

Distribution and Stratigraphic Significance of Foraminifera and Algae in Well Cores from Madison Group (Mississippian), Williston Basin, Montana

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By WILLIAM J. SANDO and BERNARD L. MAMET

C O N T R I B U T I O N S T O S T R A T I G R A P H Y

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*Determination of age and correlation of the
Madison Group in the subsurface of the
Williston Basin on the basis of foraminifers and algae*



UNITED STATES DEPARTMENT OF THE INTERIOR

JAMES G. WATT, *Secretary*

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CONVERSION FACTORS

| Metric unit | Inch-Pound equivalent | Metric unit | Inch-Pound equivalent |
|--|---|--|--|
| Length | | | |
| millimeter (mm) | \equiv 0.03937 inch (in) | liter per second (L/s) | \equiv .0353 cubic foot per second |
| meter (m) | \equiv 3.28 feet (ft) | cubic meter per second per square mile [(ft ³ /s)/mi ²] | \equiv 91.47 cubic feet per second per square mile [(ft ³ /s)/mi ²] |
| kilometer (km) | \equiv .62 mile (mi) | meter per day (m/d) | |
| Area | | | |
| square meter (m ²) | \equiv 10.76 square feet (ft ²) | meter per kilometer (m/km) | \equiv 3.28 feet per day (hydraulic conductivity) (ft/d) |
| square kilometer (km ²) | \equiv .386 square mile (mi ²) | kilometer per hour (km/h) | \equiv 5.28 feet per mile (ft/mi) |
| hectare (ha) | \equiv 2.47 acres | meter per second (m/s) | \equiv .9113 foot per second (ft/s) |
| Volume | | | |
| cubic centimeter (cm ³) | \equiv 0.061 cubic inch (in ³) | meter squared per day (m ² /d) | \equiv 3.28 feet per second |
| liter (L) | \equiv 61.03 cubic inches | cubic meter per day (m ³ /d) | \equiv 10.764 feet squared per day (ft ² /d) |
| cubic meter (m ³) | \equiv 35.31 cubic feet (ft ³) | cubic meter per second (m ³ /s) | (transmissivity) |
| cubic meter | \equiv .00081 acre-foot (acre-ft) | cubic meter per minute (m ³ /min) | |
| cubic hectometer (hm ³) | \equiv 810.7 acre-feet | liter per second (L/s) | |
| liter | \equiv 2.113 pints (pt) | liter per second per meter (L/s/m) | |
| liter | \equiv 1.06 quarts (qt) | kilometer per hour (km/h) | |
| cubic meter | \equiv .26 gallon (gal) | meter per second (m/s) | \equiv 15.85 gallons per minute |
| cubic meter | \equiv .00026 million gallons (Mgal) or 10 ⁶ gal | gram per cubic centimeter (g/cm ³) | \equiv 4.83 gallons per minute per foot (gal/min)/ft |
| cubic meter | \equiv 6,290 barrels (bbl) (1 bbl = 42 gal) | gram per square centimeter (g/cm ²) | \equiv .62 mile per hour (mi/h) |
| Weight | | | |
| gram (g) | \equiv 0.035 ounce, avoirdupois (oz avdp) | gram per cubic centimeter (g/cm ³) | \equiv 2.237 miles per hour |
| gram | \equiv .00122 pound, avoirdupois (lb avdp) | gram per square centimeter (g/cm ²) | \equiv 62.43 pounds per cubic foot (lb/ft ³) |
| metric tons (t) | \equiv 1.102 tons, short (2,000 lb) | gram per square centimeter (g/cm ²) | \equiv 2.048 pounds per square foot (lb/ft ²) |
| metric tons | \equiv 0.9832 ton, long (2,240 lb) | centimeter | \equiv .0142 pound per square inch (lb/in ²) |
| Specific combinations | | | |
| kilogram per square centimeter (kg/cm ²) | \equiv 0.96 atmosphere (atm) | degree Celsius (°C) | \equiv 1.8 degrees Fahrenheit (°F) |
| kilogram per square centimeter | \equiv .98 bar (0.9869 atm) | degrees Celsius (temperature) | \equiv [(1.8 × °C) + 32] degrees Fahrenheit |
| metric meter per second (m ³ /s) | \equiv 35.3 cubic feet per second (ft ³ /s) | | |
| Temperature | | | |

CONTRIBUTIONS TO STRATIGRAPHY

DISTRIBUTION AND STRATIGRAPHIC SIGNIFICANCE OF FORAMINIFERA AND ALGAE IN WELL CORES FROM MADISON GROUP (MISSISSIPPIAN), WILLISTON BASIN, MONTANA

By WILLIAM J. SANDO and BERNARD L. MAMET¹

ABSTRACT

Core samples of the Madison Group (Mississippian) from four wells in the Williston Basin, Montana, have yielded microfauna and microflora that are less abundant and diversified than those from outcrops of the Madison to the west because of less favorable environmental conditions in the Williston Basin area during Madison time. Mamet Zones 7, 8, 9, 10, and 11 (or possibly 12) are represented. Microfossil distribution supports the interpretation that the upper part of the Madison in the Williston Basin is no younger than early Meramecian (early Viséan) and that the Madison is bounded above by the same regional disconformity that has been widely observed in surface sections.

INTRODUCTION

In a recent review of paleontologic evidence bearing on the age and geologic history of the Madison Group in the Williston Basin, Sando (1978) presented arguments, based on the distribution of corals, that the Madison Group in the subsurface is the same age as it is in outcrops west of the Williston Basin and that it is bounded above by the same regional disconformity observed in surface sections. These arguments run counter to most previously published interpretations of the age and nature of the upper boundary of the Madison in the Williston Basin (Perry and Sloss, 1943, p. 1301–1302; Hadley, 1950, p. 44; Sloss, 1952, p. 65, 67; Nordquist, 1953, figs. 4 and 5; Gardner, 1959, p. 331–332; Sandberg, 1962, p. 64; Roberts, 1966, fig. 2; Maughan and Roberts, 1967, p. B5, fig. 1, pl. 2).

In order to test these conclusions, thin sections were made from core samples from the seven wells reported on by Sando (1978); Mamet searched the thin sections for foraminifers and algae and studied rewarding samples. The purpose of this report is to present data on the distribution and significance of these samples. No previous study of these microfossils from the Williston Basin has been published.

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We are grateful to J. T. Dutro, Jr., and J. M. Berdan, U.S. Geological Survey, for their very helpful suggestions and their critical review of the manuscript. The excellent quality of the thin sections results from the work of K. R. Moore, U.S. Geological Survey.

DISTRIBUTION OF MICROFOSSILS

Although samples from all seven wells were examined, only the Pine, Richey, Eggebrecht, and East Poplar wells along the west flank of the Williston Basin in Montana provided adequate fauna and flora for zonal determinations (wells 1-4, fig. 1). Foraminifers and algae are rare in all cores; of 176 samples studied, only 28 samples contained zonally identifiable microfossil assemblages. Identifiable samples were zoned according to the North American scheme of Mamet and Skipp (1970a, b).

Figure 2 shows the distribution of zonally identifiable microfossils. Data on corals in these wells can be found in Sando (1978), and detailed lithic descriptions of the units in the wells can be obtained from American Stratigraphic Company logs. Coral zones recognized in the wells are those of Sando and others (1969). The correlation datum is the Coral Zone C₂/D boundary.

BIOSTRATIGRAPHY

LODGEPOLE LIMESTONE

The Lodgepole Limestone, lowermost formation in the Madison Group, contains a few samples in the upper part that represent Zone 7, which ranges from uppermost Kinderhookian into lower Osagean (middle Tournaisian of Europe). The Zone 7 assemblage is widespread in the Woodhurst Member of the Lodgepole, where it coincides with Coral Zone C₁. In the Williston Basin cores, the assemblage is very sparse and poorly diversified and consists mostly of Earlandiidae and accessory Endothyridae and Tournayellidae:

Bisphaera sp.

Brunsiina sp.

Calcisphaera laevis (Williamson)

Earlandia clavatula (Howchin)

E. elegans (Rauzer-Chernoussova)

E. minima (Birina)

Latiendothyra sp.

L. of the group *L. parakosvensis* (Lipina)

Palaeospirolectammina tchernyshinensis (Lipina)

Rectoseptaglomospiranella sp.

Septabrunsiina sp.

Septaglomospiranella primaeva (Chernysheva)

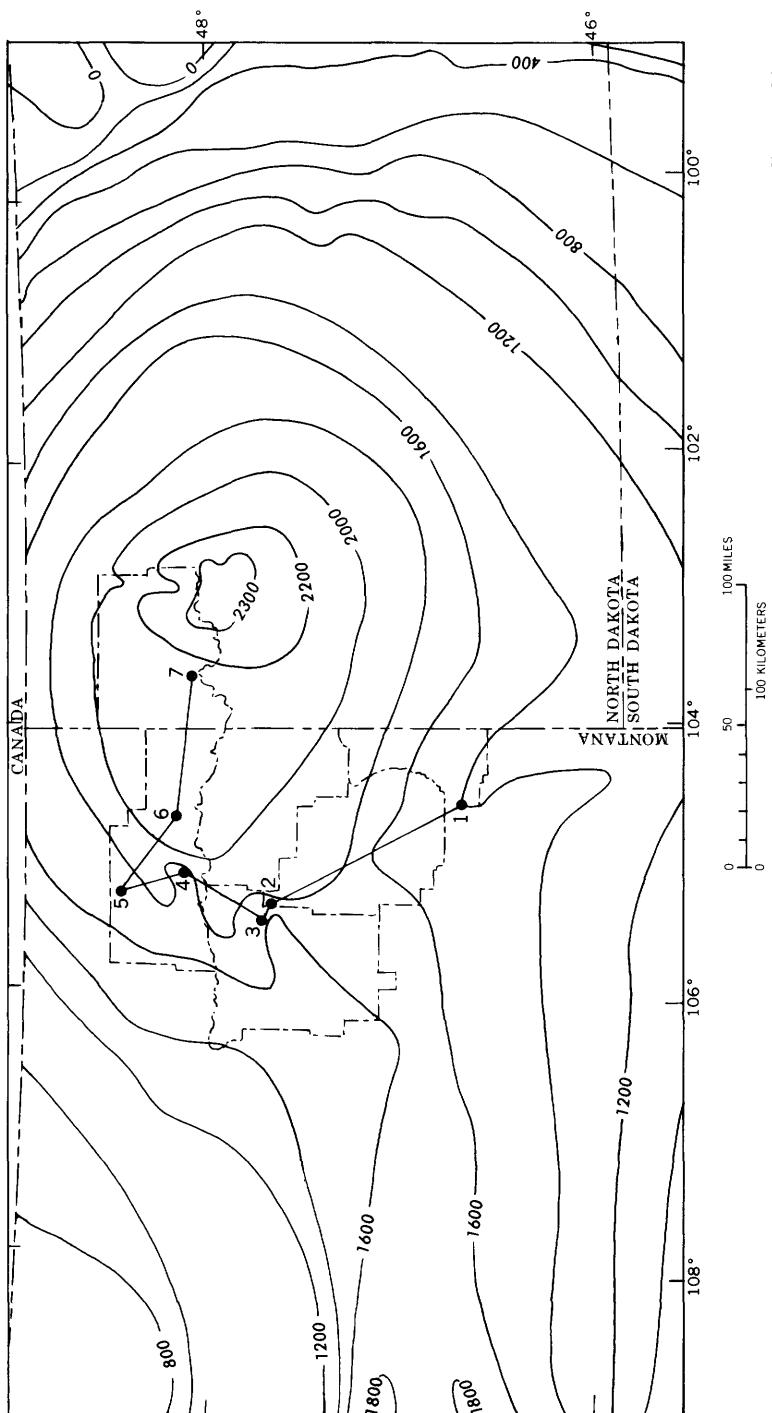


FIGURE 1.—Isopach map of Madison Group in the Williston Basin, showing locations of subsurface sections shown on figure 2 and discussed in text.
 (1) Shell Oil Co., Pine Unit No. 1, SW $\frac{1}{4}$ SW $\frac{1}{4}$ NE $\frac{1}{4}$ sec. 30, T. 12 N., R. 57 E., Wibaux County, Mont. (2) Shell Oil Co., Richey area, Northern Pacific Railroad No. 1, SW $\frac{1}{4}$ NW $\frac{1}{4}$ NW $\frac{1}{4}$ sec. 19, T. 23 N., R. 50 E., Dawson County, Mont. (3) Hodge, Smith, and Hodge Co., No. 1 Eggebrecht, CSW $\frac{1}{4}$ SE $\frac{1}{4}$ sec. 3, T. 23 N., R. 49 E., McCone County, Mont. (4) C. H. Murphy Co., East Poplar Unit No. 1, CSW $\frac{1}{4}$ NE $\frac{1}{4}$ sec. 2, T. 28 N., R. 51 E., Roosevelt County, Mont. (5) California Co., Grimm No. 1, CNE $\frac{1}{4}$ SE $\frac{1}{4}$ sec. 13, T. 32 N., R. 49 E., Roosevelt County, Mont. (6) Socony Vacuum Oil Co., Dann No. F-33, NW $\frac{1}{4}$ SE $\frac{1}{4}$ sec. 23, T. 29 N., R. 54 E., Roosevelt County, Mont. (7) Texas Co., Donahue No. 1, CSW $\frac{1}{4}$ NE $\frac{1}{4}$ sec. 23, T. 154 N., R. 100 W., Williams County, N. Dak. Isopach interval 200 ft except for 100-ft interval at center of basin. (Modified from Sandberg, 1962, fig. 15.)

CONTRIBUTIONS TO STRATIGRAPHY

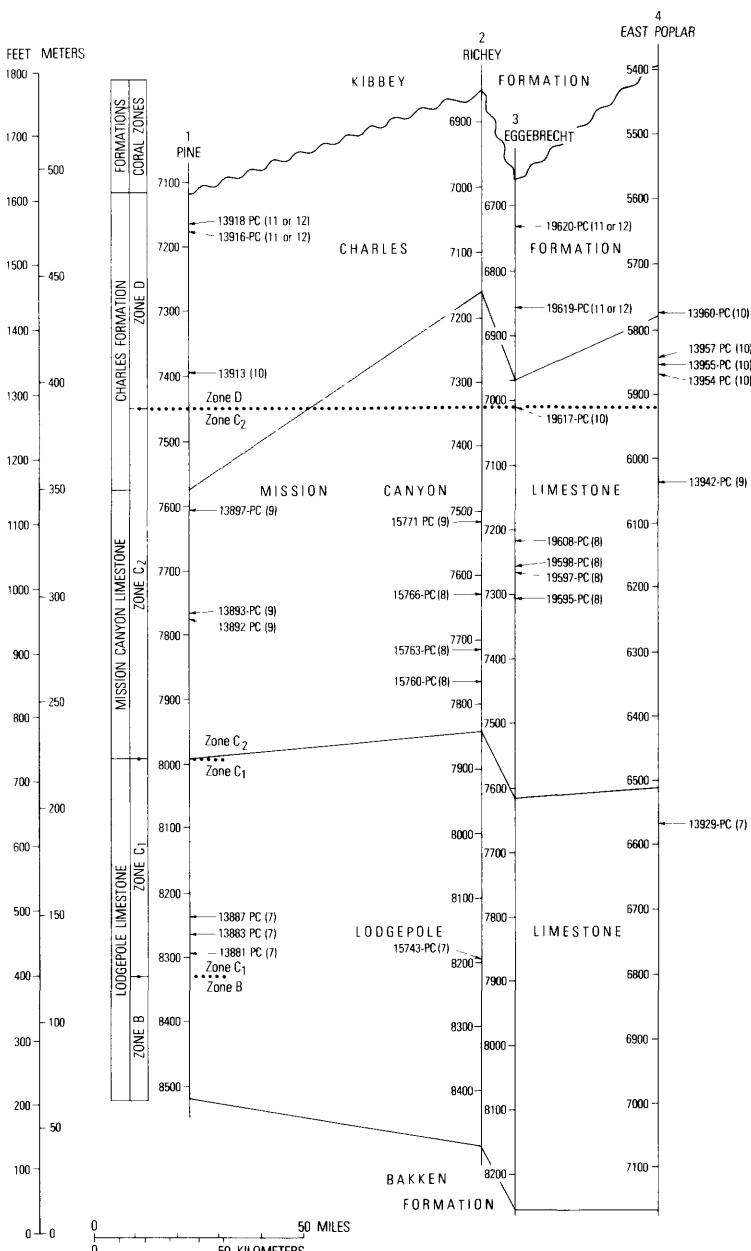


FIGURE 2.—Distribution of zonally identifiable microfossils in the Pine, Richey, Eggebrecht, and East Poplar cores. Formation boundaries (solid lines), coral zones (dotted lines), and correlation from Sando, 1978, pl. 1. Positions of microfossil samples indicated by arrows followed by USGS locality numbers and by Mamet Zone numbers (in parentheses).

MISSION CANYON LIMESTONE

The overlying Mission Canyon Limestone contains good representatives of Zones 8, 9, and 10. Zones 8 and 9 are characteristic of the lower part of the Mission Canyon elsewhere, and Zone 10 is widespread in the upper part of the formation. Zones 8 and 9 are of middle and late Osagean age (late Tournaisian of Europe) and coincide mostly with Coral Zone C₂. Zone 10 is of latest Osagean and early Meramecian age (early Viséan of Europe) and coincides with the lower part of Coral Zone D.

In the Williston Basin cores, Zone 8 is represented by an abundant and diversified microfauna and microflora. Some levels contain algal bafflestones and algal mats. *Columbiapora johnsoni* is sporadic, sometimes very abundant and sometimes absent, and encrusting *Asphaltinella* shows the same erratic pattern. Foraminifers are mostly Tournayellidae, and endothyrids are poorly represented:

- Asphaltinella* sp. (sometimes very abundant)
- Bevocastria* sp. (common in algal nodules)
- Bisphaera* sp.
- Calcisphaera* sp.
- C. laevis* Williamson
- Columbiapora johnsoni* Mamet
- Diplosphaerina* sp.
- Earlandia clavatula* (Howchin)
- E. moderata* (Malakhova)
- Issinella* sp.
- I. grandis* Reitlinger
- Kamaena* sp.
- Latiendothyra* sp.
- Nostocites* sp.
- Palaeospirolectammina tchernyshinensis* (Lipina)
- Parathurammina* sp.
- Phaeophycophyta
- Rectoseptaglomospiranella?* sp.
- Septabrunsiina parakrainica* Skipp, Holcomb, and Gutschick
- Septaglomospiranella* sp.
- Septatournayella* sp.
- cf. *Spinoendothyra?* sp. (very scarce)
- Tuberendothyra* sp.

Zone 9 is represented in the cores by an abundant and diversified microfauna and microflora. The microflora consists of *Issinella* bafflestones and *Bevocastria* nodular wackestones. Foraminifers are still dominated by Tournayellidae, along with minor Endothyridae (the *Spinoendothyra* fauna is quite sparse):

- Bevocastria* sp. (very abundant in lagoonal facies)
Calcisphaera laevis Williamson
Earlandia clavatula (Howchin)
E. elegans (Rauzer-Chernoussova)
Endothyra sp.
Eoforschia sp.
Inflatoendothyra sp.
Issinella grandis Reitlinger (very abundant in bafflestones)
Kamaena sp.
K. pirleti Mamet and Roux
Latiendothyra sp.
Palaeoberesella sp.
P. lahuseni (von Möller)
Proninella sp.
Pseudoissinella sp.
Septabrunsiina sp.
S. parakrainica Skipp, Holcomb, and Gutschick
Septaglomospiranella sp.
Septatournayella pseudocamerata (Lipina in Lebedeva)
Spinoendothyra sp.
S. spinosa (Chernysheva)

Zone 10 is represented by an abundant and diversified microfauna and microflora. Tournayellidae are slowly eliminated and replaced by Endothyridae, Globoendothyridae, and Endothyranopsidae:

- Calcisphaera laevis* (Williamson)
C. pachysphaerica (Pronina)
Dainella sp.
Diplosphaerina sp.
Earlandia clavatula (Howchin)
E. vulgaris (Rauzer-Chernoussova and Reitlinger)
cf. *Eblanaia?* sp.
Endothyra sp.
Eoparastaffella sp.
Eotuberitina sp.
cf. *Globoendothyra?* sp. (or *Eogloboendothyra?* sp.)
Globoendothyra of the group *G. baileyi* (Hall)
Inflatoendothyra sp.
Issinella sp.
Parathurammina sp.
Priscella sp.
Quasipolyderma sp.
Radiosphaerina sp.
Septatournayella(?) henbesti Skipp, Holcomb, and Gutschick

CHARLES FORMATION

The Charles Formation, uppermost formation in the Madison Group, is characterized by microfossil assemblages of Zones 10 and 11 (possibly ranging into Zone 12). In the outcrop area of the Mississippian west of the Williston Basin, these zones are found in the upper part of the Mission Canyon Limestone (Sando and others, 1969). Zones 10 and 11 are of early Meramecian age (early Viséan of Europe) and are ordinarily the youngest zones found in the Madison Group beneath a regional disconformity overlain by the Big Snowy Group in Montana and the Amsden Formation in Wyoming (Sando and others, 1969). The only exception to this rule is in the Bighorn Mountains of Wyoming and Montana, where rare occurrences of Zone 12 are known (Sando and Mamet, 1974). The rest of Meramecian time (Zones 13 through 15) is represented by an hiatus wherever the Madison is directly overlain by the Chesterian (Big Snowy Group and lower part of the Amsden Formation) (Sando and others, 1975, fig. 13). The occurrence of Zone 11 (possibly ranging into 12) microfossils near the top of the Charles Formation in the Pine and Eggebrecht wells corroborates the interpretation based on corals (Sando, 1978, p. 236)—that the top of the Charles Formation in the subsurface is bounded by the same disconformity that is so well documented in outcrops of the Madison to the west.

Zone 11 (possibly ranging into 12) microflora is abundant and diversified in a lagoonal facies. Among the foraminifers, the endothyrids, globoendothyrids, and endothyranopsids dominate; the *Eoendothyranopsis* fauna is poorly represented:

Asphaltinella sp. (abundant in algal mats)

Brunisia sp.

Calcisphaera laevis (Williamson)

C. pachysphaerica (Pronina)

Dainella sp.

Erlandia of the group *E. clavatula* (Howchin)

E. vulgaris (Rauzer-Chernoussova and Reitlinger)

Eblanaia sp.

Endothyra sp.

cf. *Eoendothyranopsis?* sp.

Eoendothyranopsis of the group *E. spiroides* (Zeller)

E. spiroides (Zeller)

Eoforschia sp.

E. moelleri (Malakhova in Dain)

Eoparastaffella sp.

Globoendothyra bridgensis Skipp

cf. *Inflatoendothyra?* sp.

Issinella sp.

Kamaena sp.
Nostocites sp.
Orthriosiphonoides sp.
Palaeoberesella sp.
Paracalligelloides sp.
Parathurammina sp.
Planoendothyra? sp.
Priscella sp.
Pseudoissinella sp.
Radiosphaerina sp.
cf. *Septabrunsiina?* sp.
Septatournayella(?) henbesti Skipp, Holcomb, and Gutschick
Skippella sp.
Stacheiinae

ENVIRONMENTAL INTERPRETATION

The Williston Basin foraminifers belong to the North America Realm (Mamet, 1977; Mamet and Skipp, 1979) but are far less abundant and diversified than in the Lodgepole and Mission Canyon Limestones of the outcrop area to the west (see microfossil lists in Sando and others, 1969). In the Williston Basin, most of the foraminiferal microfauna is observed in limestones that show signs of restricted circulation (sulfates are present in the Lodgepole and are widespread in the Mission Canyon and Charles). Pseudomorphs of sulfates are common in algal-rich lagoonal phases; this environment is basically unfavorable to foraminifers, which prefer open-marine conditions. On the other hand, the algal microflora is more abundant and diversified, including palaeoberesellids, issinellids, and bevoastridiids. Dasyclads are usually restricted to *Columbiapora*. *Asphaltinella* is quite common and is accompanied by stacheiins in the upper part of the Charles Formation.

REGIONAL CORRELATION

Figure 3 shows correlation of the Madison Group in the Pine Well core with type sections of the Madison and its component stratigraphic units in the outcrop area to the west in Montana. Figure 4 shows locations of the sections.

The correlation diagram (fig. 3) was originally based on coral zones recognized by Sando (1978). Occurrences of zonally determined microfossil collections have been added to the diagram, and microfossil zone boundaries have been drawn. Three of the four microfossil-zone boundaries coincide with coral-zone boundaries. Thus, the microfossil correlations confirm the correlations made on coral zones and support the

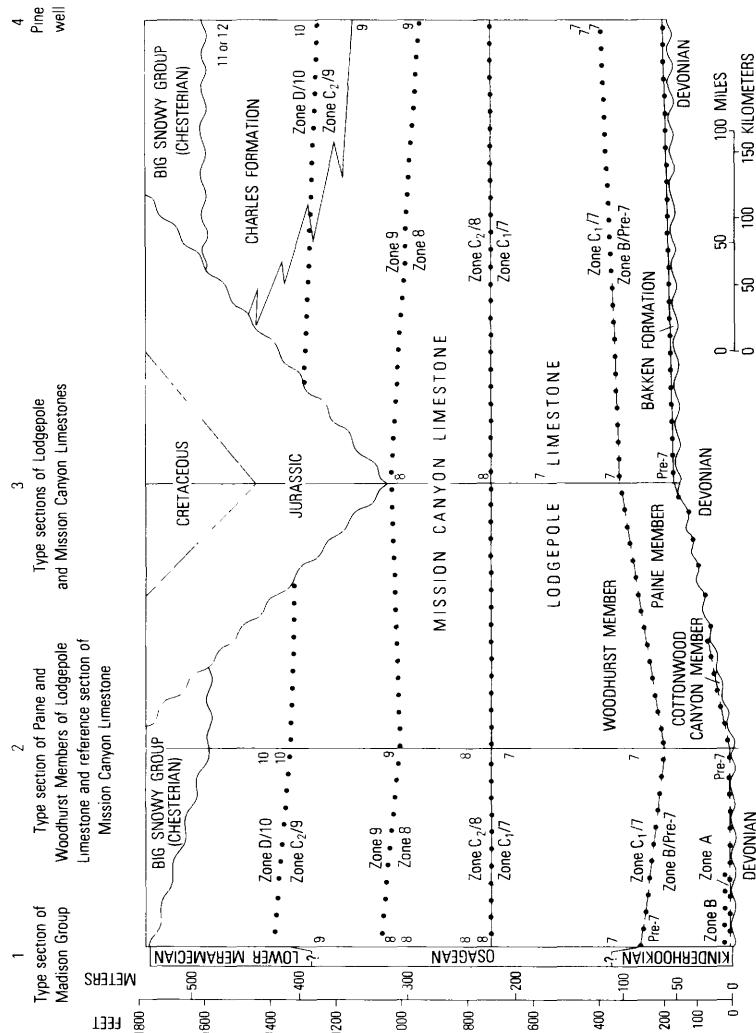


FIGURE 3.—Correlation of Madison Group in Shell Pine Unit No. 1 well (4) with (1) Logan section, (2) composite of Monarch-US 89 and Dry Fork sections, and (3) composite of Little Chief Canyon and Mission Canyon sections. Figure 4 shows locations of sections. Formation boundaries indicated by solid lines (wavy where disconformable), member boundaries indicated by dashed lines, fossil-zone boundaries indicated by dots. Positions of microfossil collections indicated by zone numbers adjacent to sections. (See Sando and Dutro (1974) for descriptions of the type sections and Sando (1960) for log of Shell Pine well.)

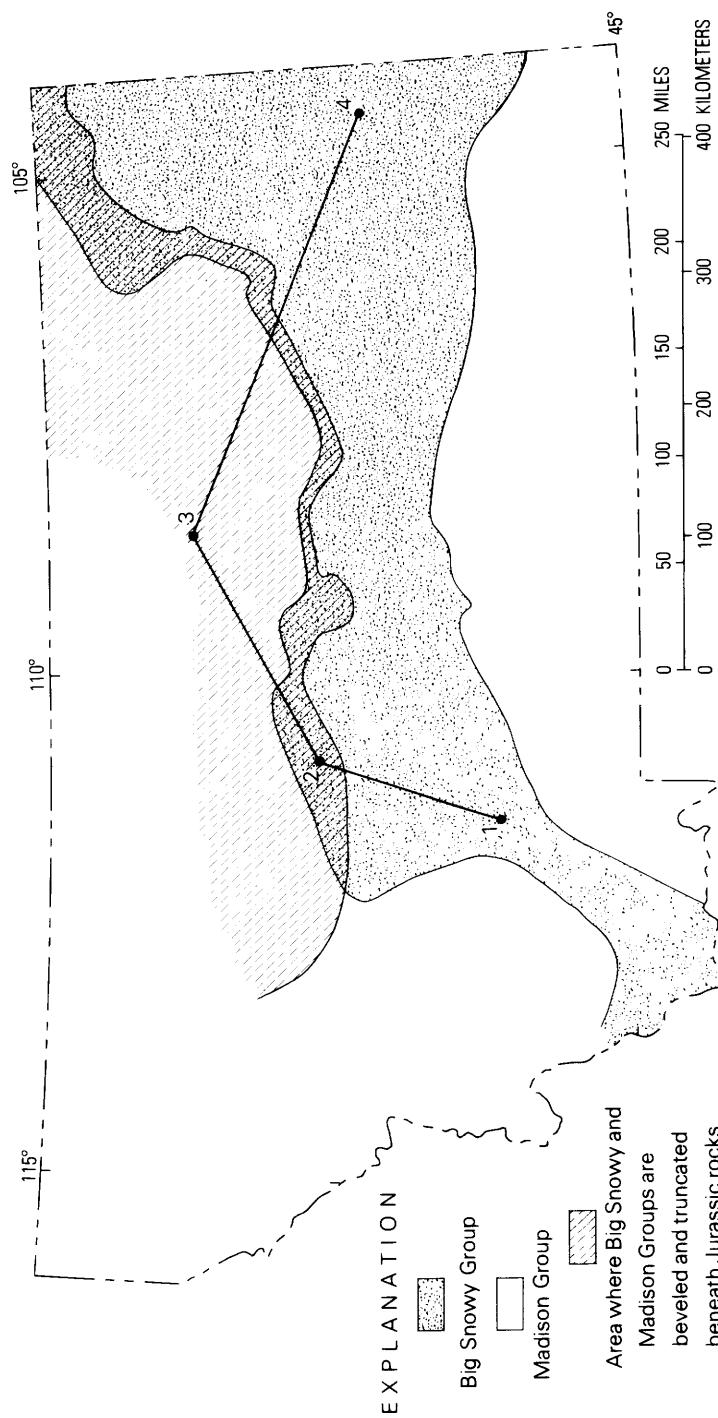


FIGURE 4.—Paleogeologic map of Montana showing distribution of Big Snowy and Madison Groups and locations of stratigraphic sections shown on figure 3. (1) Logan section ($SE \frac{1}{4} SW \frac{1}{4}$ sec. 25, T. 2 N., R. 2 E., Gallatin County). (2) Composite of Monarch-U.S. 89 and Dry Fork sections ($SE \frac{1}{4} NE \frac{1}{4}$ sec. 27, T. 16 N., R. 7 E. and SE corner sec. 35, T. 16 N., R. 7 E., Cascade County). (3) Composite of Little Chief Canyon and Mission Canyon sections ($NE \frac{1}{4} NW \frac{1}{4}$ sec. 30, T. 26 N., R. 25 E., and $SE \frac{1}{4} NW \frac{1}{4}$ sec. 32, T. 26 N., R. 24 E., Blaine County). (4) Shell Oil Co., Pine Unit No. 1 well ($SW \frac{1}{4} SW \frac{1}{4} NE \frac{1}{4}$ sec. 30, T. 12 N., R. 57 E., Wibaux County). (Modified from Maughan and Roberts, 1967, pl. 3.)

interpretation that the Madison Group in the subsurface of the Williston Basin is the same age as the Madison Group of outcrop.

CONCLUSIONS

Foraminifers and algae in the Madison Group of the Williston Basin are not as abundant or diversified as they are in other areas where the Madison microfossils have been studied, because of less favorable environmental conditions during Mississippian time in the Williston Basin. Despite these limitations, the presence of Zones 10 and 11 (possibly ranging into 12) assemblages near the top of the Charles Formation and the absence of assemblages from younger Meramecian zones indicates that the Charles is no younger than early Meramecian and that it is bounded above by the same regional disconformity that characterizes the top of the Madison Group in outcrop areas outside the Williston Basin.

REFERENCES CITED

- Gardner, L. S., 1959, Revision of Big Snowy Group in central Montana: American Association of Petroleum Geologists Bulletin, v. 43, no. 2, p. 329-349.
- Hadley, H. D., 1950, The Charles problem, in Billings Geological Society, 1st Annual Field Conference [Guidebook]: p. 44-46.
- Mamet, B. L., 1977, Foraminiferal zonation of the lower Carboniferous; methods and stratigraphic implications, in Kauffman, E. G., and Hazel, J. E., eds., Concepts and methods of biostratigraphy: Stroudsburg, Pa., Dowden, Hutchinson, & Ross, p. 445-462.
- Mamet, B. L., and Skipp, Betty, 1970a, Lower Carboniferous calcareous Foraminifera; preliminary zonation and stratigraphic implications for the Mississippian of North America: International Congress of Carboniferous Stratigraphy and Geology, 6th, Sheffield, 1967, Comptes Rendus, v. 3, p. 1129-1146.
- , 1970b, Preliminary foraminiferal correlations of early Carboniferous strata in the North American Cordillera, in Colloque sur la Stratigraphie du Carbonifère: Liège Université, Congrès et Colloques, v. 55, p. 327-348.
- , 1979, Lower Carboniferous Foraminifera; paleogeographical implications: International Congress of Carboniferous Stratigraphy and Geology, 8th, Moscow, 1975, Comptes Rendus, v. 2, p. 48-66.
- Maughan, E. K., and Roberts, A. E., 1967, Big Snowy and Amsden Groups and the Mississippian-Pennsylvanian boundary in Montana: U.S. Geological Survey Professional Paper 554-B, 27 p.
- Nordquist, J. W., 1953, Mississippian stratigraphy of northern Montana, in Billings Geological Society, 4th Annual Field Conference, Little Rocky Mountains-Montana, southwestern Saskatchewan, Sept. 1953, Guidebook: p. 68-82.
- Perry, E. S., and Sloss, L. L., 1943, Big Snowy Group, lithology and correlation in the northern Great Plains: American Association of Petroleum Geologists Bulletin, v. 27, no. 10, p. 1287-1304.
- Roberts, A. E., 1966, Stratigraphy of Madison Group near Livingston, Montana, and discussion of karst and solution-breccia features: U.S. Geological Survey Professional Paper 526-B, 23 p.
- Sandberg, C. A., 1962, Geology of the Williston Basin, North Dakota, Montana, and South Dakota, with reference to subsurface disposal of radioactive wastes: U.S. Geological Survey open-file report, 148 p.

- Sando, W. J., 1960, Corals from well cores of Madison Group, Williston Basin: U.S. Geological Survey Bulletin 1071-F, p. 157-190 [1961].
- , 1978, Coral zones and problems of Mississippian stratigraphy in the Williston Basin, in Montana Geological Society, 24th Annual Conference [Sept. 24-27, 1978], 1978 Williston Basin Symposium—The economic geology of the Williston Basin: [Billings, Mont.], p. 231-237.
- Sando, W. J., and Dutro, J. T., Jr., 1974, Type sections of the Madison Group (Mississippian) and its subdivisions in Montana: U.S. Geological Survey Professional Paper 842, 22 p.
- Sando, W. J., Gordon, Mackenzie, Jr., and Dutro, J. T., Jr., 1975, Stratigraphy and geologic history of the Amsden Formation (Mississippian and Pennsylvanian) of Wyoming: U.S. Geological Survey Professional Paper 848-A, 83 p.
- Sando, W. J., and Mamet, B. L., 1974, New evidence on the age of the top of the Madison Limestone (Mississippian), Bighorn Mountains, Wyoming and Montana: U.S. Geological Survey Journal of Research, v. 2, no. 5, p. 619-624.
- Sando, W. J., Mamet, B. L., and Dutro, J. T., Jr., 1969, Carboniferous megafaunal and microfaunal zonation in the northern Cordillera of the United States: U.S. Geological Survey Professional Paper 613-E, 29 p.
- Sloss, L. L., 1952, Introduction to the Mississippian of the Williston Basin, in Billings Geological Society, 3d Annual Field Conference, Black Hills-Williston Basin, Sept. 1952, Guidebook: p. 65-69.

