21264U 24 | 178167 | 113330 34 | 87/17 | 0/4/07 |
| 1.0 | | | 5892 | | - 4 | •257828 | _ | 221040 54 | • 257476 |
| 2.0 | 105390 54 | | .258927 | 109506 54 | 214080 54 | .257834 | 112313 54 | 219692 54 | .257481 |
| 3,0 | 104997 54 | 205066 54 | .258936 | 108994 54 | 213056 54 | . 257843 | 111656 54 | 218378 54 | .257490 |
| 4.0 | 104607 54 | 204284 54 | •258949 | 108489 54 | | • 257855 | 111015 54 | | . 257503 |
| 5.0 | | | .258965 | 107989 54 | 211046 54 | .257872 | 110390 54 | 215846 54 | .257519 |
| 0.9 | 103829 54 | 202728 54 | .258985 | | 210060 54 | .257892 | 109780 54 | 214628 54 | .257539 |
| 7.0 | 103443 54 | 201956 54 | .259008 | 107009 54 | 209086 54 | .257915 | 109184, 54 | 213436 54 | .257563 |
| 8.0 | 103058 54 | | .259035 | | | . 257942 | | | .257590 |
| 0 6 | 102674 54 | 200420 54 | .259066 | 106051 54 | 207170 54 | .257973 | 108032 54 | 211132 54 | • 257621 |
| 10.0 | | 199655 54 | .259100 | | 1 | .258007 | | | .257655 |
| 12.0 | 101533 54 | 198136 54 | .259179 | 104653 54 | 204374 54 | • 258087 | 106392 54 | 207852 54 | • 257735 |
| 0.61 | | | .259325 | | | .258233 | 04840 | 04748 | .257882 |
| 20.0 | | | ,259639 | 101122 54 | 197312 54 | • 258549 | 102408 54 | 199885 54 | • 253198 |
| 0.5 | 967396 53 | | .260043 | 0 | 193140 54 | • 258955 | 100129 54 | 95327 | .258605 |
| 30.0 | 949620 53 | 184993 54 | .260536 | 970279 53 | .+ | • 259450 | 979709 53 | 191011 54 | .259100 |
| 40.0 | | 178098 54 | .261785 | | | • 260705 | | 182946 54 | .260358 |
| 50.0 | 882139 53 | 171488 54 | .263382 | 896308 53 | 174324 54 | •262310 | 902087 53 | 175481 54 | • 261964 |
| 65.0 | 835250 53 | 162098 54 | •266413 | 846311 53 | 164313 54 | • 265355 | 5 | | .265014 |
| 80.0 | | 153316 54 | .270181 | 800264 53 | 155085 54 | • 269140 | 803572 53 | 155748 54 | • 268805 |
| 100.0 | | | .276293 | 744298 53 | | • 275279 | 746732 53 | 144342 54 | .274953 |
| 150.0 | 623696 53 | 119559 54 | .296436 | 627364 53 | 120300 54 | • 295502 | 628644 53 | 120559 54 | .295202 |
| 200.0 | 534648 53 | 101482 54 | .322278 | 536832 53 | 5 | .321435 | 537580 53 | 102080 54 | .321163 |
| 300.0 | 409770 53 | 757298 53 | .385379 | | 759200 53 | • 384709 | 410987 53 | 759844 53 | • 384494 |
| 400,00 | 5 5 | | .456640 | 30504 | 588730 53 | •456115 | 330651 53 | | • 455946 |
| o,009 | 239192 53 | 581496 53 | .603728 | 5 | | • 603409 | 239370 53 | 381930 53 | • 603306 |
| 0.006 | 176605 53 | 212096 53 | 169664 | 176636 53 | 212194 53 | • 799554 | 176646 53 | 212226 53 | • 799510 |
| | | | | | | | | | |

*106179 54 is to be read 0. $106179 \times 10^4 = 1061.79$ miles

Table 5,3: h = 350 km

| | ф soo | •317318 | •317318 | | • 317324 | •317331 | • 317341 | • 317353 | •317369 | •317387 | • 317409 | .317433 | •317460 | • 317522 | •317637 | | • 318202 | •318590 | • 319576 | • 320838 | • 323241 | • 326242 | • 331140 | • 347525 | • 369034 | .423355 | • 486896 | • 622422 | • 807780 |
|---------|----------------------------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|
| N = 400 | d1 miles | 271916 54 | | 270534 54 | 269186 54 | 267870 54 | | 265334 54 | | 262916 54 | 261746 54 | 260602 54 | 1 | 257304 54 | | 249258 54 | 244634 54 | 240234 54 | 231962 54 | | 213434 54 | 203406 54 | 191048 54 | 164328 54 | 142481 54 | 109633 54 | 867028 53 | | 322284 53 |
| | R ₁ /2 miles | 140690 54 | 140343 54 | 139999 54 | 139325 54 | 138667 54 | 138025 54 | 0 | 136787 54 | ~ | | 135033 54 | 134472 54 | m | | 129360 54 | 127047 54 | 124848 54 | 120713 54 | 116851 54 | 111455 54 | 106451 54 | 100292 54 | 870245 53 | 762577 53 | 603235 53 | 495386 53 | 366096 53 | 273293 53 |
| | ф soo | • 317591 | .317591 | .317592 | •317596 | .317603 | • 317613 | •317626 | •317642 | • 317650 | • 317681 | .317705 | •317732 | •317794 | •317909 | • 318156 | • 318474 | • 318862 | •319846 | .321107 | • 323508 | • 326506 | • 331399 | • 347769 | •369259 | . 423541 | 487047 | .622517 | •807821 |
| N = 301 | d _J miles | 265602 54 | | 264552 54 | 263514 54 | 262490 54 | 61476 | 260474 54 | 259484 54 | 258506 54 | 257538 54 | | 255632 54 | | 251032 54 | | 1 | 238290 54 | 230452 54 | 223022 54 | 1 | 202690 54 | 190509 54 | | 142296 54 | | 866586 53 | 573382 53 | 322234 53 |
| | R ₁ /2 miles | 137533 54 | 37269 | 7008 | 136489 54 | | 135470 54 | | 134474 54 | | 133500 54 | 133022 54 | 132548 54 | | 1 | 128045 54 | 125925 54 | 123876 54 | 119958 54 | 116245 54 | 1 | 106094 54 | 100024 54 | 868760 53 | | 602833 53 | 495183 53 | 366031 53 | 273277 53 |
| | φ soo | .318449 | •318449 | .318450 | •318455 | .318462 | .318472 | ,318484 | .318500 | .318518 | .318539 | ,318563 | •318590 | .318652 | 99/818 | .319013 | •319329 | .319716 | .320697 | .321954 | .324347 | .327335 | . 332213 | .348535 | • 369968 | •424129 | •487522 | .622816 | .807951 |
| 0 = N | d _l miles | | | 255890 54 | | | | | | | | | | | | | | | | | | 758 | 89010 | 89 | 82 | 09296 | 65258 | 72 | 322086 53 |
| | R ₁ /2 miles | | 32873 | 9 | 281 | 888 | 31496 | 901 | | 328 | 29941 | 55 | 29170 | | 1 | 36 | 39 | 121721 54 | | 0 | 8086 | 30 | - | 27 | 6 | 25 | 94569 | | 273231 53 |
| | m.r. | 0.0 | 0.5 | 1.0 | 2.0 | 3.0 | 0.4 | 5.0 | 0 9 | 7.0 | 8.0 | 0.6 | Ю | \sim | PQ. | 20.0 | 25.0 | 30.0 | 0.04 | 50.0 | 0.69 | 80.0 | 10000 | 150.0 | 2000 | 300.0 | 0,004 | 0.009 | 0.006 |
| | | | | | | | | | | | | | | | 8 | 38 | | | , | | | | | 1000 | | | | | |

* 13307154 is to be read 0.133071 \times 10^4 = 1330.71 miles

Table 5. 3: h = 475 km

| | ф soo | 200726 | 267,006 | .364997 | .365001 | .365007 | •365015 | • 365026 | • 365039 | .365054 | .365072 | • 365092 | •365115 | .365167 | • 365263 | .365471 | •365737 | • 366063 | • 366890 | •367950 | • 369972 | • 372502 | • 376644 | .390613 | •409186 | • 457077 | • 514457 | • 639976 | •815523 |
|----------------------|----------------------------|------------|---------|-----------|-----------|-----------|-----------|-----------|-----------|------------|-----------|-----------|---------|-----------|-----------|-----------|----------|------------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|
| N _s = 400 | d _l miles | 212000 E | 1 | 310712 54 | | 308046 54 | 306760 54 | 305506 54 | 304282 54 | 303084 54 | 301912 54 | 300766 54 | | 297456 54 | 294312 54 | 289366 54 | | .280256 54 | 271864 54 | 263984 54 | 252896 54 | 242514 54 | 229586 54 | 201050 54 | 177007 54 | 139381 54 | 111921 54 | 752796 53 | 427280 53 |
| - | R ₁ /2 miles | 143650 56 | - 1 | | 162084 54 | 161426 54 | 160783 54 | 160156 54 | 159544 54 | 158945 54 | l` | 157785 54 | | | | 152085 54 | | 147530 54 | 143335 54 | 139396 54 | 133857 54 | 128676 54 | 122233 54 | 108064 54 | 962142 53 | 779643 53 | 650499 53 | 489141 53 | 368986 53 |
| | ф воз | 368333 | 345223 | .365224 | •365227 | • 365233 | •365241 | • 365252 | • 365265 | • 365281 | • 365298 | • 365319 | 4365341 | • 365393 | • 365489 | • 365697 | • 365963 | .366288 | •367115 | • 368174 | •370194 | .372722 | .376862 | • 390820 | • 409380 | • 457242 | • 514593 | • 640064 | • 815562 |
| N ₈ = 301 | d _l miles | 30.576.6 | | | 303658 54 | 302632 54 | 301618 54 | 300614 54 | | 298642 54 | | 296710 54 | | 293886 54 | | 286702 54 | | 278280 54 | | 262738 54 | | 241766 54 | 229018 54 | 200724 54 | | 139282 54 | 111867 54 | 752594 53 | 427214 53 |
| | R ₁ /2 miles | 160276 64 | | ٦, ٦ | Q. | 158719 54 | 158211 54 | 157710 54 | 157214 54 | 156723 54 | 156238 54 | 155758 54 | | 154345 54 | | 150753 54 | | 146542 54 | 142564 54 | | 133385 54 | 128304 54 | 121951 54 | 107903 54 | 961124 53 | 779170 53 | 650251 53 | 489059 53 | 368966 53 |
| | ф воэ | 36 5.07.7. | 365944 | 365945 | .365949 | .365955 | •365963 | .365974 | .365987 | •366002 | .366020 | .366040 | .366062 | .366114 | .366210 | .366417 | .366683 | • 367007 | .367832 | •368888 | .370903 | .373425 | •377554 | .391480 | 666604* | •457769 | .515028 | •640345 | .815687 |
| 0 = x | d _l miles | 79 062306 | 324 | 9593 | 95140 | 294354 54 | 93568 | 92786 | 292004 54 | | | - 1 | | | | | | 273866 54 | 2 | 259628 54 | 99565 | 239738 54 | | 199773 54 | 6182 | 138983 54 | 104 | 751980 53 | 427016 53 |
| | R ₁ /2 miles | 155762 54* | 5564 | 536 | 972 | 154578 54 | 186 | 3794 | 153404 54 | . + | 52626 | മ | 51852 | Α. | 149935 54 | ~ | | 144334 54 | ~ | 0 | | 127291 54 | 121158 54 | 5 | 7 5 | 77739 5 | 5 | 888 | 04 5 |
| | ψ m.r. | | 200 | 1 0 | 2.0 | 3.0 | O • 4 | 5.0 | 0 • 9 | 7.0 | 8.0 | 0.6 | 10.0 | 12.0 | 15.0 | 20.0 | 25.0 | 30.0 | 0.04 | 50°0 | 65.0 | 80 00 | 100.0 | 150.0 | 200.0 | 30000 | 0.004 | 0 009 | 0.006 |
| | | | | | | | | | | | | | | | 8 | 9 | | | | | | | | | | | | | |

* 155762 54 is to be read 0.155762 \times 10⁴ = 1557.62 miles

| °10 °20 | 0 - 0 c |) 4 m = | 1 60 60 60 | , H W W K | | | | 000- | 1 4 8 61 10 1 10 10 10 10 10 | 24632 32 11683 52 17062 52 79532 51 14113 52 | ~~~~~~~ ~~~~~~~~ | . w w w w | 0000 | ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ |
|----------------|----------------------------|------------------|----------------------------|---------------|---------------------------|--------------------------------------|----------------------------|-------------------------------|---------------------------------|---|--|---------------------------------------|---------------------------------------|---------------------------------------|
| c ₉ | מטטטנ | ורטירטית | <i>,</i> ~ ~ ~ ~ | , ru ru ru ru | , 10, 10, 10, 11 | ה מימים | N N N | win www | , ru ru ru ru | 25307 32 88395 51 1 6 907 52 63717 51 15027 52 | ດດດດເ | יתיתיתית | ימממי | ນທທາ |
| °8 °18 | | | | | | | | | | 71071 51 17735 52 53282 51 17104 52 | | | | |
| °7 °17 | 5429 5 5868 5 2545 5 | 6038.5 7916.5 | 4588 5 8079 5 7991 5 | 1695 5 | 3057 6902 7264 5 | 2774 5 2774 5 3207 5 9306 5 | 5352 5 4019 5 3592 5 | 8780 5 2303 5 3477 5 | 0373 5 5665 5 4198 5 | 25540 52 59699 51 19966 52 46053 51 21492 52 | 4111 5 1834 5 7351 5 8219 5 | 1005 5 5998 5 2925 5 | 0743 0743 1761 54201 5001 | 2315 5 0608 5 1040 5 |
| °6 °16 | מטטטנ | ייטיטיי | 1000 | . ~ ~ ~ ~ ~ ~ | ימימימי | 7 02 02 02 | ת ת ת | N N N N | n n n n n | 27123 52 51888 51 24916 52 40892 51 32484 52 | מטטטט | . സ . സ . സ | ימימימי | $\sigma \sigma \sigma \sigma$ |
| °5 °15 | | | | | | | | | | 37660 52 37660 52 37153 51 87711 52 | | | | |
| °4 °14 | 0000 | | | | | | מ ש ש | * 01 W 01 | | 42487 51 10525 53 34448 51 88308 52 | ~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~ | | . W W W W | ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ |
| °3 °13 | W W W W | 1 -+ -1 ·(| , w w w | | . 4 0 0 6 | , w w w | 5000 | 5000 | | 14694 55 39769 51 93230 52 32533 51 27046 52 | 0 1 1 1 C G | ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ ~ | | 0 0 0 H |
| °2 °12 | ທທານທ | ייטיטיי | י תי תי"ת | ~~~~~ | . ~ ~ ~ ~ ~ | 1000 | NNN | N N N N | . ~ ~ ~ ~ ~ | 37975 51 29675 52 31256 51 15379 52 | տտտտտ | ~ ~ ~ ~ ~ ~ | | ກທທທ |
| °1 °11 | രവംഗം | י מי מי מי | 1000 | יהטיטיה | I N N N N | 10000 | R R R | $\alpha \alpha \alpha \alpha$ | | 36953 51 16986 52 30523 51 10544 52 | מטטטט | | | |
| ÷ i. | 1.0 | 2.0 | 4 0 | 5.0 | 7.0 | 9.0 | 12.0 | 15.0 | 30.0 | • • | 80.0 | 00 | | 600.0 |

| °10 | °20 | 5104 54 | 7185 54 | ı ıv | 69480 53 | 1 10 | 2594 53 | 5 | 5262 53 | י ע | 1978 53 | מ מ | ט יט | | 30113 53 | | | | | | | | | | | | | | 87048 51 | | | | | 9160 52 | 7129 51 | 47756 51 | 3196 51 9891 51 | 1879 51 | 4766 51 | 0925 51 |
|------------------|-----------------|---------|------------|------|----------------|------|---------|------|-------------|-----------------------|------------|-------------|-----------|-----|-------------|------------------|------|-------------|------|------|---------|------|-------|------|------|------|----|------|-------------|-----|------|-------|-----------|---------|---------|------------|--------------------|---------|---------|---------|
| 65 | c ₁₉ | 55 | ა ი 4 ი | 24 | | 54 | 53 5 | 54 4 | 53 | | 34261 53 3 | 53 | . 4s | 53 | | 40 | 53 | 45.24 | ח ה | 53 | 60 1 | 522 | 52 | 52 | | 52 | 52 | 52 | 51 | 52 | | 51 | 5,1 | 52 | 5 51 | 40292 51 4 | 51 | 51 | 51 | 51 |
| ္တ | c ₁₈ | 54 | | 55 | го 10 10 | | 54 | 53 | | 54 | 53 8 | 53 | יה שני | 53 | 5 5 5 | 5 1 1 1 | 52 | 5 5 5 | 5 L | 52 | 53 | 52 | 52 | 52 | 52 | 52 | 52 | 51 | 52857 51 65 | 52 | 52 | 51 | 51 | 51 | 51 | 51 | 51 | 51 | 51 | 51 |
| د2 | c ₁₇ | 54 2 | 53 1 | 54 1 | m = | 53 | 53 | 53 2 | 53 2 | 53 | 52 | 53 | 5 L | 52 | 55 T | 531 | 52 | e G |) in | 52 | 52 | 52 | 52 | 52 | 52 | 51, | 52 | 51 | 44377 51 52 | 52 | 52 | 51 | 51 | 51 | 51 | 51 | 5.1 5.1 | 51 | 51 | 51 |
| 9, | °16 | | | | | | 53 | 55 | | 1.0 | 52 | in ii ii | 53 | 52 | | 1 50 | 5.5 | | 52 | 52 | 52 | 52 | | 52 | 51 | 51 | 52 | 51 | 38556 51 4 | 52 | 52 | 51 | אר ה ר | 51 | 51 | 51 | 5.1 5.1 | 51 | 51 | 51 |
| °S | c ₁₅ | 53 | ა ი ე ა | 53 | n n n | 52 | 53 | 55 | 53 | 5 1 1 1 1 | 52 | r C | 1 50 | 52 | 25 | 52 | 52 | 52 | 52 | | | | | 55 | 51 | 51 | 52 | 51 | 34459 51 | 52 | 53 | 51 | 5 1 | 51 | 51 | | 5.1 5.1 | 51 | 51 | 5 |
| °4 | c ₁₄ | 53 | 5 t | 53 | 52 | 52 | 53 | 52 | | 150 | 52 | e G | 52 | 52 | 52 | 52 | 52 | 52 | 52 | | 52 | 52 | | 52 | 51 | 51 | 52 | 51 | 31555 51 | 52 | 52 | 51 | א ני | | | | | | | |
| °3 | c ₁₃ | | | | | | | | | | | | | | | | | | | | | | | | | | | | 29532 51 | | | | | | | | | | | |
| c ₂ | c ₁₂ | | | | | | | | | | | | | | | | | | | 6620 | 1076 | | | 3272 | | 3400 | | | 28198 51 | | 0783 | | | | | | | | | |
| ე <mark>.</mark> | c ₁₁ | rv i | U R | 5 | n n | , ru | r. | S | | I IU | 5 | יט מ | י יני | ıU | ת ונ | J IV | 5 | וט ת | ט וט | | rU I | | ווי ר | 5 | | | | | 27437 51 | | | | | | | | | | | |
| → ¦ | : | 5 | 1.0 | | 2•0 | 3.0 | | 4.0 | ر د د |)) | 0•9 | , | • | 8.0 | 0 | | 10.0 | 12.0 | • | 15.0 | | 20.0 | 25.0 | | 30.0 | 40.0 | | 50.0 | 0.59 | 0.0 | | 100.0 | 150.0 | • | 200.0 | | 3000 | 0.004 | | 6000 |

appear to be reasonable for most applications to use only the single value $N_s = 301$ in ionospheric calculations, and the remaining tables in this section give the convergence factor c_m for several angles, ψ , and for m = 1 to 20. The values of c_m are given since they involve the geometry in a rather complex way; however, since the ionosphere probably appears rough to the radio waves reflected at the higher layers, c_m is given only for h = 70, 90, and 110 km.

The results presented in this section were all obtained by ray tracing methods, and such methods yield reliable results only when the following two conditions are satisfied: (a) the index of refraction, n, must not change appreciably in a distance equal to a wavelength, and (b) the fractional change in the spacing between neighboring rays in a wavelength along the ray must be small compared with unity. Both of these conditions require that resort must be made to wave solutions of the problem at the lower frequencies. Condition (b) above is always violated at a caustic, and we have shown in Appendix I how such cases may be treated. It might be supposed, since n changes only from about 1.0003 to 1 for a 70 km change in h, that condition (a) would be well satisfied at frequencies even as low as 5 kc; however, it must be remembered that this small change in n actually causes appreciable bending when ψ is small, and we should, instead, require that $\Delta N/N < 0.1(2\pi/\lambda) = 0.002 f_{kc}$ if we are to expect ray tracing to apply. This more stringent requirement is met for f > 65 kc for the $N_s = 301$ atmosphere, and it appears that a wave solution will be required at lower frequencies for a precise treatment of the bending. Until an adequate wave solution becomes available, it would seem that the ray tracing solution here given should be used even for frequencies as low as 10 kc since the alternate assumption of N_s = 0 would undoubtedly yield an even poorer approximation to the actual bending.

5.2 The Total Bending

Above a height of about 70 km, the troposphere no longer bends the radio waves appreciably, and the designation τ has been given to the total bending of radio waves passing entirely through it. Table 5.5 gives τ , the critical range, R_c , and $R_e-(R_1/2)$ as a function of ψ for the CRPL Reference Radio Refractivity Atmospheres. The results provide a convenient means for determining the true elevation angle ψ_o and true range R_o of a satellite at very high heights, say $h > 70~km^*$,

^{*} These results may actually be used without appreciable error whenever $R_{\rm e} > 2~R_{\rm c}$.

Table 5.5

The Total Bending, τ , Critical Range, R_c , and R_e - $(R_1/2)$ for the C.R.P.L. Reference Radio Refractivity Atmospheres

| | | | •• | 25.0 | | | 201.0 | 1 - 1 000 | |
|-------|---------------|-------------------------|-----------------|---------------|-----------|-----------------|---------------|-----------|---------------------|
| | $N_s = 200.0$ | h _s = 10,000 | D' | $N_s = 250.0$ | h = 5,000 | | $N_s = 301.0$ | _ | |
| ψ | т | R _c | $R_e - (R_1/2)$ | т | | $R_e - (R_1/2)$ | т | | $R_{e} - (R_{1}/2)$ |
| m.r. | m.r. | km. | meters | m.r. | km. | meters | m.r. | km. | meters |
| 0.0 | 7.1845 | 200.400 | 62.2 | 9.4504 | 198.752 | 78.1 | 12.1522 | 196.588 | 95.40 |
| 0.5 | 7.1022 | 199.007 | 61.5 | 9.3356 | 197.274 | 77.1 | 11.9871 | 195.049 | 94.1 |
| 1.0 | 7.0211 | 197.632 | 60.8 | 9.2226 | 195.817 | 76.2 | 11.8252 | 193.537 | 92.9 |
| 2.0 | 6.8630 | 194.937 | 59.4 | 9.0025 | 192.966 | 74.3 | 11.5109 | 190.592 | 90.5 |
| 3.0 | 6.7100 | 192.309 | 58.1 | 8.7899 | 190.195 | 72.5 | 11.2090 | 187.744 | 88.3 |
| 4.0 | 6, 5622 | 189.743 | 56.7 | 8,5848 | 187.498 | 70.8 | 10.9194 | 184.987 | 86.1 |
| 5.0 | 6.4193 | 187.235 | 55.4 | 8.3869 | 184.870 | 69.1 | 10.6416 | 182,313 | 83.9 |
| 6.0 | 6.2813 | 184.782 | 54.3 | 8.1962 | 182.308 | 67.6 | 10.3754 | 179.717 | 82.0 |
| 7.0 | 6.1480 | 182.381 | 53.1 | 8.0124 | 179.807 | 66.1 | 10,1202 | 177.190 | 80.0 |
| 8.0 | 6.0194 | 180.027 | 51.9 | 7.8353 | 177.363 | 64.7 | 9.8757 | 174.730 | 78.1 |
| 9.0 | 5.8951 | 177.719 | 50.8 | 7.6645 | 174.973 | 63.2 | 9.6413 | 172.331 | 76.3 |
| 10.0 | 5.7751 | 175.454 | 49.8 | 7.4999 | 172.635 | 61.9 | 9.4165 | 169.988 | 74.6 |
| 12.0 | 5.5470 | 171.046 | 47.8 | 7.1882 | 168.104 | 59.3 | 8.9940 | 165.461 | 71.4 |
| 15.0 | 5.2326 | 164.720 | 45.0 | 6.7602 | 161.641 | 55.7 | 8.4207 | 159.027 | 67.0 |
| 20.0 | 4.7721 | 154.868 | 40.97 | 6.1379 | 151.664 | 50.58 | 7.5998 | 149.131 | 60.57 |
| 25.0 | 4.3771 | 145.812 | 37.44 | 5.6086 | 142.576 | 46.11 | 6.9125 | 140.141 | 55.19 |
| 30.0 | 4.0352 | 137.488 | 34.54 | 5.1539 | 134.283 | 42.31 | 6.3292 | 131.954 | 50.52 |
| 40.0 | 3.4745 | 122.821 | 29.49 | 4.4155 | 119.781 | 36.33 | 5.3943 | 117.660 | 43.17 |
| 50.0 | 3.0364 | 110.443 | 25.80 | 3.8448 | 107.623 | 31.50 | 4.6808 | 105.694 | 37.39 |
| 65.0 | 2.5379 | 95.332 | 21.46 | 3,2016 | 92.847 | 26.22 | 3.8847 | 91.165 | 31.17 |
| 80.0 | 2.1689 | 83.422 | 18.30 | 2.7295 | 81.232 | 22.29 | 3.3051 | 79.749 | 26.50 |
| 100.0 | 1.8076 | 71.142 | 15.30 | 2.2702 | 69.270 | 18.55 | 2.7442 | 67.997 | 21.96 |
| 150.0 | 1.2606 | 51.401 | 10.68 | 1.5794 | 50.052 | 13.02 | 1.9053 | 49,122 | 15.39 |
| 200.0 | 0.9586 | 39.975 | 8.19 | 1.2000 | 38.928 | 9.97 | 1.4464 | 38.201 | 11.80 |
| 300.0 | 0.6379 | 27.600 | 5. 59 | 0.7978 | 26.879 | 6.80 | 0.9611 | 26.375 | 8.03 |
| 400.0 | 0.4693 | 21.156 | 4.26 | 0.5868 | 20.604 | 5.19 | 0.7068 | 20.217 | 6. 12 |
| 600.0 | 0.2912 | 14.699 | 2.94 | 0.3641 | 14.316 | 3.59 | 0.4384 | 14.047 | 4.25 |
| 900.0 | 0.1584 | 10.630 | 2.13 | 0.1980 | 10.354 | 2.59 | 0.2384 | 10.158 | 3.06 |

 $Table \ 5.5$ The Total Bending, τ , Critical Range, R_c , and R_e - ($R_1/2$) for the C.R.P.L. Reference Radio Refractivity Atmospheres

| | $N_s = 350.0$ | h = 01 | | N _s = 400.0 | h = 01 | | $N_s = 450.0$ | h _s = 0' | |
|-------|---------------|---------|-----------------|------------------------|---------|-----------------|---------------|---------------------|---------------------|
| ψ | s T | | $R_e - (R_1/2)$ | s τ | R C | $R_e - (R_1/2)$ | _ | R C | $R_{e} - (R_{1}/2)$ |
| m.r. | m.r. | km | e 1 meters | m.r. | c km | meters | m.r. | c km | e 1 meters |
| 0.0 | 15.6236 | 192,552 | 111.10 | 20.4292 | 189.878 | 128.70 | 27.5747 | 191.865 | 152.6 |
| 0.5 | 15.3818 | 190.840 | 109.40 | 20,0507 | 187.842 | 126.60 | 26.9134 | 189.128 | 149.3 |
| 1.0 | 15.1452 | 189.165 | 107.80 | 19.6822 | 185.861 | 124.30 | 26.2746 | 186.497 | 146.1 |
| 2.0 | 14.6878 | 185.921 | 104.80 | 18.9747 | 182.063 | 120.20 | 25.0635 | 181.540 | 140.2 |
| 3.0 | 14.2512 | 182.809 | 101.80 | 18.3060 | 178.467 | 116.40 | 23.9391 | 176.959 | 134.6 |
| 4.0 | 13.8349 | 179.818 | 99.00 | 17.6748 | 175.055 | 112.70 | 22.8975 | 172.716 | 129.4 |
| 5.0 | 13.4383 | 176.939 | 96.3 | 17.0798 | 171.811 | 109.3 | 21.9340 | 168.775 | 124.7 |
| 6.0 | 13.0605 | 174.163 | 93.8 | 16.5190 | 168.717 | 106.0 | 21.0429 | 165.099 | 120.3 |
| 7.0 | 12.7008 | 171.479 | 91.4 | 15.9905 | 165.759 | 103.0 | 20.2186 | 161.655 | 116.2 |
| 8.0 | 12.3583 | 168,881 | 89.1 | 15.4923 | 162.923 | 100.1 | 19.4552 | 158.414 | 112.4 |
| 9.0 | 12.0320 | 166.361 | 86.9 | 15.0224 | 160.199 | 97.2 | 18.7472 | 155.350 | 108.8 |
| 10.0 | 11.7210 | 163.914 | 84.8 | 14.5787 | 157.574 | 94.6 | 18.0892 | 152.443 | 105.5 |
| 12.0 | 11.1413 | 159.218 | 80.7 | 13.7626 | 152.593 | 89.8 | 16.9049 | 147.026 | 99.3 |
| 15.0 | 10.3651 | 152.608 | 75.4 | 12.6912 | 145.688 | 83.3 | 15.3978 | 139.701 | 91.4 |
| 20.0 | 9.2736 | 142.573 | 67.9 | 11.2232 | 135.398 | 74.3 | 13.4115 | 129.072 | 80.7 |
| 25.0 | 8.3768 | 133.571 | 61.49 | 10.0480 | 126.325 | 66.93 | 11.8776 | 119.907 | 72.17 |
| 30.0 | 7,6271 | 125.452 | 56.08 | 9.0849 | 118.249 | 60.80 | 10.6525 | 111.868 | 65.12 |
| 40.0 | 6.4460 | 111.434 | 47.70 | 7.6000 | 104.501 | 51.28 | 8.8124 | 98.383 | 54.48 |
| 50.0 | 5.5604 | 99.827 | 41.23 | 6.5103 | 93.278 | 44.12 | 7.4944 | 87.524 | 46.62 |
| 65.0 | 4. 5872 | 85.861 | 34.05 | 5.3338 | 79.931 | 36.33 | 6.0979 | 74.749 | 38,22 |
| 80.0 | 3.8877 | 74.969 | 29.02 | 4.5005 | 69.621 | 30.82 | 5.1232 | 64.964 | 32.30 |
| 100.0 | 3,2172 | 63.818 | 24.05 | 3.7104 | 59.138 | 25.51 | 4.2089 | 55.077 | 26.79 |
| 150.0 | 2.2247 | 46.007 | 16.76 | 2.5542 | 42.519 | 17.71 | 2.8853 | 39.506 | 18.50 |
| 200.0 | 1.6860 | 35.747 | 12.83 | 1.9321 | 32.998 | 13.6 | 2.1789 | 30,628 | 14.13 |
| 300.0 | 1.1188 | 24.662 | 8.73 | 1.2803 | 22.745 | 9.21 | 1.4419 | 21.093 | 9.60 |
| 400.0 | 0.8224 | 18,899 | 6.65 | 0.9405 | 17.423 | 7.03 | 1.0588 | 16.153 | 7.33 |
| 600.0 | 0.5099 | 13.128 | 4.62 | 0.5830 | 12.100 | 4.86 | 0.6560 | 11.216 | 5.06 |
| 900.0 | 0.2772 | 9.494 | 3.33 | 0.3169 | 8.749 | 3,51 | 0.3566 | 8.108 | 3.66 |

in terms of its observed elevation angle ψ , its observed radio range R_e and the surface value, N_s , of refractivity at the observing point:

$$\psi_{0} = \psi - \tau [1 - (R_{c}/R_{e})]$$
 (47)

$$R_o = (R_1/2) - R_c \left[\frac{1}{3} - \frac{R_c}{2R_e} \right] \sin^2 \tau$$
 (48)

6. Tropospheric Scatter

The distance d_{Lt} to the radio horizon of a transmitting antenna of height, h_t, above a smooth spherical earth of radius a' and for a linear gradient atmosphere may be determined from:

$$d_{Lt} = \sqrt{2 k a^{\dagger} h_t}$$
 (49)

where k is defined by (36). For the particular CRPL Reference $N_s=30l$ atmosphere, ka'=5,280 miles and, if h_t is expressed in feet and d_{Lt} in miles: $d_{Lt}=\sqrt{2\ h_t}$. When h_t is greater than one kilometer, (49) no longer applies, and reference should then be made to the preceding section or to references 67 and 70; in particular, since the horizon is defined by the ray corresponding to $\psi=0$, Fig. 47 shows the relation over a smooth earth between d_{Lt} and h_t at large heights and for several values of N_s . If we let h_r and d_{Lr} denote the height and distance to the radio horizon for the receiving antenna, then receiving antennas at a distance $d>d_{Lt}+d_{Lr}$ from the transmitting antenna lie below the horizon ray of the transmitting antenna, and it becomes convenient to calculate the transmission loss at such distances in terms of the angular distance, θ . 3/l Over a smooth, spherical earth and in a linear gradient atmosphere, θ may be determined by:

$$\theta = \frac{d - d_{Lt} - d_{Lr}}{ka!} \text{ radians}$$
 (50)

The angular distance, θ , is a particularly convenient parameter for making appropriate allowance for the effects of irregularities in the terrain, and a detailed explanation of methods for calculating the cumulative distribution of transmission loss in propagation over irregular terrain and for a wide range of atmospheric conditions is given in

a recent report by Rice, Longley, and Norton. $\frac{70}{}$ Thus it is shown in that report that the received field on these beyond-the-horizon paths may be considered to consist of diffracted and scattered components. The scattered component may be explained quantitatively for the winter afternoon hours in terms of scatter by a turbulent atmosphere using the mixing-in-gradient hypothesis as the basis for describing the turbulence, i.e., the (r/ℓ_0) $K_1(r/\ell_0)$ correlation function may be used to describe the correlation in the variations of refractive index at points a distance r apart in the atmosphere. $\frac{57}{}$ Some direct experimental evidence for this description of atmospheric turbulence is given in a recent paper by the author. $\frac{58}{}$ The extension of these estimates of the transmission loss for winter afternoons to all-day, all-year values is then done empirically, using the angular distance as a parameter in this empirical analysis.

Fig. 48 shows the values of median basic transmission loss separately for the diffracted wave and for tropospheric scatter as calculated in the manner described in reference 70 for the range of frequencies from 10 to 10,000 Mc and for transmitting and receiving antennas both at a height of 30 feet. If we assume that the short term variations in the scatter fields are Rayleigh distributed, and that the diffracted waves are relatively steady, then we may determine the expected combined median basic transmission loss, $L_{\rm bm}$, in terms of the diffracted wave transmission loss $L_{\rm bd}$ and the median basic scatter transmission loss, $L_{\rm bms}$ as follows:

$$L_{bm} = L_{bd} - R(0.5)$$
 (51)

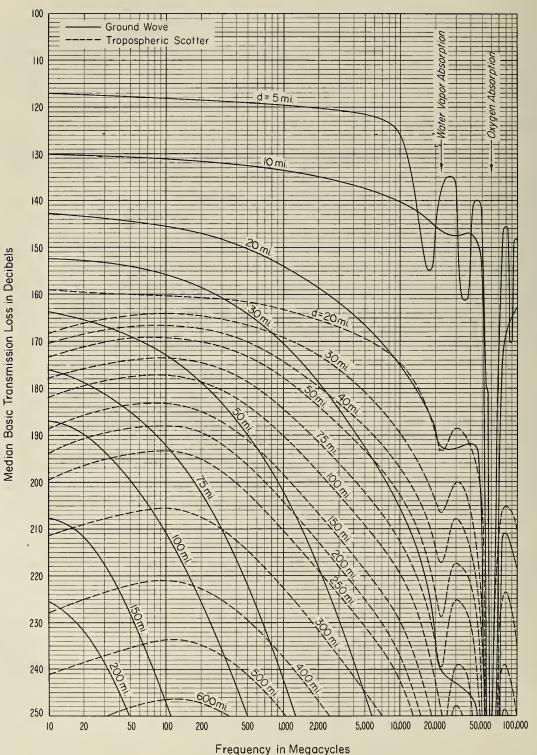
where $K = L_{\rm bd} - L_{\rm bms} + 1.592$ is the ratio in decibels of the average scattered power to the diffracted wave power, and R(0.5) is given graphically and in tables in reference 71. When K is less than -16.5 db, $L_{\rm bm}$ differs from $L_{\rm bd}$ by less than 0.1 db, and when K is greater than 19.5 db, $L_{\rm bm}$ differs from $L_{\rm bms}$ by less than 0.1 db.

Finally, to determine the expected values, $L_b(p)$, of basic transmission loss exceeded by (100 - p) per cent of the hourly medians during a year, we may simply subtract V(p, θ) as given on Fig. 49 from L_{bm} as calculated above from the values shown on Fig. 48.

Fig. 50 shows the influence on the median basic transmission loss at 100 Mc of changing one antenna height while keeping the other antenna height fixed at 30 feet. The values given are for a smooth

MEDIAN BASIC TRANSMISSION LOSS FOR THE GROUND WAVE AND TROPOSPHERIC SCATTER MODES OF PROPAGATION OVER A SMOOTH SPHERICAL EARTH

Over Land σ = 0.005 mhos/meter ϵ = 15 Polarization: Horizontal Transmitting and Receiving Antennas Both 30 Feet Above the Surface



THE VARIANCE OF TRANSMISSION LOSS IN TROPOSPHERIC PROPAGATION

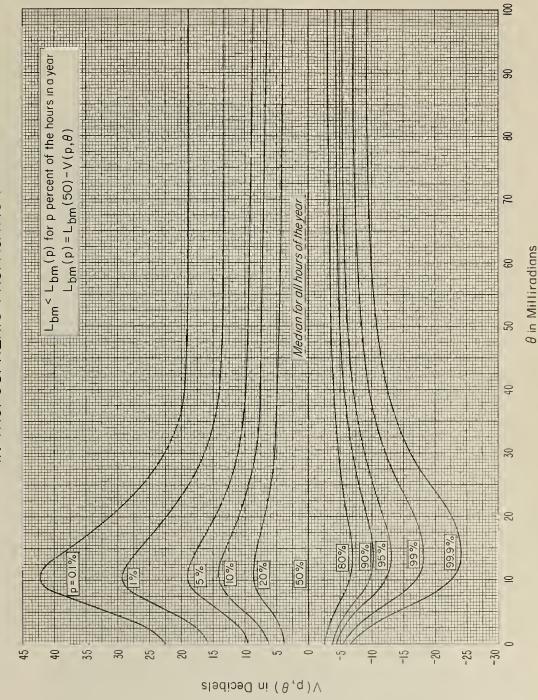


Figure 49

MEDIAN BASIC TRANSMISSION LOSS AT 100 MC

Smooth Spherical Earth and a CRPL Reference N_S=301 Atmosphere Horizontal Polarization Over Land

One Antenna at 30 Feet and the Other Antenna at the Heights Indicated

- Locotion of Maximum Fields (minimum transmission losses)
- (L_{bd}) Ground Wave
- ---- (L_{bms}) Tropospheric Scotter for Winter Afternoon Hours

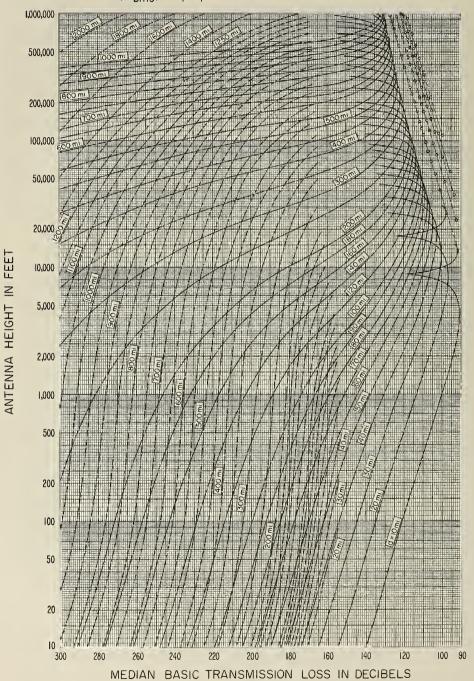


Figure 50

earth and a CRPL Reference $N_s=301$ atmosphere. The first two oscillations of the field are shown for d=10 and 20 miles, but for the other distances only one oscillation is shown. The six points of field maxima are shown for all of the distances as circled points. Note that the total number of maxima to be expected (as a function of range at a given height or as a function of height at a given range) for a particular antenna height is equal to the number of half wavelengths contained in this height; in the present case of 100 Mc, 30 feet represents 6 half-wavelengths; in this connection, see (8), page 11.

The scatter curves on Fig. 50 correspond to the winter afternoon hours, and the reader is referred to reference 70 for curves suitable for translating these values to transmission losses exceeded for several percentages of various periods of time. The scatter loss predictions on Fig. 50 are shown only up to heights just short of the radio horizon since the method of estimation given in reference 70 is not applicable to line-of-sight paths.

7. Point-to-Point Radio Relaying by Tropospheric Scatter

As an example of the method of using transmission loss in systems design, we will consider the problem of estimating the effective maximum range of a radio relay system using tropospheric scatter. As an illustration of typical ranges to be expected, we will assume that the terrain is smooth, and will base our predictions on a CRPL Reference Radio Refractivity Atmosphere with N_c = 301. We will assume that either two 28-foot or two 60-foot parabolic antennas are used at both ends of the path, with their centers 30 feet above the ground and connected in a quadruple diversity system. With these assumptions, we may use the methods described in reference 70 to determine the transmission loss, L(99), which we would expect one per cent of the actual hourly median transmission losses to exceed throughout a period of one year; the use of these one per cent losses implies that the specified service will be available for 99% of the Tables 7.1 and 7.2 give for the 28' and 60' antennas the free space gains $G_t + G_r$, and the path antenna gains as a function of frequency and distance, while Tables 7.3 and 7.4 give L(99) as a function of frequency and distance.

The power required to provide a specified type and grade of service for 99% of the hours may now be obtained from the equation:

$$P_{+} = L_{+} + L(99) + R + F + B - 204$$
 (52)

Table 7.1 Path Antenna Gain in Decibels for 28-Foot Parabolic Antennas 30 Feet Above a Smooth Spherical Earth with a CRPL Model Radio Refractivity Atmosphere Corresponding to $N_{\rm S}=301$

| f _{Mc} | G _t +G _r | G in Decibels | | | | | | | | |
|-----------------|--------------------------------|---------------|-------|-------|-------|-------|-------|-------|--|--|
| | db | d=100 mi. | 150 | 200 | 300 | 500 | 700 | 1000 | | |
| 100 | 33.02 | 33.02 | 33.02 | 33.02 | 32.92 | 32.82 | 32.72 | 32.67 | | |
| 150 | 40.07 | 40.07 | 39.97 | 39.97 | 39.87 | 39.57 | 39.47 | 39.37 | | |
| 200 | 45.06 | 45.06 | 44.96 | 44.86 | 44.66 | 44.36 | 44.16 | 44.06 | | |
| 300 | 52.11 | 52.03 | 51.91 | 51.71 | 51.31 | 50.81 | 50.51 | 50.41 | | |
| 500 | 60.98 | 60.75 | 60.38 | 60.08 | 59.38 | 58.48 | 58.18 | 58.18 | | |
| 700 | 66.83 | 66.35 | 65.83 | 65.23 | 64.33 | 63.23 | 62.83 | 62.93 | | |
| 1000 | 73.02 | 72.12 | 71.22 | 70.42 | 69.22 | 67.92 | 67.32 | 67.22 | | |
| 1500 | 80.07 | 78.32 | 76.57 | 75.87 | 74.27 | 72.57 | 71.67 | 71.57 | | |
| 2000 | 85.06 | 82.48 | 80.66 | 79.26 | 77.16 | 75.46 | 74.36 | 73.86 | | |
| 3000 | 92.11 | 87.71 | 84.21 | 83.41 | 80.91 | 78.51 | 77.61 | 77.31 | | |
| 5000 | 100.98 | 93.28 | 90.18 | 87.68 | 84.68 | 82.18 | 81.28 | 80.98 | | |
| 7000 | 106.83 | 96.53 | 92.83 | 90.03 | 86.83 | 84.12 | 83.33 | 82.93 | | |
| 10000 | 113.02 | 99.32 | 95.22 | 92.52 | 89.02 | 86.02 | 85.22 | 84.92 | | |

Table 7.2 Path Antenna Gain in Decibels for 60-Foot Parabolic Antennas 30 Feet Above a Smooth Spherical Earth with a CRPL Model Radio Refractivity Atmosphere Corresponding to $N_{\rm S}=301$

| f _{Mc} | G _t +G _r | $egin{array}{c} G_{p} & in \ Decibels \end{array}$ | | | | | | | | |
|-----------------|--------------------------------|--|-------|-------|-------|-------|-------|-------|--|--|
| | db | d=100 mi. | 150 | 200 | 300 | 500 | 700 | 1000 | | |
| 100 | 46.26 | 46.16 | 46.16 | 46.06 | 45.86 | 45.46 | 45.26 | 45.16 | | |
| 150 | 53.31 | 53.21 | 53.11 | 52.91 | 52.51 | 52.11 | 51.56 | 51.48 | | |
| 200 | 58.30 | 58.10 | 57.90 | 57.60 | 57.00 | 56.30 | 55.90 | 55.86 | | |
| 300 | 65.35 | 64.95 | 64.45 | 63.95 | 62.85 | 62.05 | 61.70 | 61.90 | | |
| 500 | 74.22 | 73.22 | 72.67 | 71.02 | 70.22 | 68.72 | 68.12 | 68.12 | | |
| 700 | 80.07 | 78.37 | 77.07 | 75.87 | 74.17 | 72.57 | 71.67 | 71.57 | | |
| 1000 | 86.26 | 83.46 | 81.46 | 79.96 | 77.86 | 76.26 | 75.01 | 74.66 | | |
| 1500 | 93.31 | 88.51 | 85.91 | 84.01 | 81.51 | 79.11 | 78.21 | 77.81 | | |
| 2000 | 98.30 | 91.70 | 88.80 | 86.60 | 83.70 | 81.10 | 80.20 | 79.80 | | |
| 3000 | 105.35 | 95.75 | 92.25 | 89.65 | 86.35 | 83.55 | 82.85 | 82.45 | | |
| 5000 | 114.22 | 100.02 | 95.62 | 92.92 | 89.52 | 86.52 | 85.62 | 85.42 | | |
| 7000 | 120.07 | 102.37 | 97.87 | 94.97 | 91.17 | 88.22 | 87.57 | 87.07 | | |
| 10000 | 126.26 | 104.56 | 99.96 | 96.66 | 93.06 | 90.06 | 89.26 | 88.76 | | |

Table 7.3

Transmission Loss L(99) (Corresponding to Fields Exceeded 99% of the time) Expected Between Two 28-Foot Parabolic Antennas at a Height of 30 Feet Above a Smooth Spherical Earth with a CRPL Model Radio Refractivity Atmosphere Corresponding to N_S = 301

| T / | / O O ' | ٠: | Da | ai h | ~1 c |
|-----|---------|-------|----|------|------|
| וע | (フフ, | in in | De | CID | GIP |

| f Mc | d=100 mi. | 150 | 200 | 300 | 500 | 700 | 1000 |
|---------|--------------|--------|--------|--------|--------|--------|--------|
| 100 | 160.21 | 164.33 | 166.95 | 182.31 | 207.79 | 233.29 | 276.57 |
| 150 | 154.20 | 158.56 | 161.57 | 176.52 | 200.95 | 225.91 | 269.52 |
| 200 | 150.24 | 154.72 | 157.92 | 173.08 | 197.38 | 221.76 | 265.56 |
| 300 | 145.32 | 150.27 | 153.55 | 169.33 | 192.88 | 217.45 | 260.36 |
| 500 | 140.10 | 145.87 | 149.64 | 165.94 | 189.74 | 214.13 | 255.99 |
| 700 | 137.39 | 143.81 | 147.98 | 164.83 | 188.55 | 213.08 | 254.05 |
| 1000 | 135.11 | 142.31 | 146.98 | 164.07 | 188.08 | 212.72 | 251.88 |
| 1500 | 133.32 | 141.88 | 146.32 | 164.50 | 188.91 | 213.41 | 254.20 |
| 2000 | 132.54 | 141.36 | 146.76 | 165.62 | 190.08 | 214.46 | 255.50 |
| 3000 | 132.70 | 143.69 | 148.71 | 168.01 | 193.39 | 217.50 | 257.99 |
| 5000 | 134.79 | 145.77 | 152.99 | 173.37 | 199.19 | 223.38 | 264.01 |
| 7000 | 137.37 | 149.86 | 158.13 | 179.75 | 206.98 | 231.32 | 272.13 |
| 10000 | 144.68 | 160.12 | 170.68 | 196.11 | 223.53 | 247.87 | 288.69 |

Table 7.4

Transmission Loss L(99) (Corresponding to Fields Exceeded 99% of the time) Expected Between Two 60-Foot Parabolic Antennas at a Height of 30 Feet Above a Smooth Spherical Earth with a CRPL Model Radio Refractivity Atmosphere Corresponding to N = 301

L(99) in Decibels

| f _{Mc} | d=100 mi. | 150 | 200 | 300 | 500 | 700 | 1000 |
|-----------------|--------------|--------|--------|--------|--------|--------|--------|
| | | | | | | | |
| 100 | 147.07 | 151.19 | 153.91 | 169.37 | 195.15 | 220.75 | 264.08 |
| 150 | 141.06 | 145.52 | 148.63 | 163.88 | 188.41 | 213.83 | 257.41 |
| 200 | 137.20 | 141.78 | 145.19 | 160.74 | 185.44 | 210.02 | 253.76 |
| 300 | 132.40 | 137.73 | 141.31 | 157.79 | 181.64 | 206.26 | 248.87 |
| 500 | 127.63 | 133.58 | 138.70 | 155.10 | 179.50 | 204.19 | 246.05 |
| 700 | 123.67 | 132.57 | 137.34 | 154.99 | 179.21 | 204.24 | 245.41 |
| 1000 | 123.77 | 132.07 | 137.44 | 155.43 | 179.74 | 205.03 | 244.44 |
| 1500 | 123.13 | 132.54 | 138.18 | 157.26 | 182.37 | 206.87 | 247.96 |
| 2000 | 123.32 | 133.22 | 139.43 | 159.08 | 184.45 | 208.62 | 249.56 |
| 3000 | 124.66 | 135.65 | 142.47 | 162.57 | 188.35 | 212.26 | 252.67 |
| 5000 | 128.05 | 140.34 | 147.75 | 168.53 | 194.85 | 219.04 | 259.57 |
| 7000 | 131.53 | 144.82 | 153.19 | 175.41 | 202.89 | 227.18 | 267.79 |
| 10,000 | 139.44 | 155.38 | 166.54 | 192.07 | 219.49 | 243.83 | 284.85 |

Each of the terms in (52) is expressed in decibels; P_t is the transmitter power expressed in decibels above one watt; L, is the loss in the transmitting antenna circuit and the transmitting antenna transmission line (this term is set equal to one db for the calculations in this report); R is the median pre-detection signal-to-r.m.s. noise ratio required for the specified grade of service; F is the effective receiver noise figure and includes the effects of the antenna noise as well as the receiver noise together with the receiving antenna circuit and transmission line loss; 1/ it is assumed that the receiver incorporates gain adequate to ensure that the first circuit noise is detectable; $B = 10 \log_{10}(b_0 + b_m)$ is the effective receiver bandwidth factor with bo and bm expressed in cycles per second; bo allows for the drift between the transmitter and receiver oscillators, while bm allows for the band occupied by the modulation; the constant term (-204) is 10 log₁₀ k T where k is Boltzmann's constant and the reference temperature is taken to be 288.44° Kelvin; this is just the noise power in a one cycle per second bandwidth in db relative to one watt.

For the calculations in this report, the transmitter and receiver oscillators were each assumed to have a stability of one part in 10^8 and to vary independently so that $b_0 = \sqrt{2} f_{Mc} \cdot 10^{-2}$. Table 7.5 gives the values of b_m assumed for the various types of service considered. The effective receiver noise figure has been estimated as $F = 5 \log_{10} f_{Mc}$ - 5. Table 7.5 also gives the values of R for the various kinds of service on the assumption that quadruple diversity is used. The value of R for the FM Multichannel system is expected to provide a service with less than an 0.01% teletype character error rate. The FM Multichannel System consists of 36 voice channels, each of which can accommodate sixteen 60 words per minute teletype circuits. The values of R given in Table 7.5 were determined by methods given in a recent report by Watt. 73/ The value of R for the FM Multichannel system corresponds to typical fading encountered at 1000 Mc, and this value of R may change by a few db with frequency as the fading changes, but such changes have so far not been evaluated quantitatively; furthermore, R will also change as the fading changes from hour to hour.

Table 7.6 gives as a function of frequency the maximum permissible hourly median transmission loss for a transmitter power of 10 kw: $L_{\rm M} = 204 + P_{\rm t} - L_{\rm t} - R - F - B$ corresponding to the kinds of service described above. By combining the information in Tables 7.3, 7.4, and 7.6, we can estimate the maximum range for a quadruple diversity system with 10 kw transmitters. These ranges are shown on Fig. 51 as a function of the radio frequency.

Table 7.5

| Type of Service | b _m cycles/sec. | R* decibels | Signal Bandwidth cycles/sec. | Post Detection Signal-to noise ratio decibels |
|---------------------------------------|-------------------------------|----------------|--|---|
| Transmission Loss Measure- ment | 0 | o# | 0 | <u>-</u> |
| FM Multichan- nel System | 3,750,000 | 9.5 | 36 Voice channels each capable of use for sixteen 60 words per min. teletype circuits | 0.01% teletype character error rate |
| FM Music | 150,000 | 26.5 | 15,000 | 50 ** |
| U.S. Standard Television | 3,750,000 | 32.7 | 3,750,000 | 30 ** |

^{*} Ratio between the median intermediate frequency Rayleigh distributed signal and the r.m.s. Rayleigh distributed noise.

^{**} This ratio will be exceeded with a quadruple diversity system for 99% of each hour for which the corresponding value of R is maintained in each receiver.

[#] Diversity reception not involved in this case.

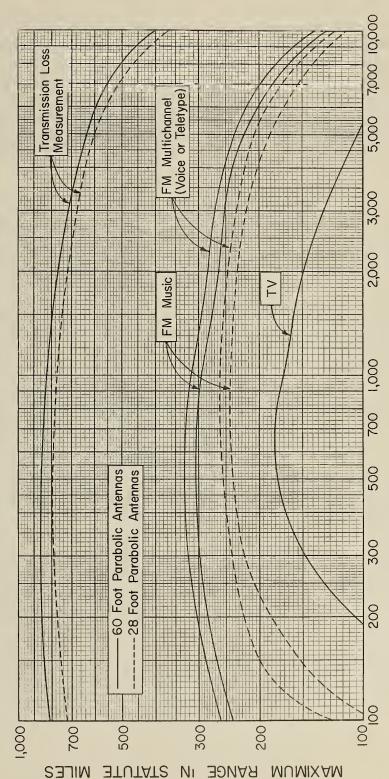
Maximum Permissible Transmission Loss L_{M} for 10 KW Transmitters Using the Parameters of Table 7.5

Table 7.6

| | Transmission loss | FM | FM | U.S. Standard |
|-----------------|-------------------|--------------|--------|------------------|
| f _{Mc} | Measurement | Multichannel | Music | Television |
| 100 | 236.50 | 162.76 | 159.74 | 139.56 |
| 150 | 233.85 | 161.88 | 158.86 | 138.68 |
| 200 | 231.98 | 161.26 | 158.23 | 138.06 |
| 300 | 229.34 | 160.37 | 157.35 | 137.17 |
| 500 | 226.01 | 159.27 | 156.25 | 136.07 |
| 700 | 223.82 | 158.54 | 155.51 | 135.34 |
| 1000 | 221.50 | 157.76 | 154.74 | 134.56 |
| 1500 | 218.85 | 156.88 | 153.86 | 133.68 |
| 2000 | 216.98 | 156.25 | 153.23 | 133.05 |
| 3000 | 214.34 | 155.37 | 152.35 | 132.17 |
| 5000 | 211.01 | 154.27 | 151.25 | 131.07 |
| 7000 | 208.82 | 153.53 | 150.51 | 130.33 |
| 10000 | 206.50 | 152.76 | 149.74 | 129.56 |

MAXIMUM DISTANCE AT WHICH SATISFACTORY SERVICE OF THE TYPES INDICATED MAY BE PROVIDED FOR 99% OF THE HOURS USING KILOWATT TRANSMITTERS AND QUADRUPLE DIVERSITY

Atmospheric Absorption and Rain Attenuation Typical of the Washington, D.C. Area (Smooth Spherical Earth; N_S=301; h_t=h_r=30 feet)



FREQUENCY IN MEGACYCLES PER SECOND

Figure 51

Acknowledgements: The author has had extensive assistance in preparing this report from many members of the Radio Propagation Engineering Division Staff. Acknowledgement of most of this assistance is given in the text or by reference, but in addition the following contributions should be mentioned. Don Watt assisted extensively with the theory of ionospheric roughness outlined in Appendix I. Anita Longley was largely responsible for the sky wave computations for Figs. 31 to 40, inclusive, while Peter Ratcliffe did most of the calculations for Figs. 41 to 45, inclusive. Lew Vogler calculated the diffracted fields shown on Fig. 50, while Anita Longley calculated the scattered fields. Ralph Johler calculated the values given in Table I-2. Gertrude Qvale typed the report, and John Harman drafted most of the figures.

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Appendix I

The Attenuation of Radio Waves Propagated Between a Perfectly Reflecting Spherical Ionospheric Layer and a Spherical Earth

The attenuation with which we are here concerning ourselves is that due to the spreading of the energy over larger and larger areas as it progresses further and further from the transmitting antenna.

For the sake of clarity in presentation, several simpler problems will be solved first in order to illustrate the principles involved. Consider first the attenuation of waves emanating from an isotropic radiator in free space as in Fig. I-1. The total energy passing through the differential elements of area, dA₁ and, dA, normal to the radius vector and at the unit of distance R₁ and at R, respectively, will be equal.

$$dA_1 = R_1 d\psi r_1 d\chi = R_1^2 \cos \psi d\psi d\chi$$
 (I-1)

$$dA = R^2 \cos \psi \, d\psi \, d\chi \tag{I-2}$$

Now, if we let p_1 and p_2 represent the energy density per unit area at the distances R_1 and R_2 , we obtain:

$$p_1 dA_1 = p_2 dA$$
 (I-3)

$$(p_2/p_1) = (dA_1/dA) = R_1^2/R^2$$
 (I-4)

Thus we see that the field intensity (i.e., the energy density) is inversely proportional in free space to the square of the distance

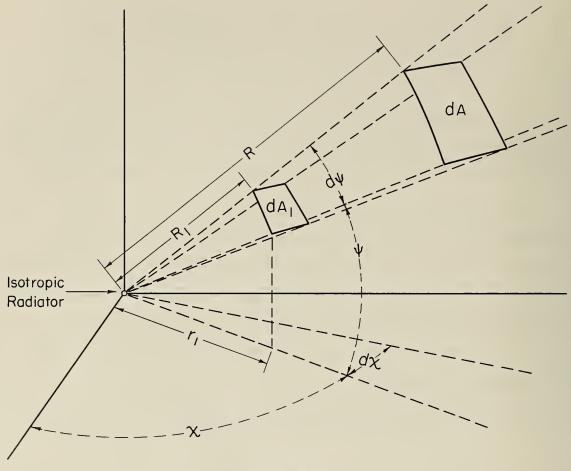


Figure I -1

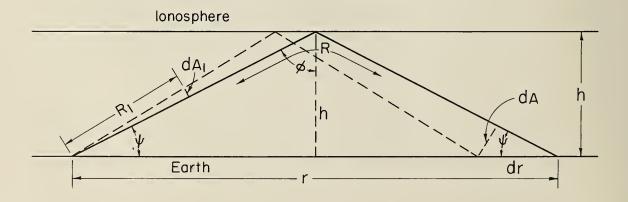


Figure I - 2

from the source. In terms of transmission loss, this result may be expressed:

$$L_{bf}(R) = L_{bf}(R_1) + 20 \log_{10}(R/R_1)$$
 (I-5)

Consider next—see Fig. I-2—the attenuation of waves reflected from a plane perfectly conducting ionosphere at a height, h, above a plane earth and with no atmospheric refraction; in this case we have:

$$dA = -r d\chi \sin \psi dr = R^2 \cos \psi d\psi d\chi \qquad (I-6)$$

In this case again we find that the attenuation of waves reflected from a plane ionosphere is the same as in (I-5).

Finally consider—see Fig. 1-3—the attenuation of waves reflected from a perfectly conducting spherical ionosphere and a perfectly conducting spherical earth with the effects of atmospheric refraction included:

$$dA = -y d\chi \sin \psi dr \qquad (I-7)$$

$$dr = 2 a d \theta ag{I-8}$$

$$y = a \sin 2 \theta \qquad (I-9)$$

$$dA = -2a^{2} \sin 2\theta \sin \psi d\theta d\chi \qquad (I-10)$$

By Snell's law:

$$\frac{\cos (\psi + \theta - \tau)}{\cos \psi} = C = \frac{\{1 + N_o \cdot 10^{-6} \exp(-c_h)\}(a + h_s)}{\{1 + N_o \cdot 10^{-6} \exp(-c_h)\}(a + h)}$$
(I-11)

$$C = \cos (\theta - \tau) - \tan \psi \sin (\theta - \tau) \qquad (I-12)$$

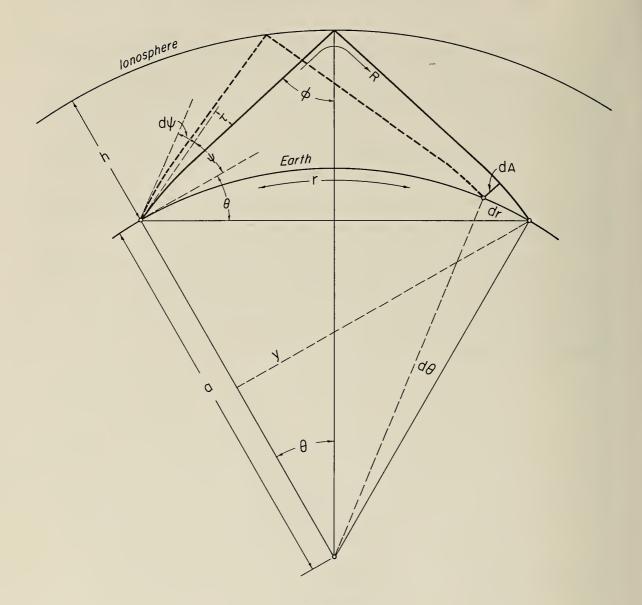


Figure I-3

Since C is a constant, independent of θ , τ and ψ , we find:

$$dC = \frac{\partial C}{\partial \theta} d\theta + \frac{\partial C}{\partial \tau} d\tau + \frac{\partial C}{\partial \psi} d\psi = 0$$
 (I-13)

Thus

$$(d\theta/d\psi) = -\left\{\frac{\partial C}{\partial \psi} / \frac{\partial C}{\partial \theta}\right\} - \left\{\frac{\partial C}{\partial \tau} / \frac{\partial C}{\partial \theta}\right\} \frac{d\tau}{d\psi}$$
 (I-14)

$$(d\theta/d\psi) = -\frac{\sin(\theta-\tau)}{\cos\psi\cos\phi} + \frac{d\tau}{d\psi}$$
 (I-15)

In the particular case when $\psi = 0$:

$$(d\theta/d\psi)_{\psi=0} = -1 + (d\tau/d\psi)_{\psi=0} = -k$$
 (I-16) *

$$L(R) = L_{bf}(R_1) + 10 \log_{10}(\frac{dA}{dA_1}) = L_{bf}(R) - C_1(R)$$
 (I-17)

Thus, if we substitute (I-10) and (I-1) in the above and solve for $C_1(R)$, we obtain:

$$C_1(R) = 10 \log_{10} \left(\frac{R^2 dA_1}{R_1^2 dA} \right) = 10 \log_{10} \left\{ \frac{R^2 \cot \psi}{-2a^2 \sin 2\theta (d\theta/d\psi)} \right\}$$
(I-18)

The function $C_1(R)$ is a convergence factor, expressed in decibels, which measures how much stronger the field intensity at the receiving point is for one reflection at a spherical ionosphere than it would be if it were plane.

The generalization of this expression to m reflections at the ionosphere may be obtained by noting that the only changes required in the above analysis are:

^{*} Note that k here refers to the ratio between the effective and actual radii of the earth.

$$dr = 2 m a d \theta ag{I-19}$$

$$y = a \sin 2 m \theta \qquad (I-20)$$

Thus the convergence factor for m ionospheric reflections may be written:

$$C_{m}(R) = 10 \log_{10} \left\{ \frac{R}{-2 \operatorname{ma} \sin \psi \left(d\theta / d\psi \right)} \right\}_{v} \cdot \left\{ \frac{R \cos \psi}{a \sin 2 \operatorname{m} \theta} \right\}_{h}$$
 (I-21)

The above expression has been divided into two factors with subscripts v and h so that the convergence of the rays in the vertical and horizontal planes, respectively, can be considered separately.

If we introduce the approximate expression $R \cong 2\,m\,a\,\sin\,\theta/\sin\,\varphi \,\,\text{into (I-21) above, we obtain:}$

$$C_{m}(R) \cong 10 \log_{10} \left\{ \frac{\sin \theta}{-\sin \phi \sin \psi (d\theta/d\psi)} \right\}_{v} \cdot \left\{ \frac{2 m \sin \theta \cos \psi}{\sin \phi \sin 2 m \theta} \right\}_{h} \quad (I-21a)$$

The above expression indicates that the convergence in the vertical plane at points far removed from a caustic is independent of the number of ionospheric reflections for a given angle, ψ .

Note that the convergence in the vertical plane becomes infinite when ψ approaches zero; this infinite convergence is demonstrated on Fig. 25. Similarly, the convergence in the horizontal plane becomes infinite at the antipode of the transmitter where $2 \text{ m } \theta = \pi$. Actually, of course, the received energy is finite at these points, and we may use Airy's integral to evaluate the convergence in the vertical plane when $\psi \leq 0$. It can be shown $\frac{1}{2}$ by the solution of the two dimensional wave equation that:

$$C_{V}(R) = 10 \log_{10} \left\{ \frac{-\Omega}{\Omega''(\Psi)} \right\}_{V}$$
 (I-22)

where $\Omega = 2\pi R/\lambda$ is the phase of the waves at the receiving point. If

we multiply the numerator and denominator of the first term in (I-21) by $2\pi/\lambda$ and compare the results with (I-22), we find:

$$\Omega$$
 " $(\psi) = 4 \pi m (a/\lambda) \sin \psi (d\theta/d\psi)$

Now we see that the above second derivative of the phase is equal to zero when the convergence factor becomes infinite; this is the definition of a caustic and the convergence at this caustic may be evaluated by means of the third derivative at this point:

$$\Omega \parallel (\psi) = -4\pi m k(a/\lambda) \qquad (\psi \le 0)$$

Thus, at the caustic in the vertical plane, C_v(R) may be expressed: 1

$$C_{v}(R) = 10 \log_{10} \left\{ \frac{\Omega 2\pi \left\{ A_{i}(0) \right\}^{2}}{\left[\frac{1}{2} \Omega''' \right]^{2/3}} \right\}$$
 (1-23)

In the above $A_i(0)$ is the Airy integral $\frac{2}{}$ with argument zero: $2\pi \{A_i(0)\}^2 = 0.79196357$.

From the above results we obtain the following expression for $C_{m}(R)$ at the caustic in the vertical plane:

$$C_{m}(R) = 10 \log_{10} \left\{ \frac{2\pi (R/\lambda) \ 0.792}{\left[2\pi m \ k(a/\lambda)\right]^{2/3}} \right\}_{v} \left\{ \frac{R}{a \sin 2m(\theta_{m} - k\psi)} \right\}_{h} (\psi \le 0)$$
(I-24)

The convergence at the antipode of the transmitter is of a somewhat different nature. Note, in particular, that there is no point of stationary phase with respect to variations in the azimuth angle, χ , since the waves appear to be arriving from all directions at this particular point. The following treatment of this problem is due to J. R. Wait. $\frac{10}{}$

Using a cylindrical coordinate system centered at the antipode (i.e., ρ , χ , z), we may obtain the following axially symmetric solution of the wave equation, for a time factor exp(i ω t):

$$E_z = A \exp \left[-i k \left(R_a + z \sin \psi \right) \right] J_o(k \rho \cos \psi) \qquad (I-25)$$

$$H_{\chi} = B \exp \left[-ik \left(R_a + z \sin \psi\right) J_1(k \rho \cos \psi) \right] \qquad (I-25a)$$

where $k=2\pi/\lambda$, R_a is the distance along the ray path to the antipode, A and B are constants, and J_0 and J_1 denote Bessel functions. For the ground wave and for those ionospheric modes for which m is sufficiently small so that ψ is negative at the antipode, ψ should be set equal to zero in (I-25) and (I-25a). Note that, in addition to the oscillations with time, the magnitudes of E_z and H_z oscillate with the distance ρ from the antipode, E_z having its maximum value at the antipode while H_z is equal to zero at the antipode. Note also that the variation with ρ is the same, independent of the azimuth angle, χ ; this would be expected since we have assumed that our source radiates uniformly in all directions. When $k \rho \cos \psi > 1$, we may replace the Bessel functions by the first terms in their asymptotic expansions and obtain:

$$E_{z} = A \exp(-ikz \sin \psi) \left\{ \frac{\exp\{i[k(\rho \cos \psi - R_{a}) - \pi/4]\} + \exp\{-i[k(\rho \cos + R_{a}) - \pi/4]\}}{\sqrt{2\pi k \rho \cos \psi}} \right\}$$

 $H_{\phi} = -i \operatorname{B} \exp(-i \operatorname{kz} \sin \psi \left\{ \frac{\exp\left\{i \left[k(\rho \cos \psi - R_{a}) - \pi/4\right]\right\} - \exp\left\{-i \left[k(\rho \cos \psi + R_{a}) - \pi/4\right]\right\}}{\sqrt{2 \pi k \rho \cos \psi}} \right\}$

(I-26a)

The two exponential terms in the above may be identified with waves arriving from opposite directions at a receiving point at a distance ρ from the antipode along great circle paths of lengths R_a - ρ cos ψ and

 $R_a + \rho \cos \psi$, respectively. It is the interference between these two waves which causes the oscillations in the magnitude of the field near the antipode.

To complete our solution we need only evaluate the constants A and B. Rather than doing this directly, we note by (I-21) that the geometrical theory indicates that the focusing in the horizontal plane not too near the antipode is given by:

$$c_{h} = \frac{R \cos \psi}{a \sin 2 m \theta} \qquad (2 m \theta < \pi) \qquad (I-27)$$

and if we multiply (I-27) by the square of the ratio of $|E_z|$ as given by (I-25) and by the first term in (I-26), we obtain the following expression for c_h which must be used instead of (I-27) at points very near the antipode:

$$c_{h} = \frac{R \cos \psi}{a \sin 2 m \theta} \cdot [J_{o}(k \rho \cos \psi)]^{2} 2\pi k \rho \cos \psi \qquad (I-28)$$

When we note that $\rho = a \sin(\pi - 2m \theta)$, the above reduces to:

$$c_h = 2\pi k R \cos^2 \psi \left[J_0(k \rho \cos \psi) \right]^2 \qquad (I-29)$$

For the ground wave $R = \pi a$ and $\psi = 0$; thus, at the antipode $c_h = 2\pi^2 k \, a$ and $C_h = 10 \log_{10} c_h = 34.210 + 10 \log_{10} f_{kc}$, and this clearly represents an extremely large focusing effect for the ground wave at and near this point. The focusing in the horizontal plane for the sky wave modes is only slightly different, but we must add to this the focusing in the vertical plane to obtain the total focusing for these modes. (I-29) is for a vertical electric dipole receiving antenna; if a horizontal magnetic dipole were used for reception, then the J_o should be replaced by J_1 .

^{*} Note that $\cos \psi > 0.995$ for $m \le 16$, and $\cos \psi = 1$ for $m \le 8$ when h = 90 km.

Using a cylindrical coordinate system centered at the antipode (i.e., ρ , χ , z), we may obtain the following axially symmetric solution of the wave equation, for a time factor exp(i ω t):

$$E_z = A \exp \left[-ik \left(R_a + z \sin \psi \right) \right] J_o(k \rho \cos \psi) \qquad (I-25)$$

$$H_{\chi} = B \exp \left[-ik \left(R_a + z \sin \psi\right) J_1(k \rho \cos \psi) \right] \qquad (I-25a)$$

where $k=2\pi/\lambda$, R_a is the distance along the ray path to the antipode, A and B are constants, and J_0 and J_1 denote Bessel functions. For the ground wave and for those ionospheric modes for which m is sufficiently small so that ψ is negative at the antipode, ψ should be set equal to zero in (I-25) and (I-25a). Note that, in addition to the oscillations with time, the magnitudes of E_z and H_z oscillate with the distance ρ from the antipode, E_z having its maximum value at the antipode while H_z is equal to zero at the antipode. Note also that the variation with ρ is the same, independent of the azimuth angle, χ ; this would be expected since we have assumed that our source radiates uniformly in all directions. When $k\rho$ cos ψ > 1, we may replace the Bessel functions by the first terms in their asymptotic expansions and obtain:

$$E_{z} = A \exp(-ikz \sin \psi) \left\{ \frac{\exp\left\{i\left[k(\rho \cos \psi - R_{a}) - \pi/4\right]\right\} + \exp\left\{-i\left[k(\rho \cos + R_{a}) - \pi/4\right]\right\}}{\sqrt{2\pi k \rho \cos \psi}} \right\}$$

(I-26)

$$H_{\phi} = -i B \exp(-ikz \sin \psi \left\{ \frac{\exp\{i[k(\rho \cos \psi - R_a) - \pi/4]\} - \exp\{-i[k(\rho \cos \psi + R_a) - \pi/4]\}}{\sqrt{2\pi k \rho \cos \psi}} \right\}$$

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The two exponential terms in the above may be identified with waves arriving from opposite directions at a receiving point at a distance ρ from the antipode along great circle paths of lengths R_a - ρ cos ψ and

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$$c_{h} = \frac{R \cos \psi}{a \sin 2 m \theta} \cdot [J_{o}(k \rho \cos \psi)]^{2} 2\pi k \rho \cos \psi \qquad (I-28)$$

When we note that $\rho = a \sin(\pi - 2m \theta)$, the above reduces to:

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^{*} Note that $\cos \psi > 0.995$ for $m \le 16$, and $\cos \psi = 1$ for $m \le 8$ when h = 90 km.

All of the above discussion applies to the case when the effective reflecting surface of the ionosphere is smooth and concentric with the surface of the earth. In practice, as the sun rises and sets, or as the geomagnetic latitude of the reflection point is varied, the surface of the ionosphere will undoubtedly change in such a way that its radius of curvature and slope relative to a tangent plane on the earth will vary over appreciable ranges, and this will cause C_m(R) to vary up and down relative to the values expected on the basis of the above analysis. However, except near the antipode, it seems plausible to assume that the median values of C_m(R) may not be much influenced by such changes. The magnitude of the antipodal anomaly will be substantially reduced by these macroscopic perturbations of the spherical concentric shell model assumed for our calculations. Note also that there will be concentric rings around the antipode at which the expected field will be equal to zero. The radii of these concentric rings are the same, regardless of the number, m, of ionospheric reflections, and are determined by the zeros of the Bessel functions; for the electric field, the first two such rings have radii equal to 0.38% and 0.88%. Note, however, that the geographical location of the centers of the antipodal anomalies may be expected to be somewhat different for the different modes for the actual non-concentric ionosphere, and thus these zeros are not likely to be observable unless some means is used to exploit their different times of arrival. The shifts in the geographical locations of the anomalies caused by these macroscopic changes in the ionosphere would be expected to result in a net increase in the fading range in the neighborhood of the antipode.

In addition to these systematic macroscopic changes in the ionosphere, the reflecting surface of the ionosphere will be locally rough, and we will see in the following analysis how this local roughness may be expected to reduce the median values of convergence as computed above for a smooth concentric ionosphere.

We have seen above that the convergence depends, at a given receiving point, upon the smooth variation of the phase of the received waves with changes in elevation angle and azimuth. If we let σ_{Ω} denote the standard deviation of the phase of the waves received via m hops, then we may use Rayleigh's criterion of ionospheric roughness (see Section 2 for a discussion of Rayleigh's criterion as applied to ground roughness) to calculate σ_{Ω} in terms of the variance, σ_{h}^{2} , of the local effective reflection heights, h, of the ionosphere at the points of stationary phase.

$$\sigma_{\Omega} = \frac{720^{\circ} \sigma_{h} \cos \phi \sqrt{m}}{\lambda}$$
 (I-30)

Note that the variance, σ_{Ω}^2 , of phase consists of components arising from (a) a drift of a fixed pattern of ionospheric irregularities relative to a reference great-circle-smooth-concentric-ionosphere path, (b) changes in the shape of these irregularities with time, and (c) changes in the locations of the reflection points in the horizontal plane.

Brennan and Phillips $\frac{3}{}$ find that variations with time of the intensity and phase of a one-hop transmission at 543 kc over a 380 mile path at night indicate rather conclusively that they may be described adequately most of the time by assuming that the received waves consist of a steady component with constant phase and approximately constant amplitude plus a random Rayleigh distributed component of relative intensity k^2 and random relative phase; the amplitude distribution expected in this case is given on Fig. 5 with $K \equiv 10 \log_{10} k^2$. The fixed component may be identified with a

specular reflection expected from the reference smooth surface while the random component arises from the surface roughness. It now becomes clear how ionospheric roughness may be expected to affect the convergence; the specular component will be increased $C_m(R)$ db whereas the random component, since its phase is random, will not be increased at all. If we write $C_m(R) \equiv 10 \log_{10} c_m$, then we find the following expression for the convergence factor $C_m(R, p)$ exceeded 100 p % of the time in terms of the values of $k^2(1-p)$ exceeded 100(1-p)% of the time:

$$C_{m}(R, p) = 10 \log_{10} \left\{ \frac{c_{m} + k^{2}(1-p)}{1+k^{2}(1-p)} \right\}$$
 (I-31)

Note that $k^2(1-p)$ approaches zero as p approaches zero, and thus $C_m(R, p)$ approaches $C_m(R)$, the value expected for a smooth ionosphere, as p approaches zero. On the other hand, $k^2(1-p)$ approaches ∞ for a perfectly rough ionosphere, and in this case there will be no convergence and $C_m(R, p)$ approaches zero.

The values of k^2 to be used in (I-31) may, in principle, be determined from observations of the variations in either the amplitude or the phase of the waves corresponding to a single mode of propagation. However, it is ordinarily better in practice to use the variations in phase as an index to k^2 since the amplitudes of the received waves also vary with ionospheric absorption, and it is sometimes difficult to separate out these absorption variations from the amplitude variations arising from surface roughness alone. In Table I-1 are tabulated some experimental measurements of σ_{Ω} . In some cases the required variance was observed directly, but in other cases it had to be estimated from phase difference measurements made on paths with one common terminal, but with their other terminals

Table I-l Estimates of the Variance of Phase on Ionospheric Paths

| ٩. | km 1 | 0.829 | 1.424 | 0.265 | 1.093 | 0.726 | 1.563 | 0.643 | 0.775 | 0.588 | 0.503 | 0.077 | 0.193 | 255 | 0.062 | 0.216 | 0.085 | 0.068 |
|----------|--------------------------------|------------|------------|-------------|--------------|-------------|-------------|--------------|--------------|-------------|--------------|------------|-------------|-------------|-------------|-------------|------------------|--------------------|
| - | + | | | | | | - | | | | | | | 0. | | | 0. | |
| | os o | 0.146 | 0.170 | 0.174 | 0.206 | 0.174 | 0.206 | 0.144 | 0.169 | 0.144 | 0.169 | 0.146 | 0.165 | 0,189 | 0, 158 | 0.167 | 0.371 | 0.334 |
| MODE | E | 3 | 3 | - | 1 | 1 | П | 3 | 3 | 3 | 3 | 3 | П | 1 | - | - | Н | 1 |
| ь (| 75 | 8.06 | 16.1 | 1.9 | 9.3 | 5.2 | 13.3 | 7.1 | 10.0 | 6.5 | 6.5 | 2.8 | 7.64 | 13,3 | 2.7 | 36.1 | 41.0 | 30.2 |
| | $\sqrt{\frac{2(1-\rho^2)}{2}}$ | | | 1.00 | 1.00 | 1.00 | 1.00 | 1.00 | 1.00 | 1.00 | 1.00 | | 1.41 | 1.00 | 1.00 | 1.41 | 0,83 | 0.63 |
| , N | ~ | | | 10 | 10 | 10 | 10 | 10 | 10 | 10 | 10 | | V 40 | 11 | 11 | 41 | 3,39 | 1.85 |
| , 6 | Degrees | 8.06 | 16.1 | 1.9 | 9.3 | 5.2 | 13.3 | 7.1 | 10.0 | 6.5 | 6.5 | 2.8 | 10.8 | 13.3 | 2.7 | 51 | 34 | 19 |
| MONTH | | FEB. | FEB. | SEPT. | SEPT. | JAN. | JAN. | JUNE | JUNE | DEC. | DEC. | DEC. | APRIL | MARCH | MARCH | MARCH | JAN. TO APRIL | APRIL MAY |
| TIME | | DAY | NIGHT | DAY | NIGHT | DAY | NIGHT | DAY | NIGHT | DAY | NIGHT | DAY | NIGHT | NIGHT | DAY | NIGHT | NIGHT | NIGHT |
| SOURCE | REFERENCE | PIERCE (4) | PIERCE (4) | REDGMENT(5) | REDGMENT (5) | REDGMENT(5) | REDGMENT(5) | REDGMENT (5) | REDGMENT (5) | REDGMENT(5) | REDGMENT (5) | PIERCE (4) | DOHERTY (6) | FLORMAN (7) | FLORMAN (7) | FLORMAN (7) | BRENNAN (3) | REDGMENT (5) NIGHT |
| Distance | Statute Miles | 3230 | 3230 | 640 | 640 | 640 | 640 | 3488 | 3488 | 3488 | 3488 | 3230 | 2350 | 780 | 780 | 1250 | 380 | 454 |
| Fred. | kc | 16 | 16 | 17.2 | 17.2 | 17.2 | 17.2 | 18.4 | 18.4 | 18.4 | 18,4 | 09 | 100 | 115 | 115 | 418 | 543 | 556 |
| No. | | | | | | | | | | | | | | | 14 | 15 | 16 | 17 |

separated by S wavelengths. In this latter case the observed phase difference was related to σ_{Ω}^2 as follows:

$$\sigma_{\rm obs.}^2 = 2(1 - \rho^2) \sigma_{\Omega}^2$$
 (I-32)

In the above ρ is the correlation between the phase variations along the two independent paths. The correlation will vary from $\rho=1$ for S=0 to $\rho\cong 0$ for $S>40\lambda$. Although direct measurements of ρ do not appear to be available in the literature, it seems reasonable to assume that ρ will be of the same order of magnitude as the correlation between the amplitude variations on paths separated a distance S at one end. Measurements of the latter correlation were reported in reference 3, and these data constituted the basis for the estimates in Table I-1.

Using (I-30) estimates of σ_h can also be made, and these are also given in Table I-1 and shown on Fig. I-4. Although the data are quite scattered, the curved lines labelled day and night, respectively, represent the estimates used in this report for calculating the median values of σ_Ω using (I-30). It should be noted that σ_Ω is itself a random variable which changes over wide ranges from hour to hour and from day to day. For example, an analysis of the data in reference 3 shows that the observed phase differences, $\sigma_{\rm obs.}$, ranged from 5.3° to more than 180°, 10% of the values exceeded 118°, 50% exceeded 34° and 90% exceeded 14°. We will see below that the maximum value of $\sigma_{\rm obs.}$ to be expected in practice is 103.9 x $\sqrt{2}$ = 147°, which corresponds to $k^2 = \infty$, and only 5% of their observed values exceeded this value.

The relation between σ_{Ω} and k^2 has been obtained on the assumption that the data fit the Rice $\frac{8}{}$ distribution of a constant vector plus a Rayleigh distributed vector. By integrating over the joint

MEDIAN EFFECTIVE IONOSPHERIC ROUGHNESS PARAMETER, $\sigma_{\rm h}$ OBTAINED FROM OBSERVATIONS OF σ_{Ω} (DEGREES) = 2.4 f_{kc} cos ϕ $\sigma_{\rm h}$ (km)

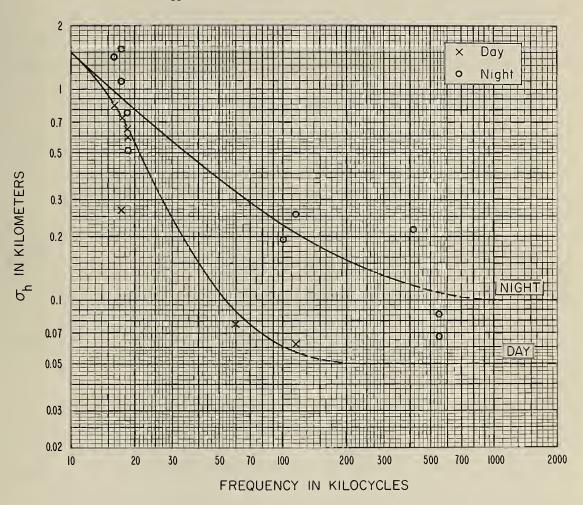


Figure I-4

probability distribution given by Rice for all of the variables except Ω , we obtain the following expression for the probability density function for Ω with k^2 as a parameter:

$$2\pi p(\Omega) = \{1 + \sqrt{\pi} z \exp(z^2)[1 + \exp(z)]\} \exp(-1/k^2)$$
 (I-33)

where $z \equiv \frac{\cos\Omega}{k}$. Fig. I-5 shows $p(\Omega)$ is symmetrically distributed about zero for all values of k^2 . When k^2 is very small, $\sin\Omega$ is distributed approximately normally about zero with variance $\sigma_{\sin\Omega}^2 = k^2/2$. When k^2 approaches infinity, Ω is uniformly distributed between - 180° and +180° and σ_{Ω} approaches 103.923 degrees. Fig. I-6 gives the cumulative distribution defined by:

$$P[\Omega(t) > \Omega(P)] = 1 - \int_{-\pi}^{\Omega(P)} p(\Omega) d\Omega \qquad (\ddagger -34)$$

The mean absolute value $|\overline{\Omega}|$ and variance σ_{Ω}^2 are also of interest:

$$|\overline{\Omega}| = 2 \int_{0}^{\pi} \Omega p(\Omega) d\Omega$$
 (I-35)

$$\sigma_{\Omega}^{2} = 2 \int_{0}^{\pi} \Omega^{2} p(\Omega) d\Omega \qquad (I-36)$$

For many applications the cumulative distribution of the absolute value of $|\Omega(t)|$ is of greater interest:—

$$P' \lceil |\Omega(t)| > \Omega(P') \rceil = 2 P \lceil \Omega(t) > \Omega(P) \rceil$$
 (I-37)

The distribution of P' is given in a recent paper $\frac{9}{}$ and several of its percentage points, together with $|\overline{\Omega}|$ and σ_{Ω} are shown on Fig. I-7 as a function of k^2 . By using the median values of σ_{Ω} determined from Fig. I-4, we may use the results shown on Fig. I-7 to determine the median values of k^2 (0.5) required for the evaluation of $C_{\mathbf{m}}(R, 0.5)$.

THE PROBABILITY DENSITY FUNCTION $p(\Omega)$

$$p(\Omega) = \frac{1}{2\pi} \left\{ 1 + \sqrt{\pi} \frac{\cos \Omega}{k} \exp\left(\frac{\cos^2 \Omega}{k^2}\right) \left[1 + \operatorname{erf}\left(\frac{\cos \Omega}{k}\right) \right] \right\} \exp\left(-\frac{1}{k^2}\right)$$

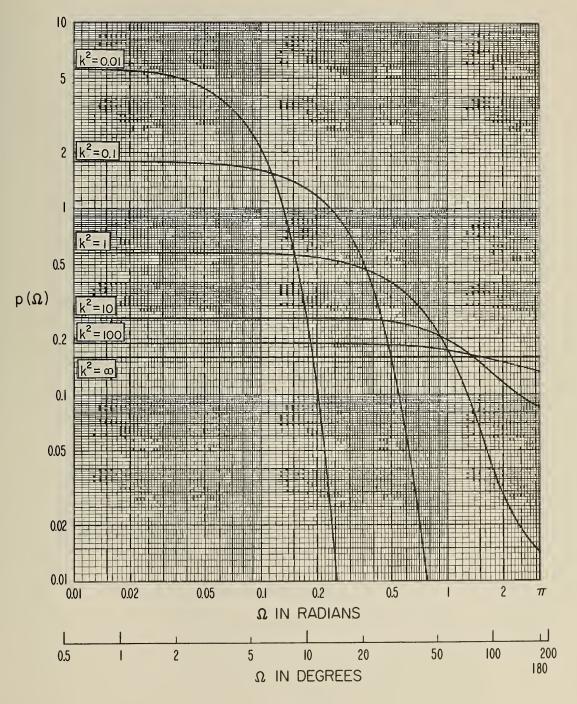


Figure I-5

THE CUMULATIVE DISTRIBUTION OF Ω (†)

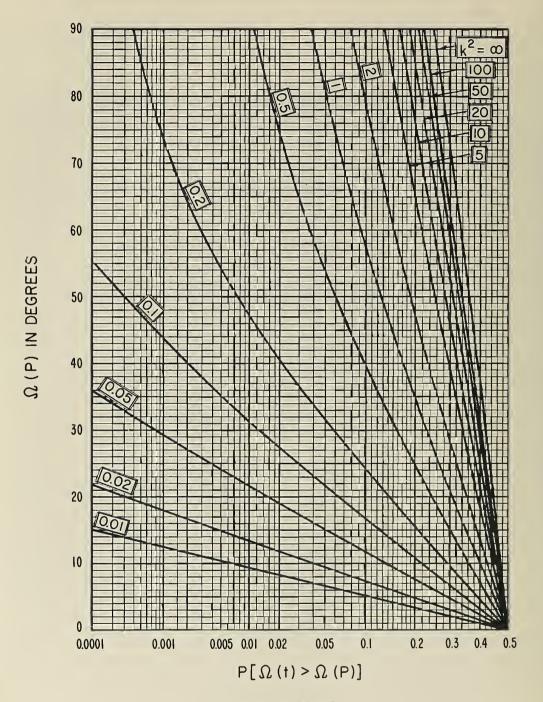


Figure I-6

STATISTICS OF THE DISTRIBUTION OF THE PHASE OF A CONSTANT VECTOR PLUS A RAYLEIGH DISTRIBUTED VECTOR

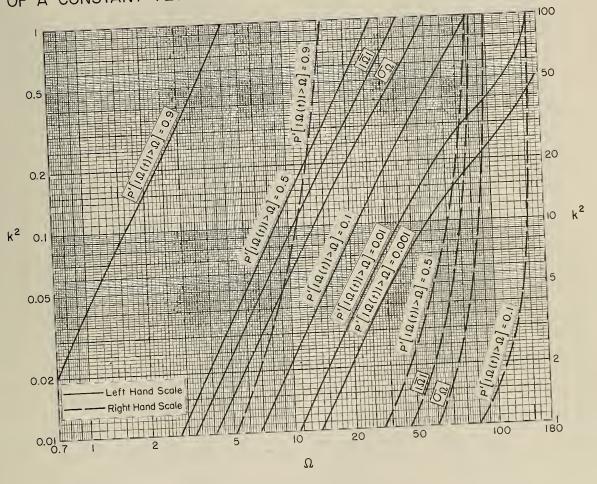


Figure I-7

Table I-2 below gives the values of $|\overline{\Omega}|$ and of σ_{Ω} for several values of $k^2.$

Table I-2

| k ² | $ \overline{\Omega} $ radians | σ_{Ω} radians | $ \overline{\Omega} $ degrees | σ_{Ω} degrees |
|----------------|-------------------------------|---------------------------|-------------------------------|---------------------------|
| 0.01 | 0.056514 | 0.070890 | 3.2380 | 4.0617 |
| 0.02 | 0.080059 | 0.10051 | 4.5871 | 5.7590 |
| 0.05 | 0.12726 | 0.16023 | 7.2915 | 9.1803 |
| 0.1 | 0.18172 | 0.23013 | 10.412 | 13.185 |
| 0.2 | 0.26330 | 0.34032 | 15.086 | 19.499 |
| 0.5 | 0.44605 | 0.60664 | 25.557 | 34.758 |
| 1 | 0.64346 | 0.87134 | 36.868 | 49.924 |
| 2 | 0.85196 | 1.1175 | 48.813 | 64.031 |
| 5 | 1.0876 | 1.3661 | 62.313 | 78.270 |
| 10 | 1.2217 | 1.4972 | 69.999 | 85.785 |
| 20 | 1.3212 | 1.5907 | 75.701 | 91.142 |
| 50 | 1.4119 | 1.6735 | 80.896 | 95.884 |
| 100 | 1.4582 | 1.7149 | 83.547 | 98.259 |
| 200 | 1.4911 | 1.7441 | 85.431 | 99.929 |
| 500 | 1.5203 | 1.7698 | 87.107 | 101.40 |
| 1000 | 1.5351 | 1.7827 | 87.953 | 102.14 |
| σ | π/2 | $\pi/\sqrt{3}$ | 90.000 | 103.92 |

An analysis was made in reference 3 of both the amplitude variations on single paths as well as the phase differences between paths separated by 3.39 wavelengths at one end. Using their observed median standard deviation of $\sigma_{\rm obs.} = 34^{\circ}$, the median value $\sigma_{\Omega} = 41^{\circ}$ given in Table I-1 was estimated; this value corresponds by Fig. I-7 to $k^2(0.5) = 0.68$. The analysis of their amplitude variations gives directly the estimate $k^2(0.5) = 2/a^2 = 0.854$, and this latter estimate is somewhat larger, as might have been expected, since the amplitude variations are biased by changes in absorption.

References to Appendix I

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- 2. British Association Mathematical Tables, Part Volume B; British University Press, Cambridge, 1946.
- 3. D. G. Brennan and M. Lindeman Phillips, "Phase and amplitude variability in medium-frequency ionospheric transmission," Technical Report No. 93, Massachusetts Institute of Technology Lincoln Laboratory, September, 1957; note that the a in this report, which is used as a parameter describing the Rice distribution of a constant plus a Rayleigh distributed vector, is equal to $\sqrt{2}$ /k where k^2 denotes the relative intensity of the random component as used in this appendix.
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- 7. Private communication from E. F. Florman of the Boulder Laboratories of the National Bureau of Standards.
- 8. S. O. Rice, "Properties of a sine wave plus random noise,"
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 the joint probability distribution is given by equation (4.6).
- 9. K. A. Norton, E. L. Shultz, and H. Yarbrough, "The probability distribution of the phase of the resultant vector sum of a constant vector plus a Rayleigh-distributed vector," Jour. Appl. Phys., vol. 23, pp. 137-141; January, 1952. Note that the k in this reference is a power ratio rather than a voltage ratio and that the formulas and graphs give the distribution of |φ|.
- 10. Private communication from J. R. Wait of the National Bureau of Standards, Boulder Laboratories.

Appendix II

The Physics of Ionospheric Scatter Propagation

Wheelon $\frac{54}{5}$ $\frac{56}{5}$ gives the following formula for the scattered power, p_s, relative to the power, p_f, expected for propagation over the same distance in free space:

$$(p_s/p_f) = 4\pi \text{ b sec } \phi \text{ } \sigma(k^2)$$
 (II-1)

$$\sigma(k^2) = \text{constant } r_e^2 < [d N_e/dh]^2 \ell_o^5 > f(k^2)$$
 (II-2)

$$f(k^{2}) = [1 + k^{2} l_{o}^{2}]^{-5/2} [1 + (k^{2} l_{s}^{2})^{2/3}]^{-2} [1 + (k^{2} l_{s}^{2})^{2}]^{-4/3}$$
 (II-3)

The dimensionless constant in (II-2) is of the order of unity, and will be set equal to one in the subsequent analysis; k^2 is defined by (32). [d N_e/dh] is the gradient of the electron density expressed in electrons/cubic meter/meter; b is the effective thickness of the scattering layer expressed in meters; the classical electron radius, $r_e = 2.81785 \times 10^{-15}$ meters and ℓ_o is the scale of turbulence expressed in meters.

Note that when $f = f_{MUF}$ we have $k^2 = 0$ and f(0) = 1. The constant S(0.5) = -8.4 db determined from the radio data may be readily identified with:

$$S(0.5) = 10 \log_{10} 4\pi \ b \ \sigma(0) = -8.4$$
 (II-4)

Consequently it follows that:

$$4\pi \, b \, r_e^2 < \left[d \, N_e / dh \right]^2 \, l_o^5 > = 0.1445$$
 (II-5)

If we let b = 10,000 meters, we obtain:

$$< [d N_e/dh]^2 I_o^5 > = 14.48 \times 10^{22}$$
 (II-6)

Note that the average indicated by <> is taken over the scattering volume. In the troposphere it has been found that l_0 is a random variable with respect to time at a fixed point and with respect to location at a fixed time; more specifically, $L = 10 \log_{10} l_0$ has been found to be normally distributed about its median value l_{om} with $\sigma_L = 5$ db. It seems not unreasonable to assume a similar variability for l_0 in the ionosphere. Similarly we may assume that $\left[d N_e / dh \right]^2$ is log-normally distributed about its median value $\left[d N_e / dh \right]^2_m$ with a similar standard deviation, i.e., about 5 db. On these assumptions it can be shown $\frac{55}{}$ by simple statistical analysis that:

$$< [d N_e/dh]^2 \ell_0^5 > = [d N_e/dh]_m^2 \ell_{om}^5 \exp[0.02651 \sigma^2]$$
 (II-7)

In the above, σ denotes the standard deviation, expressed in decibels, of $10 \log_{10} \{[d N_e/dh]^2 l_o^5\}$; if we neglect any correlation between the variations of $[d N_e/dh]$ and of l_o , then $\sigma^2 \cong (5)^2 + (25)^2 = 650$ and $\exp [0.02651 \sigma^2] = 3.045 \times 10^7$. If we combine (II-7) and (II-6) and set $l_o = 100$ meters, we obtain $[d N_e/dh]_m = 690$ electrons/c.c/kilometer. If this value is compared with the value 3,800 electrons/c.c./kilometer expected with $\sigma^2 = 0$, we see the importance of allowing for this statistical correction; the actual value probably lies somewhere between these two estimates, and can be estimated more precisely only when more adequate information becomes available relative to the variances of these variables.

The above analysis refers to the scatter expected for frequencies just above the E layer MUF. The forward scatter on the higher

frequencies where S was actually evaluated becomes independent of ℓ_o since $[1+k^2\ell_o^2]^{-5/2}=k^{-5}\ell_o^{-5}$ when $k^2\ell_o^2>>1$. In this case the correction factor should be determined for $\sigma^2=25$, i.e., $\exp[0.0265 \ \sigma^2]=1.940$ and the expected median gradient on these assumptions is then 2700 electrons/c.c./km.

It appears from the above analysis that S may increase with decreasing frequency because of the increasing importance of the variance of ℓ_0 at these lower frequencies. This statistical factor should not be ignored in analyses of ionospheric scatter data. However, it was suppressed in the present analysis because of the lack of definitive data on σ^2 .

Note that scale lengths of the order of $l_{\rm om}$ = 100 meters and electron density gradients of the order of 1,000 electrons/c.c./km. are not unreasonable values to assume for the lower ionosphere, and we conclude that Wheelon's theory provides a useful description of ionospheric turbulence which is not inconsistent with our knowledge of the ionosphere.



Appendix III

An Additional Height-Gain Factor in Transmission Loss

In free space the field strength e, expressed in volts per meter at a distance d, expressed in meters, from an isotropic transmitting antenna radiating p_r watts may be determined from the relation:

$$\frac{e^2}{z} = \frac{p_r}{4\pi d^2} \tag{III-1}$$

(Radiation from an isotropic antenna in free space)

where $z = 4\pi \text{ c} \cdot 10^{-7}$ = impedance of free space expressed in ohms, and $c = 2.997925 \cdot 10^8$ meters per second = velocity of light in free space.

Now consider the intensity of the radiation field of a short vertical electric dipole antenna of length l and at a height h_a above a perfectly conducting plane. By re-distributing the field in the space above the plane, the radiation resistance is modified by the presence of the surface as follows:

$$R_{e} = \frac{2\pi z \ell^{2}}{3\lambda^{2}} [1 + \Delta_{a}]$$
 (III-2)*

$$\Delta_{\rm a} = \frac{3}{(2k \, h_{\rm a})^2} \left[\frac{\sin{(2k \, h_{\rm a})}}{2k \, h_{\rm a}} - \cos{(2k \, h_{\rm a})} \right]$$
 (III-3)*

^{*} These relations are derived by S. A. Schelkunoff in Chapters VI and IX of the book "Electromagnetic Waves," D. Van Nostrand Company, 1943.

In the above $k = 2\pi/\lambda = 2\pi f/c$, i.e., λ is the wavelength in free space. Note that Δ_a approaches zero at large heights above the surface, and R_e approaches its free space value. On the other hand $\Delta_a = 1$ for $h_a = 0$, and the radiation resistance is then just twice its free space value. Using (III-2) we find that the field intensity of the short dipole over the perfectly conducting plane surface may be expressed:

$$\frac{e^2}{z} = \frac{p_r(3/2) [2 \cos \psi \cos (k h_a \sin \psi)]^2}{4\pi d^2 [1 + \Delta_a]}$$
 (III-4)

(Radiation from a short vertical electric dipole over a perfectly conducting surface)

Note that the factor (3/2) is just the free space gain of the short dipole antenna. Since $\Delta_a = 1$ for $h_a = 0$, the field intensity is 3 db greater when $\psi = 0$ for a dipole on the surface of a perfectly conducting plane than for a short dipole in free space. In more familiar units (III-4) with $h_a = \psi = 0$ may be expressed:

$$e(\mu v/meter) = 299,896.2 \sqrt{p(kw)/d_{km}}$$
 (h_a = ψ = 0) (III-5)

Furthermore, the effective absorbing area of a short vertical electric dipole antenna at a height h above a perfectly conducting plane may be expressed:

$$a_e = \frac{\lambda^2 (3/2) \cos^2 \psi}{4\pi [1 + \Delta_b]}$$
 (III-6)

where Δ_b is defined by (III-3) with h_a replaced by h_b .

Combining (III-4) and (III-6) we may express the transmission loss in decibels between short vertical electric dipoles at heights h_a and h_b above a perfectly conducting plane as follows:

$$L = L_{bf} - G_f - G_r + A \qquad (III-7)$$

$$A = -20 \log_{10} [2 \cos^2 \psi \cos (k h_a \sin \psi)] + L_a + L_b$$
 (III-8)

$$L_{a, b} = 10 \log_{10} [1 + \Delta_{a, b}]$$
 (III-9)

Note that $L_{\rm bf}$ is the basic transmission loss expected in free space, $G_{\rm t}=G_{\rm r}=1.761$ db, and that the transmission loss, A, relative to free space contains two height gain factors which are not ordinarily considered in field strength calculations. No allowance was made in the calculations in this report for the additional losses $L_{\rm a}$ and $L_{\rm b}$ which arise from the redistribution of the field intensity in space which, in turn, is associated with the proximity of the antennas to the ground. Thus the transmission losses shown on Figs. 7, 8, 16, 17, 18, 19, 20, 21, 27, 31, 32, 33, 34, 35, 36, 37, 38, 39, 40, etc. are too small by an amount ranging from about 6 db at very low frequencies and low antenna heights to zero at the higher frequencies. Fig. III-1 shows this additional loss as a function of antenna height (h/ λ), expressed in wavelengths, for the case of a perfectly reflecting surface, and this should also represent a good approximation in those cases where the antennas are erected over large ground screens.

It is of interest, although not surprising, to note that the transmission loss between vertical electric dipoles on the surface of a <u>perfectly conducting</u> plane ($h_a = h_b = \psi = 0$) is the same as if the dipoles were in free space, even though the field intensity at the surface is 3 db greater.

It should be noted that Schelkunoff identified the factors in (III-2) somewhat differently; thus he considered the ground to be an integral part of the antennas, and set $G_t = 10 \log_{10} \{(3/2) \cdot 2/[1 + \Delta_a]\}$,

 $G_r = 10 \log_{10} \{(3/2) \cdot 2/[1 + \Delta_b]\}$ and $A = -20 \log_{10} [\cos^2 \psi \cos(k h_a \sin \psi)]$. It seems to this writer that the terms $10 \log_{10} [1 + \Delta_a, b]$ should be excluded from G_t and G_r . Thus, according to this approach, G_t and G_r are the free space gains of the antennas, independently of their location, and the path antenna gain G_p , when measured by replacing the actual antennas by isotropic antennas, will still be approximately equal to $G_t + G_r$.

Suppose now that we use small loop antennas of area S, with their axes normal to the plane of propagation, parallel to the perfectly conducting surface and at heights h and h respectively. In this case:

$$R_{\rm m} = \frac{8\pi^3 z S^2}{3\lambda^4} [1 + \Delta_{\rm b}^{\rm i}]$$
 (III-10)

$$\Delta_{b}^{!} = (3/2) \left[\left(1 - \frac{1}{(2k h_{b})^{2}} \right) \frac{\sin (2k h_{b})}{2k h_{b}} + \frac{\cos (2k h_{b})}{(2k h_{b})^{2}} \right]$$

(III-11)

$$a_e = \frac{\lambda^2 (3/2)}{4\pi [1 + \Delta_b']}$$
 (III-12)

$$A = -20 \log_{10} [2 \cos (k h_a \sin \psi)] + L_a^{\dagger} + L_b^{\dagger}$$
 (III-13)

Note that Δ_b' approaches zero at large heights and $\Delta_b' = 1$ for $h_b = 0$.

Consider next the transmission loss between two small loop antennas at heights h and h, respectively, above a perfectly conducting surface with their axes normal to this surface. In this case:

$$R_{\rm m} = \frac{8\pi^3 z s^2}{3\lambda^4} [1 - \Delta]$$
 (III-14)

$$A = -20 \log_{10} [2 \cos^2 \psi \sin (k h_a \sin \psi)] + L_a'' + L_b''$$
 (III-15)

$$L_{a, b}^{"} = 10 \log_{10} [1 - \Delta_{a, b}]$$
 (III-16)

The factor L" is also shown as a function of (h/λ) on Fig. III-1.

Finally consider the transmission loss between two short horizontal electric dipoles of length ℓ , normal to the plane of propagation and at heights h_a and h_b , respectively, above a perfectly conducting plane surface. In this case:

$$R_{e} = \frac{2\pi z t^{2}}{3\lambda^{2}} [1 - \Delta']$$
 (III-17)

A = -20
$$\log_{10} [2 \sin (k h_a \sin \psi)] + L_a^m + L_b^m$$
 (III-18)

$$L_{a, b}^{"} = 10 \log_{10} [1 - \Delta'_{a, b}]$$
 (III-19)

Note that L^m and Lⁿ both approach (- ∞) as h approaches zero, but the radiation resistance R simultaneously approaches zero, and it would be difficult in practice to keep the radiated power constant as the antennas are brought nearer and nearer to the surface. When h and h are both much less than a wavelength, A, as defined by (III-15) for horizontal loops becomes independent of these heights and equal to $A = 20 \log_{10}(kd/5)$; similarly A, as defined by (III-18) for horizontal electric dipoles approaches $20 \log_{10}(2 \text{ kd/5})$ for h and h much less than a wavelength.

Since the factors L_a and L_b were omitted in calculating the transmission losses shown on Figs. 9 and 10, these values are much too large at the lower frequencies since the 30-foot antennas are in this case only a small fraction of a wavelength above the surface.

All of the above results refer to the case of a perfectly conducting plane surface and to distances $d >> \lambda$. For a finitely conducting ground, the factor 2 cos (k h sin ψ) in (III-8) and (III-13) or the factor 2 sin (k h sin ψ) in (III-15) and (III-18) must be replaced by the appropriate attenuation factor, 2 W, relative to the free space field. For example, for electric dipoles over a flat earth of finite conductivity and with $h_a = h_b = 0$:

$$W = |1 + i \sqrt{\pi p} \exp(-p) \operatorname{erfc}(-i \sqrt{p})$$
 (III-20)

Here p denotes Sommerfeld's numerical distance as defined in reference 13 where a comprehensive discussion is given of the radiation fields of electric and magnetic dipoles over a finitely conducting plane earth. Furthermore, Δ and Δ' will be modified when the antennas are located over a finite ground, $\frac{1}{2} \frac{2}{3} \frac{4}{4}$ but this difference will often largely be cancelled in practice if a large ground screen is used under the antennas. Although (III-14) indicates that $R_{\rm m}$ approaches zero as the vertical magnetic dipoles approach the perfectly conducting surface, Wait $\frac{1}{4}$ has shown that $R_{\rm m}$ becomes very large when such loops are brought near a finitely conducting ground.

It is sometimes convenient to be able to relate the basic transmission loss, $\ell_{\rm h}$, to the field strength e:

$$p_a/\ell_b = p_a = (e^2/z) \cdot (\lambda^2/4\pi)$$
 (III-21)

(Isotropic antennas in free space)

Expressed in decibels, we obtain from (III-21):

$$L_b = 77.216 + 20 \log_{10} f_{kc} + P_r - E_b$$
 (III-22)

(Isotropic antennas in free space)

In the above, P_r is the radiated power expressed in db above one kilowatt, and E_b is the field strength in db above one microvolt per meter. If antennas with free space gains G_t and G_r are used, we find that $E_t = E_b + G_t$, and the transmission loss between these antennas in free space may be expressed:

$$L = L_b - G_t - G_r = 77.216 + 20 \log_{10} f_{kc} + P_r - E_t - G_r$$
 (III-23)

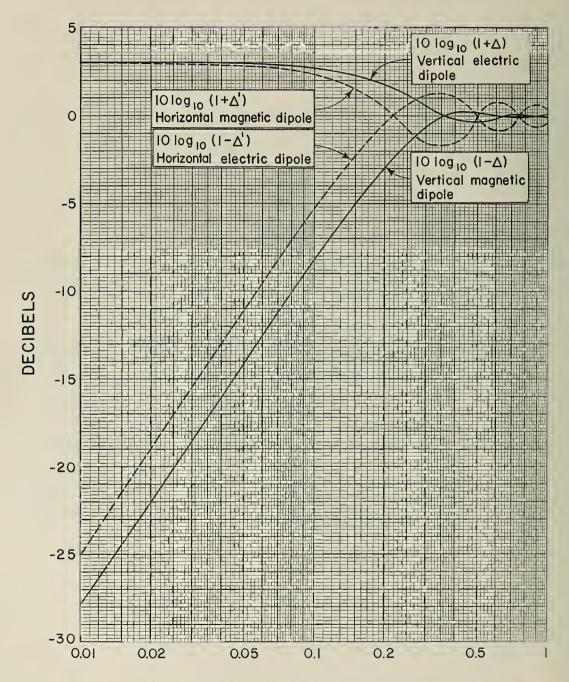
(Antennas with gains G_{t} and G_{r} in free space)

For a half wave dipole transmitting antenna $G_t = 2.15$ db, and we obtain from the above the relation given at the bottom of page two of the report.

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TRANSMISSION LOSS ARISING FROM A CHANGE IN THE RADIATION RESISTANCE OF SHORT DIPOLE ANTENNAS NEAR A PERFECTLY CONDUCTING SURFACE



(h/λ) HEIGHT OF ANTENNA IN WAVELENGTHS ABOVE THE SURFACE





